



第四章:

中纬度的经向环流系统(IV)

- *Ferrel cell, baroclinic eddies
and the westerly jet*

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E-P flux, TEM and Residual Circulation - Summary

Review

- E-P flux: $\mathcal{F} \equiv -[u^*v^*] \mathbf{j} + f \frac{[v^*\theta^*]}{\partial\theta_s/\partial p} \mathbf{k}$

- In a **steady**, **adiabatic** and **frictionless** flow:

$$[v] = \frac{1}{f} \frac{\partial}{\partial y} ([u^*v^*]) \quad [\omega] = -\frac{\partial}{\partial y} \left(\frac{[\theta^*v^*]}{\partial\theta_s/\partial p} \right) \quad \nabla \cdot \mathcal{F} = 0$$

- Residual mean circulations:

$$[\tilde{\omega}] = [\omega] + \frac{\partial}{\partial y} \left(\frac{[v^*\theta^*]}{\partial\theta_s/\partial p} \right), \quad [\tilde{v}] = [v] - \frac{\partial}{\partial p} \left(\frac{[v^*\theta^*]}{\partial\theta_s/\partial p} \right)$$

- TEM equations: $\frac{\partial[u]}{\partial t} = f[\tilde{v}] + \nabla \cdot \mathcal{F} + [F_x], \quad \frac{\partial[\theta]}{\partial t} = -[\tilde{\omega}] \frac{\partial\theta_s}{\partial p} + \left(\frac{p_o}{p} \right)^{R/c_p} \frac{[Q]}{c_p}$



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#1

$$[v^*q^*] = -\frac{\partial}{\partial y} [u^*v^*] + f_o \frac{\partial}{\partial p} \frac{[v^*\theta^*]}{\partial\theta_s/\partial p} = \nabla \cdot \mathcal{F}$$

#2

$$\frac{\partial \mathcal{A}}{\partial t} + \nabla \cdot \mathcal{F} = 0$$

#3

$$\vec{\mathcal{F}} = \mathbf{c}_g \mathcal{A}$$



$$\frac{\partial \mathcal{A}}{\partial t} + \nabla \cdot (\mathcal{A} \mathbf{c}_g) = 0$$

- Residual mean circulations:

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Outline



- Observations
- The Ferrel Cell
- Baroclinic eddies
 - Review: baroclinic instability and baroclinic eddy life cycle
 - Eddy-mean flow interaction, E-P flux
 - Transformed Eulerian Mean equations
- Eddy-driven jet
- The energy cycle



Baroclinic eddy life cycle

- An E-P flux view

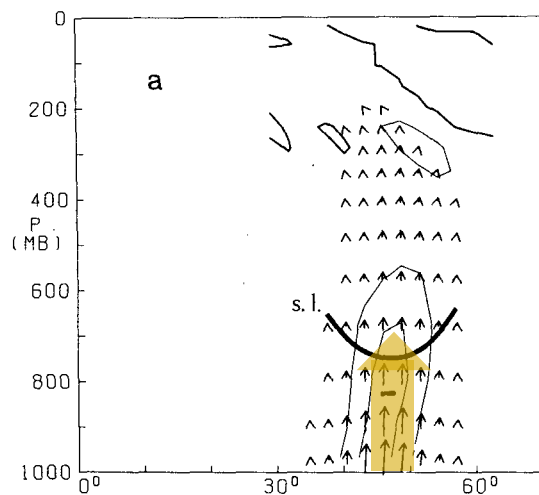
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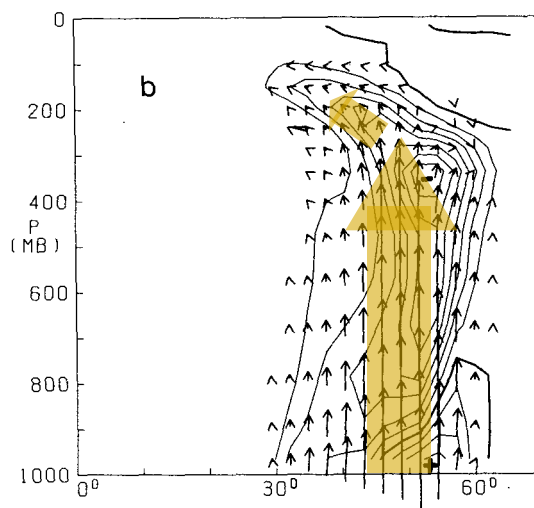
Eddies: generate at lower level,
propagate **upwards** and **away** from the
eddy source region

Review

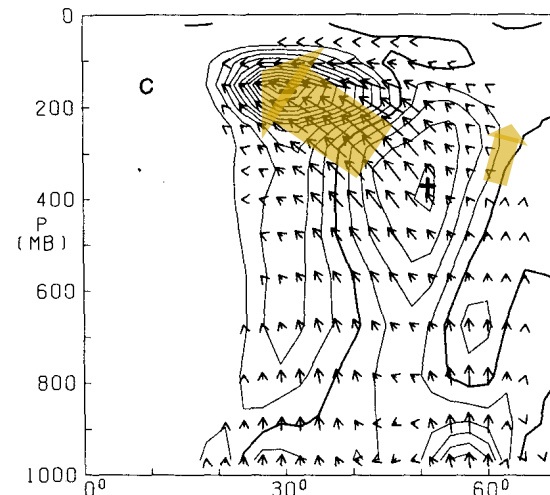
Numerical simulations
Simmons and Hoskins,
1978, JAS



TOTAL E-P FLUX DIVERGENCE
DAY .00



TOTAL E-P FLUX DIVERGENCE
DAY 5.00



TOTAL E-P FLUX DIVERGENCE
DAY 8.00



E-P flux

- The westerly jet

Review

Wave energies.
propagate **upwards** and
away from the center of
the jet

$$\mathcal{F} \equiv -[u^* v^*] \mathbf{j} + f \frac{[v^* \theta^*]}{\partial \theta_s / \partial p} \mathbf{k}$$

$$\vec{\mathcal{F}} = \mathbf{c}_g \mathcal{A}$$

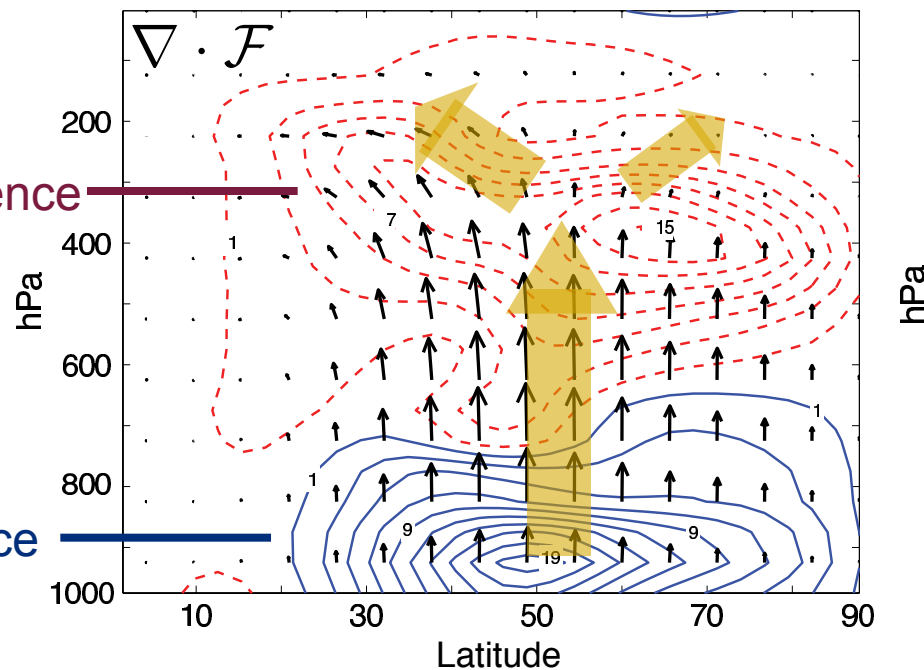
$$\frac{\partial[u]}{\partial t} = f[\tilde{v}] + \nabla \cdot \mathcal{F} + [F_x]$$

In the vertical direction:

Accelerating the lower jet
decelerating the upper jet
reduce the vertical shear of U

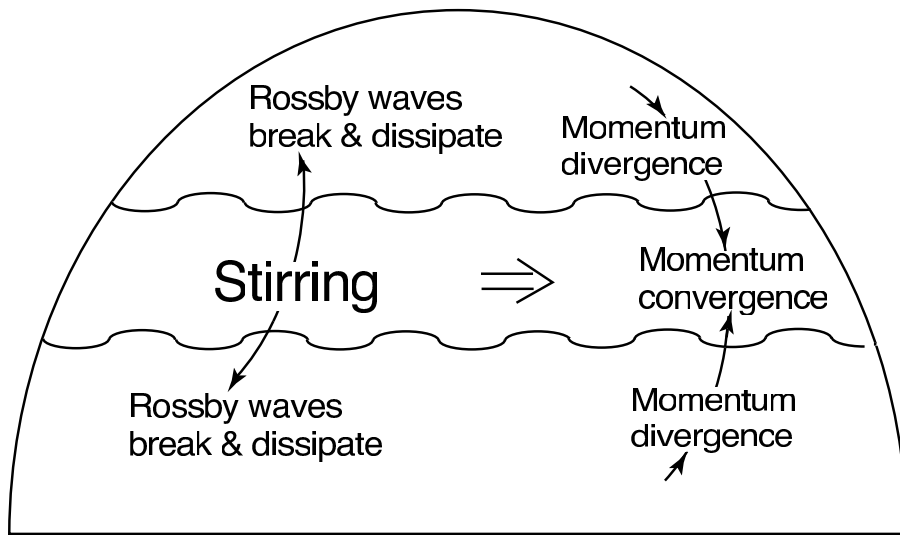
Convergence

Divergence





$$\mathcal{F} \equiv -[u^* v^*] \mathbf{j} + f \frac{[v^* \theta^*]}{\partial \theta_s / \partial p} \mathbf{k}$$



$$\frac{\partial}{\partial t} \langle [u] \rangle = -\frac{\partial}{\partial y} \langle [u^* v^*] \rangle - r[u_{\text{surf}}]$$

$\langle \rangle$ means vertical average



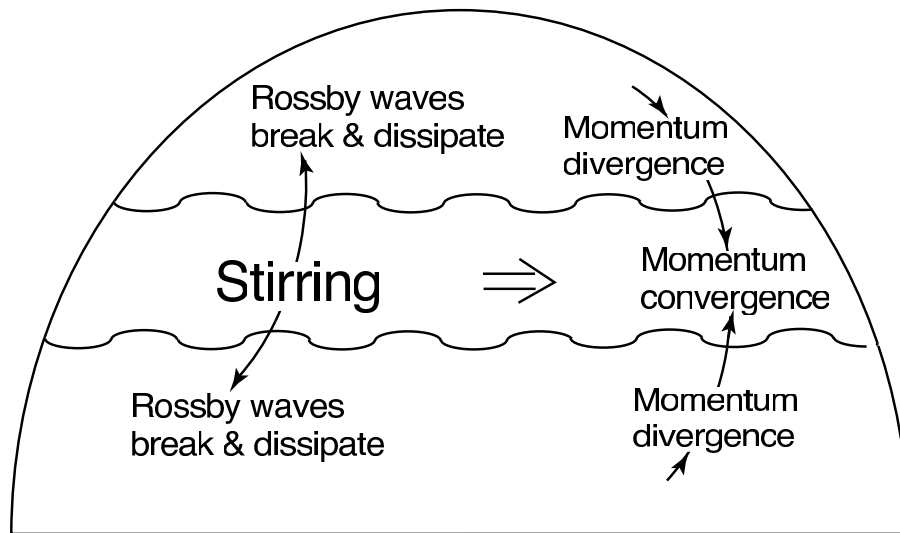
Wave energies.
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Wave energies.
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In equilibrium: $\vec{\mathcal{F}} = \vec{c}_g \mathcal{A}$

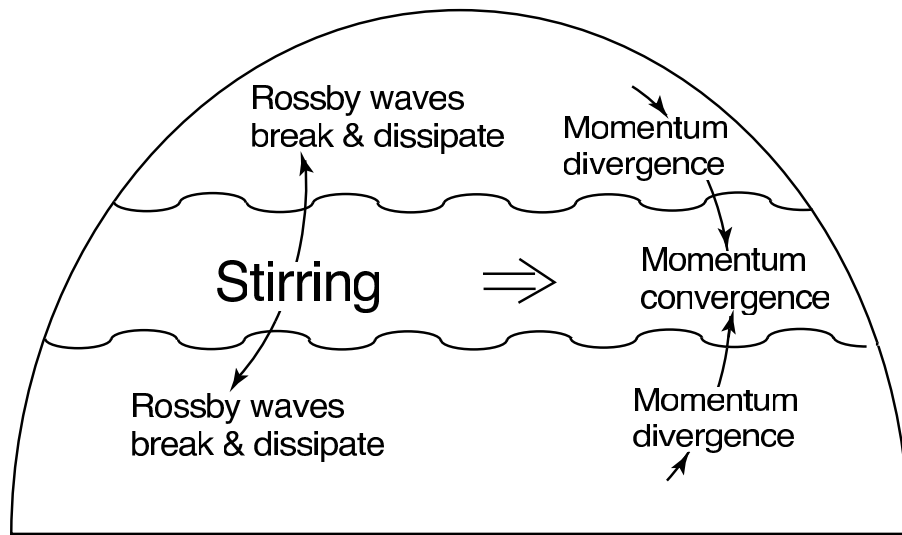
$$r[u_{\text{surf}}] \sim -\frac{\partial}{\partial y} \langle [u^* v^*] \rangle$$

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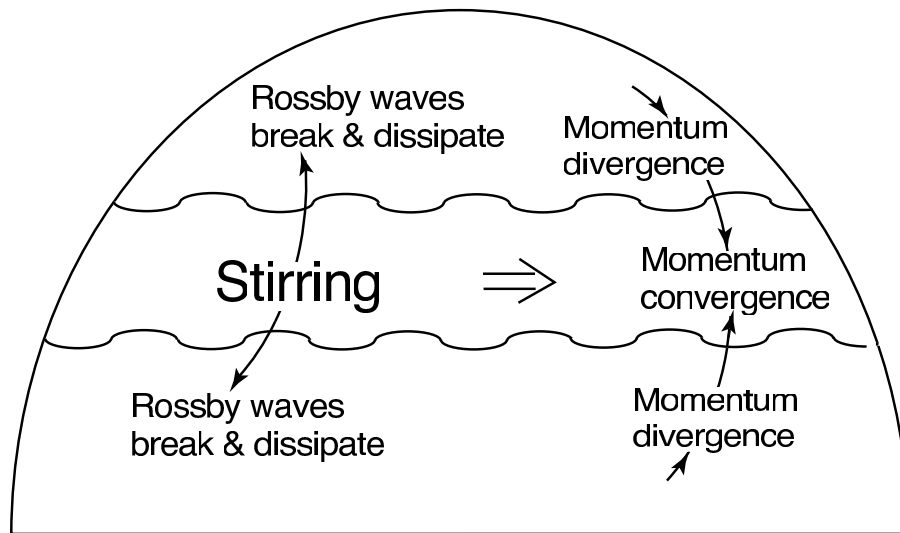
There MUST be **surface westerlies** at midlatitudes.



Eddy-driven jet:

- the momentum budget

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Review

Wave energies propagate **upwards** and **away** from the center of the jet

In equilibrium:

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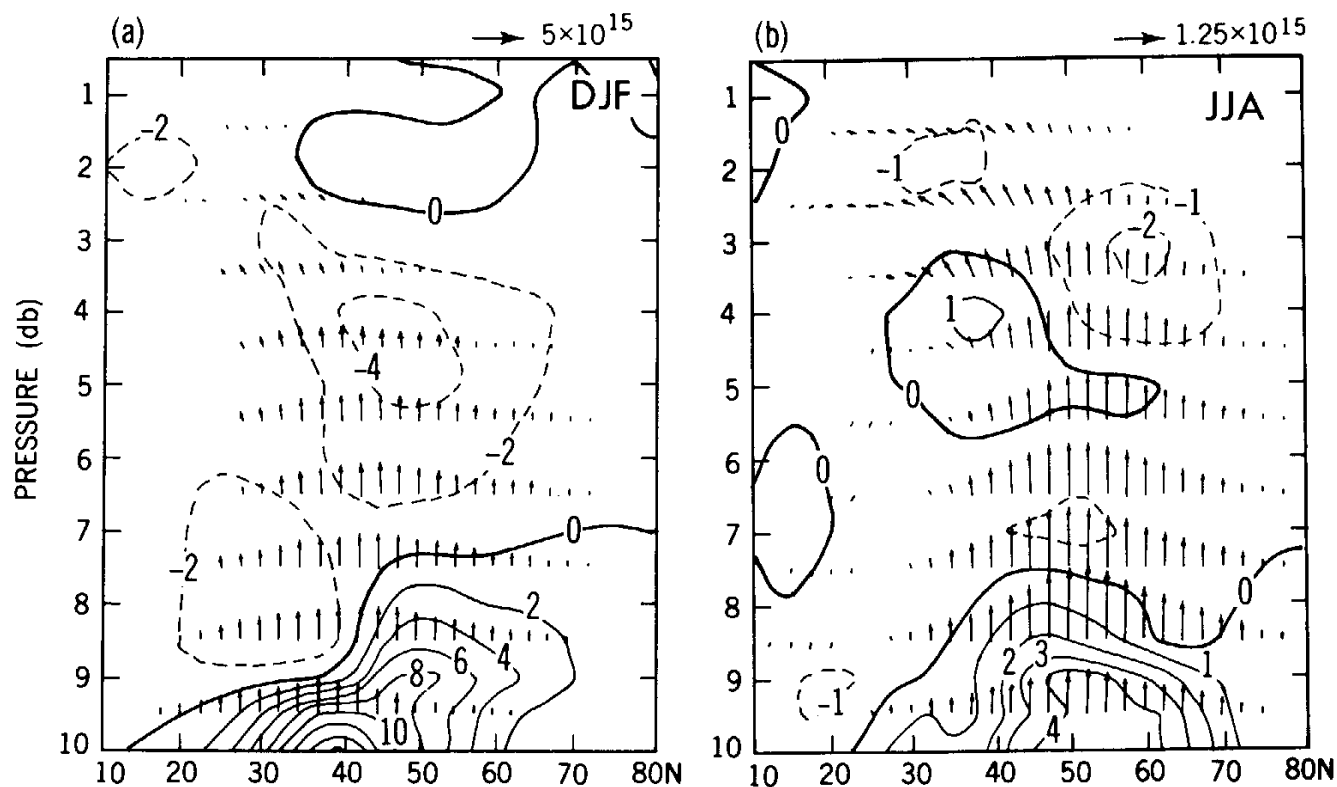
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E-P FLUX TRANSIENT EDDIES





E-P flux

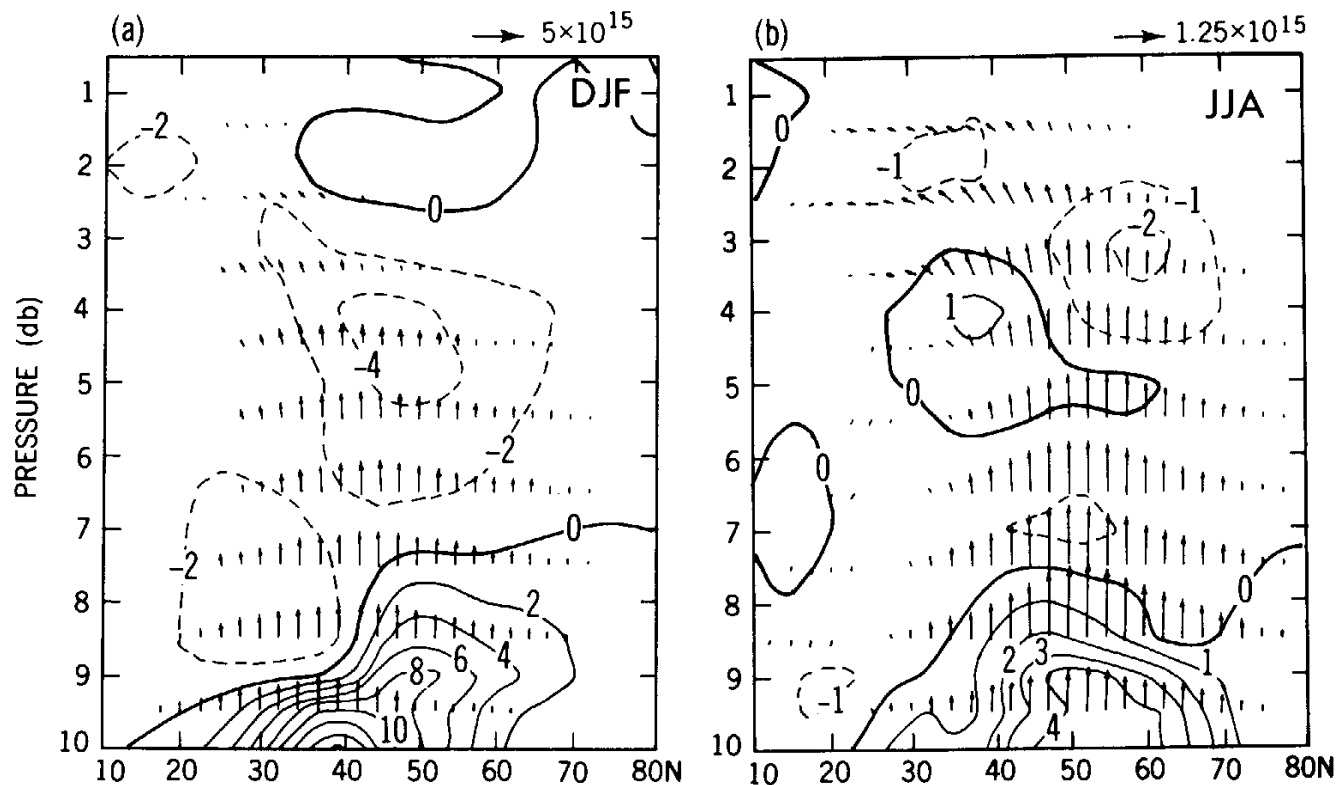
- in the real atmosphere



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E-P FLUX TRANSIENT EDDIES





E-P flux

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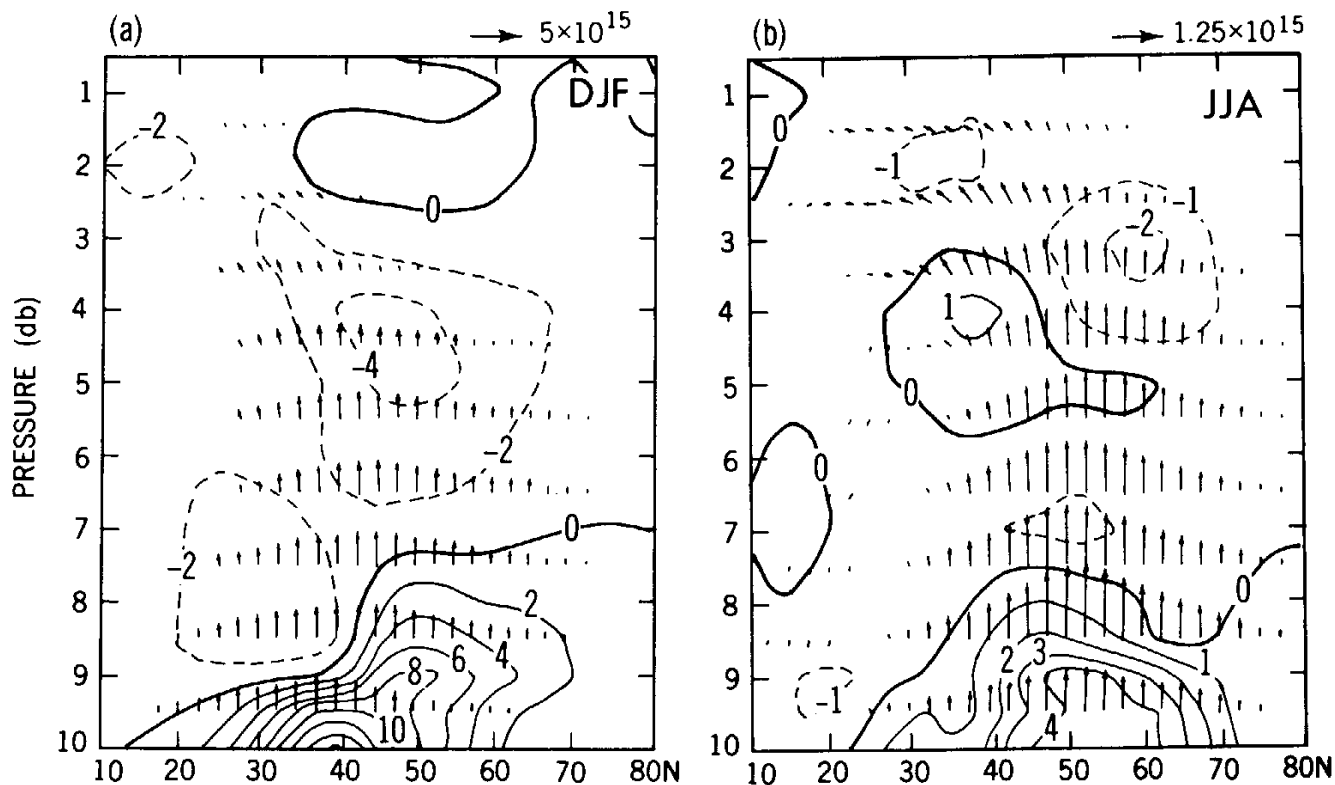


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Vertical component
is dominant.

E-P FLUX TRANSIENT EDDIES





E-P flux

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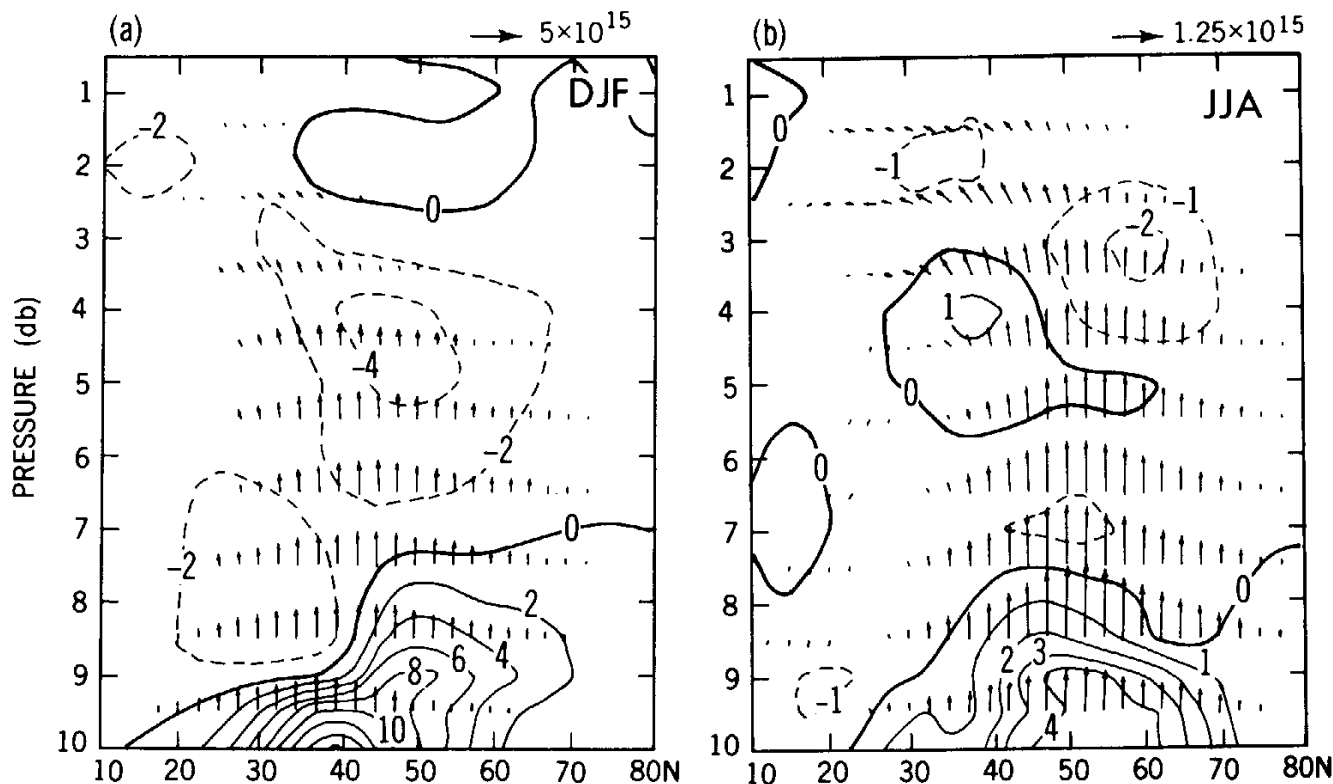
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Vertical component
is dominant.

EP divergence in
the lower layers;
convergence in the
upper layers.

E-P FLUX TRANSIENT EDDIES





E-P flux and the eddy-driven jet

-summary



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E-P flux and the eddy-driven jet -summary



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- Numerical results and observations: eddies **generate** in the lower level, propagate **upwards** and **away** from the eddy source region.

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- **Accelerating** the lower jet, **decelerating** the upper jet, reduce the vertical shear of U
- **Momentum budget** indicates that there MUST be **surface westerlies** in the eddy source latitude.



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Energy cycles

in the baroclinic eddy-mean flow interactions



- Basic forms of energy:



Energy cycles

in the baroclinic eddy-mean flow interactions



■ Basic forms of energy:

- Kinetic energy (动能):
$$K = \frac{1}{2}(u^2 + v^2 + w^2) \approx \frac{1}{2}(u^2 + v^2)$$



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- Kinetic energy (动能): $K = \frac{1}{2}(u^2 + v^2 + w^2) \approx \frac{1}{2}(u^2 + v^2)$
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■ Total energy:

$$E = I + \Phi + LH + K$$



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■ Total potential energy:

$$\int_0^\infty \rho(I + \Phi)dz = \frac{1}{g} \int_0^{p_s} (c_v T + RT)dp = \frac{1}{g} \int_0^{p_s} c_p T dp$$



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Energy cycles

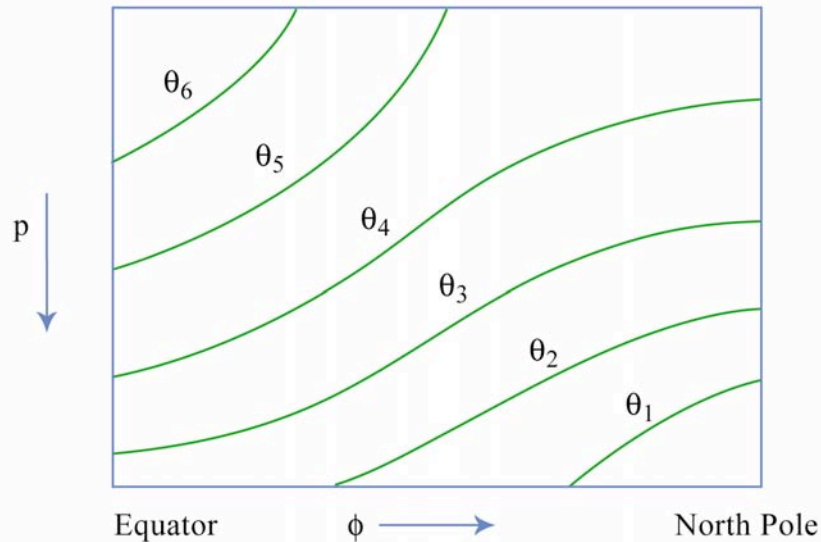
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From Stone's class notes



Energy cycles

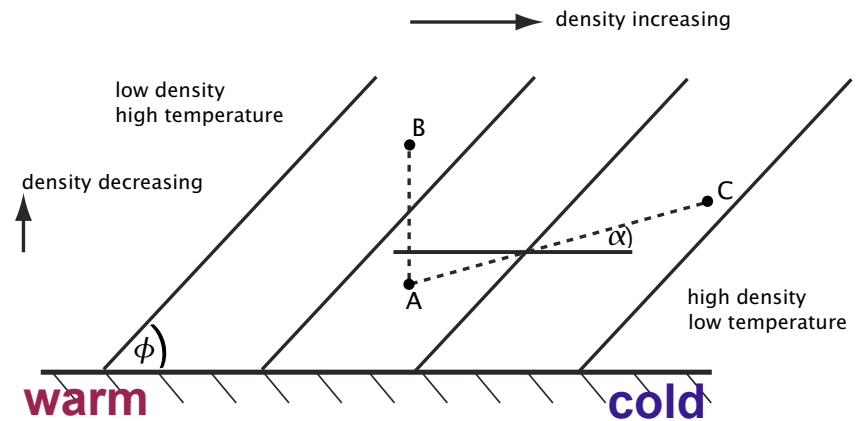
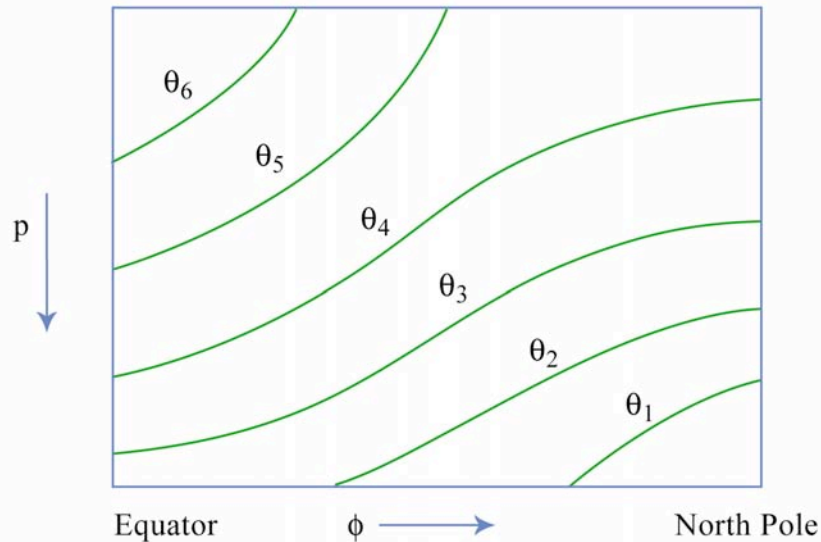
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Energy cycles

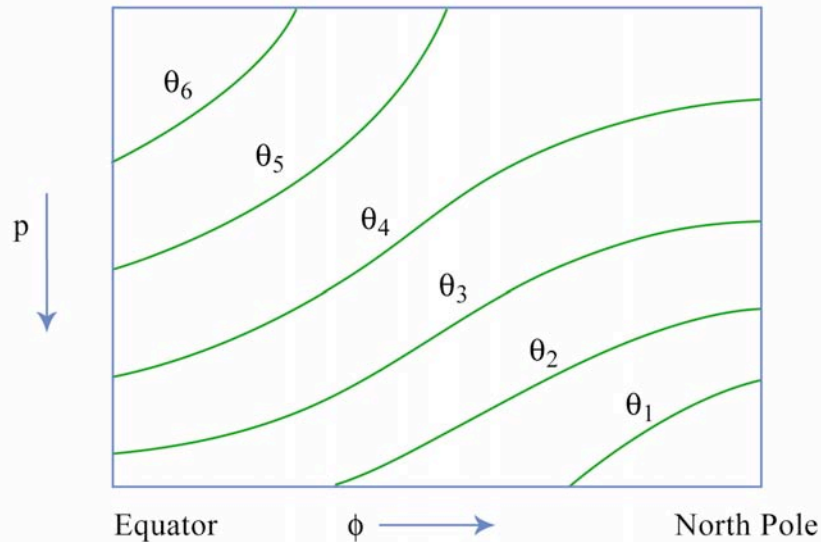
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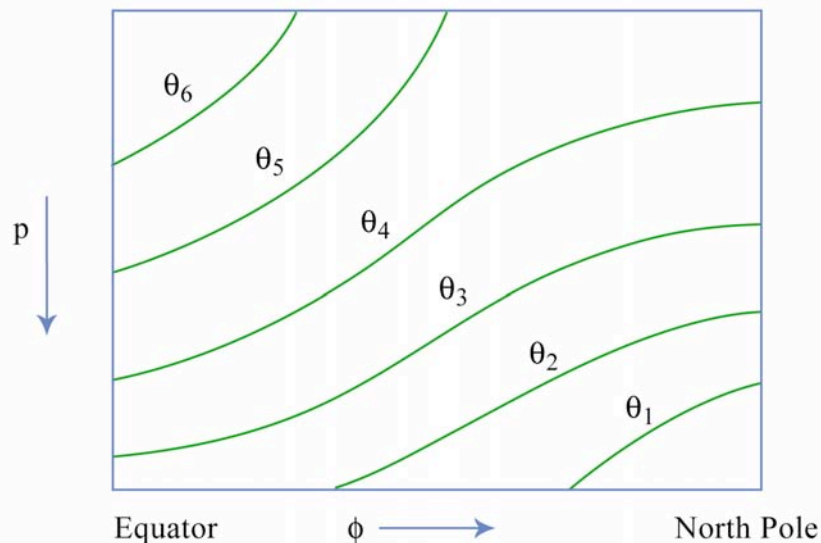
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Energy cycles

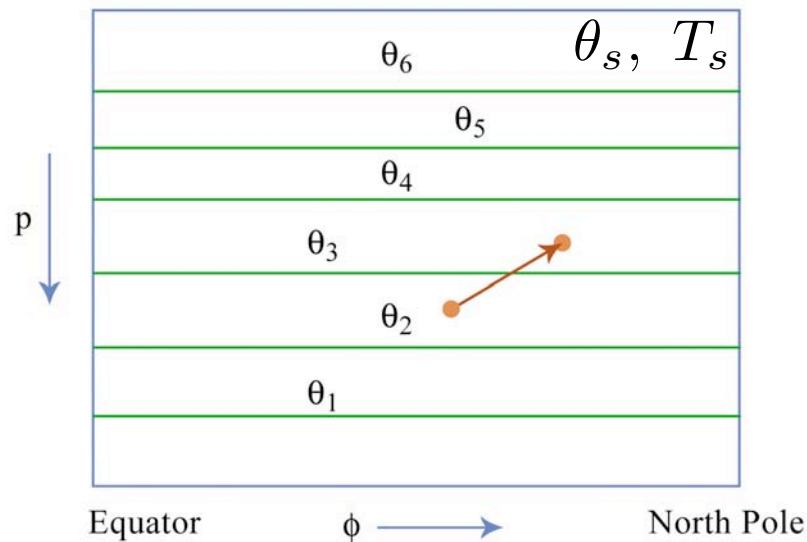
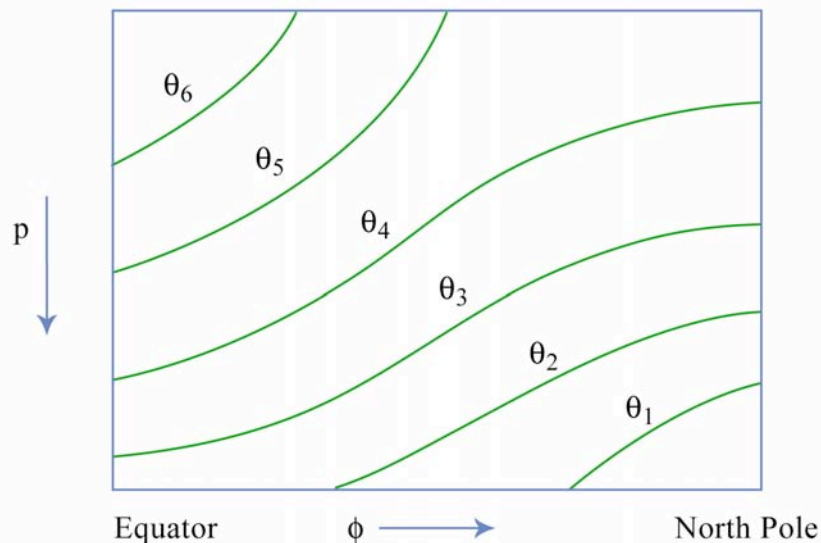
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Energy cycles

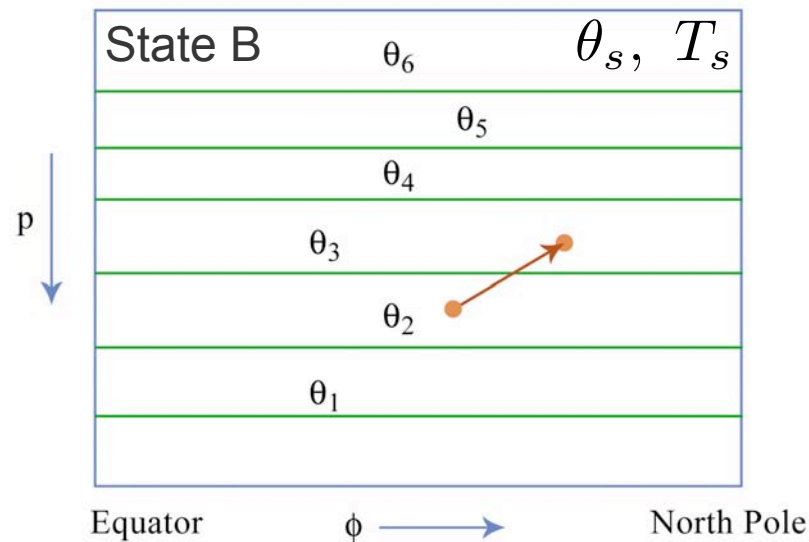
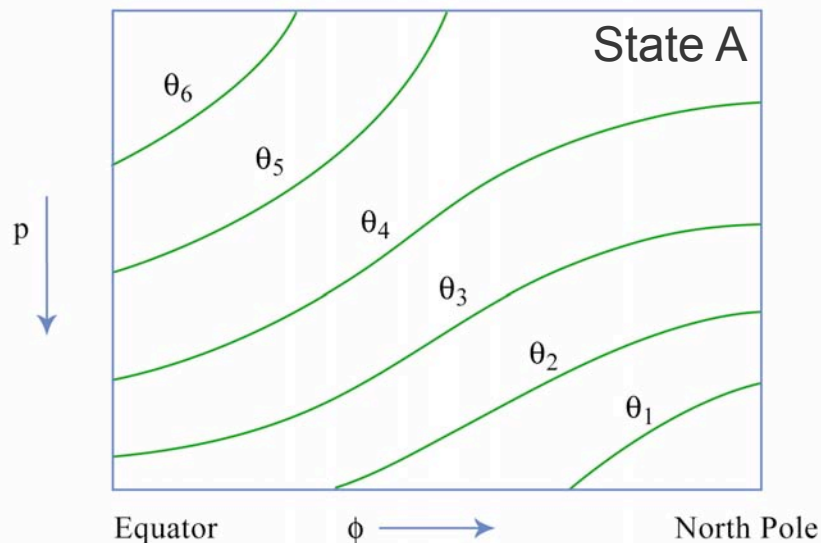
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Energy cycles

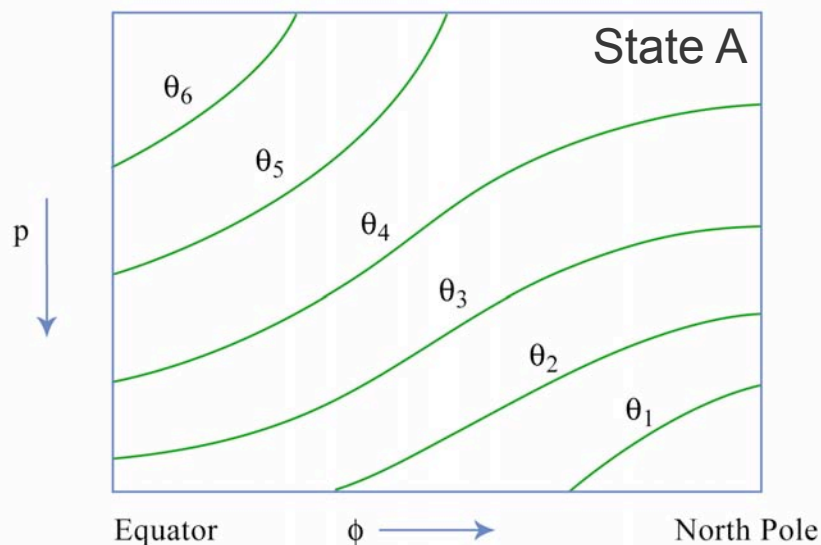
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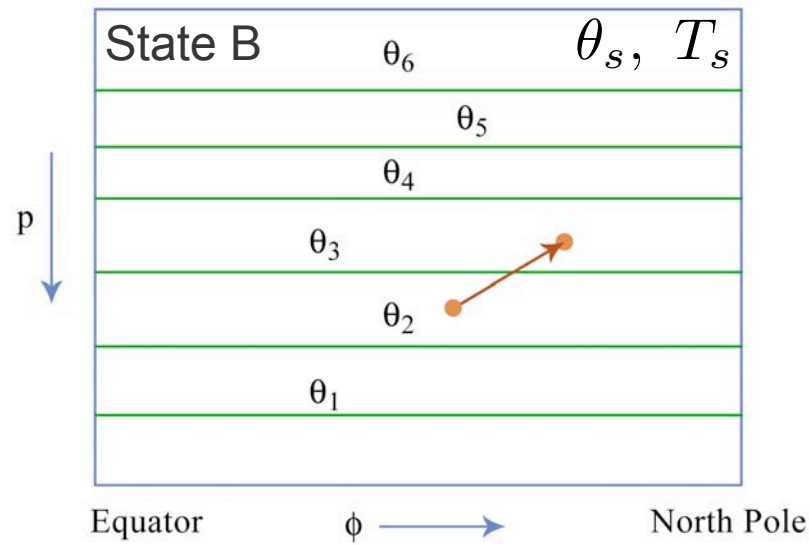
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—



= Available potential energy



Energy cycles

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State A — State B = **Available potential energy**



Energy cycles

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- Available potential energy (有效位能):

$$P = \frac{1}{2} \int_0^{p_s} \frac{T_s}{\gamma_d - \gamma_s} \left(\frac{T - T_s}{T_s} \right)^2 dp = \frac{c_p}{2g} \int_0^{p_s} \Gamma (T - T_s)^2 dp$$

$$\Gamma = -\frac{R}{c_p p} \left(\frac{p_s}{p} \right)^{\frac{R}{c_p}} \left(\frac{\partial \theta_s}{\partial p} \right)^{-1}$$

$$= (\gamma_d / T_s) (\gamma_d - \gamma_s)^{-1}$$



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From the “approximate” expression
of Lorenz (1955)

$$\begin{aligned} \Gamma &= -\frac{R}{c_p p} \left(\frac{p_s}{p} \right)^{\frac{R}{c_p}} \left(\frac{\partial \theta_s}{\partial p} \right)^{-1} \\ &= (\gamma_d / T_s) (\gamma_d - \gamma_s)^{-1} \end{aligned}$$



Energy cycles

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■ Basic forms of energy:

■ Kinetic energy (动能):
$$K = \frac{1}{2}(u^2 + v^2 + w^2) \approx \frac{1}{2}(u^2 + v^2)$$

■ Available potential energy (有效位能):

$$P = \frac{1}{2} \int_0^{p_s} \frac{T_s}{\gamma_d - \gamma_s} \left(\frac{T - T_s}{T_s} \right)^2 dp = \frac{c_p}{2g} \int_0^{p_s} \Gamma (T - T_s)^2 dp$$

■ Tendency equations:



Energy cycles

in the baroclinic eddy-mean flow interactions



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Q - diabatic heating



Energy cycles

in the baroclinic eddy-mean flow interactions



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Energy cycles

in the baroclinic eddy-mean flow interactions



- Zonal mean and eddy components:

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Energy cycles

in the baroclinic eddy-mean flow interactions



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Energy cycles

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Energy cycles

in the baroclinic eddy-mean flow interactions



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Energy cycles

in the baroclinic eddy-mean flow interactions



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Energy cycles

in the baroclinic eddy-mean flow interactions



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Energy cycles

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Energy cycles

in the baroclinic eddy-mean flow interactions



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Energy cycles

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Energy cycles

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Energy cycles

in the baroclinic eddy-mean flow interactions



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Energy cycles

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Energy cycles

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Energy cycles

in the baroclinic eddy-mean flow interactions



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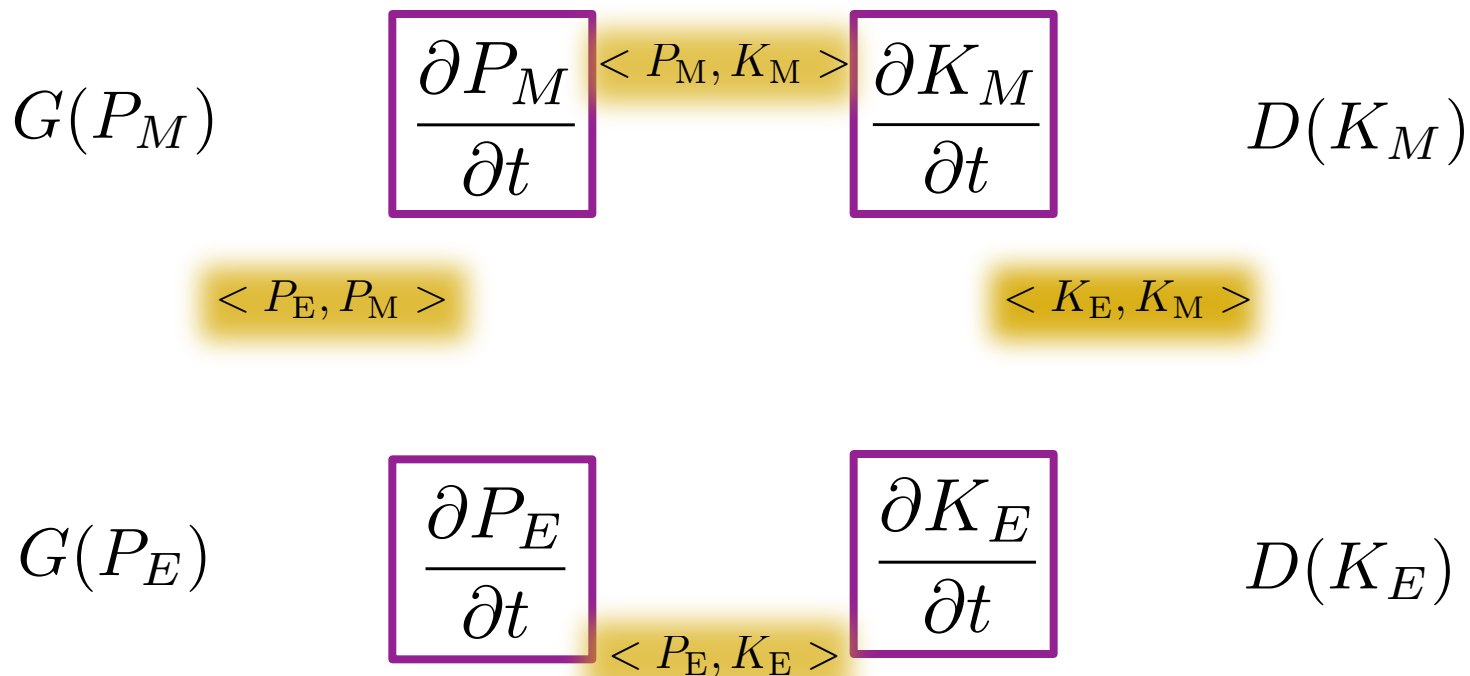
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Energy cycles

in the baroclinic eddy-mean flow interactions





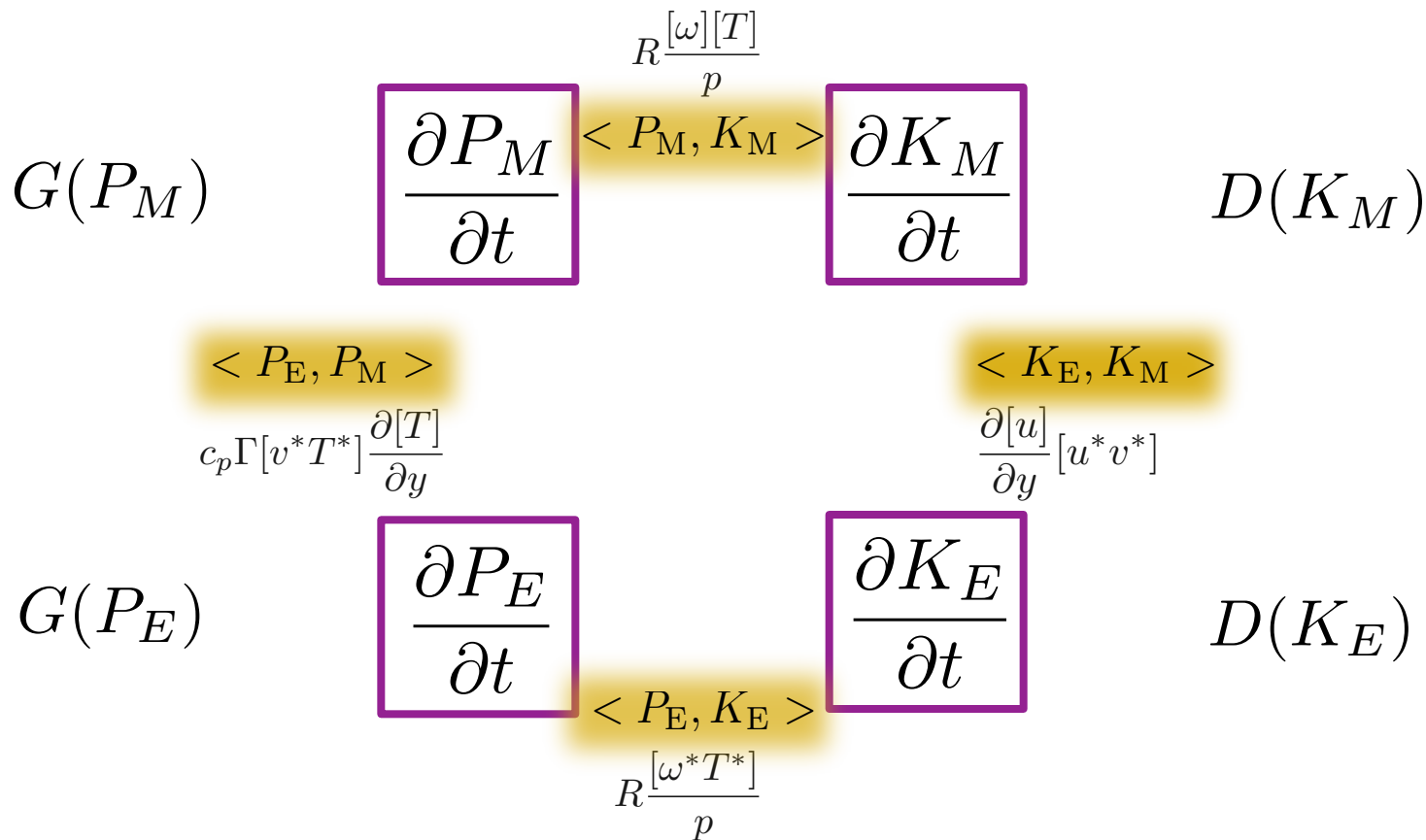
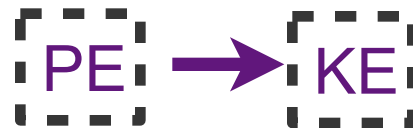
in the baroclinic eddy-mean flow interactions

授课教师：张洋 20



Energy cycles

in the baroclinic eddy-mean flow interactions





in the baroclinic eddy-mean flow interactions

-
- PE → KE

Diagram illustrating a thermodynamic cycle with four states: (P_M, K_M) , (P_E, K_M) , (P_E, K_E) , and (P_M, K_E) .

The cycle is represented by a square with purple boxes at the corners. The left side is labeled $G(P)$ and the right side $D(K)$.

The top horizontal process (from (P_M, K_M) to (P_E, K_M)) is labeled $< P_M, K_M >$ and $\frac{\partial P_M}{\partial t}$.

The bottom horizontal process (from (P_E, K_E) to (P_M, K_E)) is labeled $< P_E, K_E >$ and $\frac{\partial P_E}{\partial t}$.

The left vertical process (from (P_M, K_M) to (P_E, K_M)) is labeled $< P_E, P_M >$ and $\frac{\partial [T]}{\partial y}$.

The right vertical process (from (P_E, K_E) to (P_M, K_E)) is labeled $< K_E, K_M >$ and $\frac{\partial [u]}{\partial y}$.

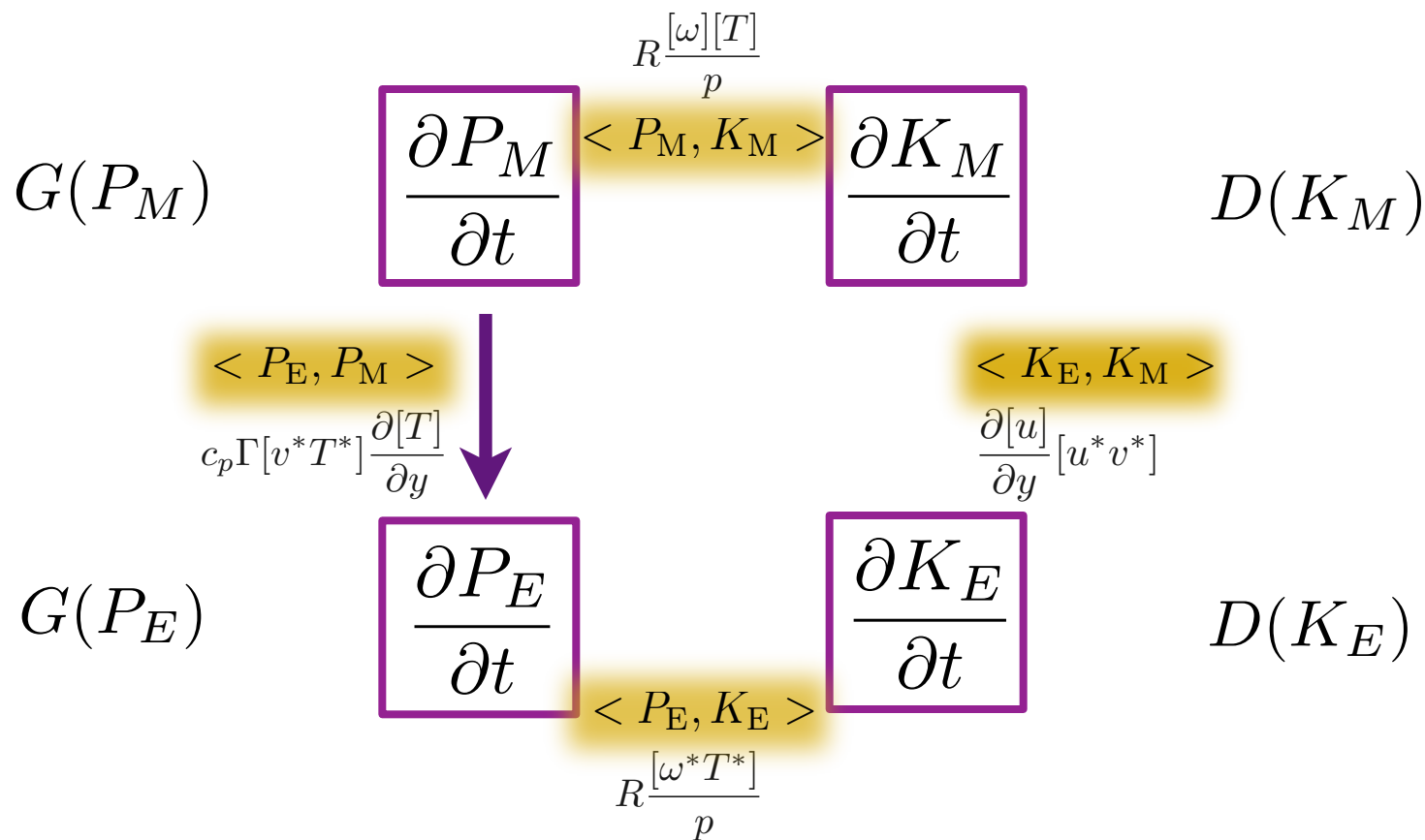


Energy cycles

in the baroclinic eddy-mean flow interactions



- Energy cycles in eddy life cycle: $\boxed{\text{PE}} \rightarrow \boxed{\text{KE}}$



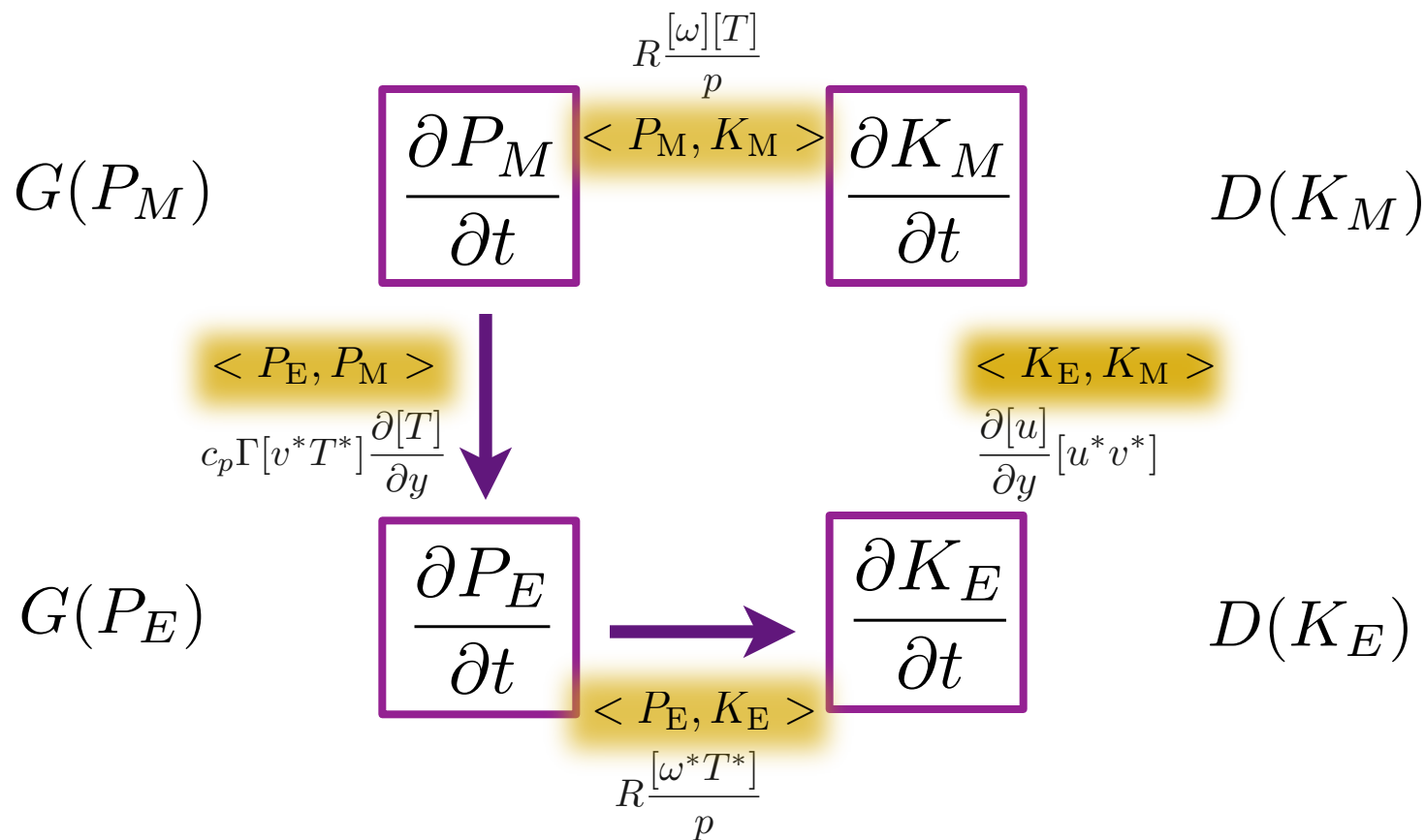


Energy cycles

in the baroclinic eddy-mean flow interactions



- Energy cycles in eddy life cycle: $\boxed{\text{PE}} \rightarrow \boxed{\text{KE}}$



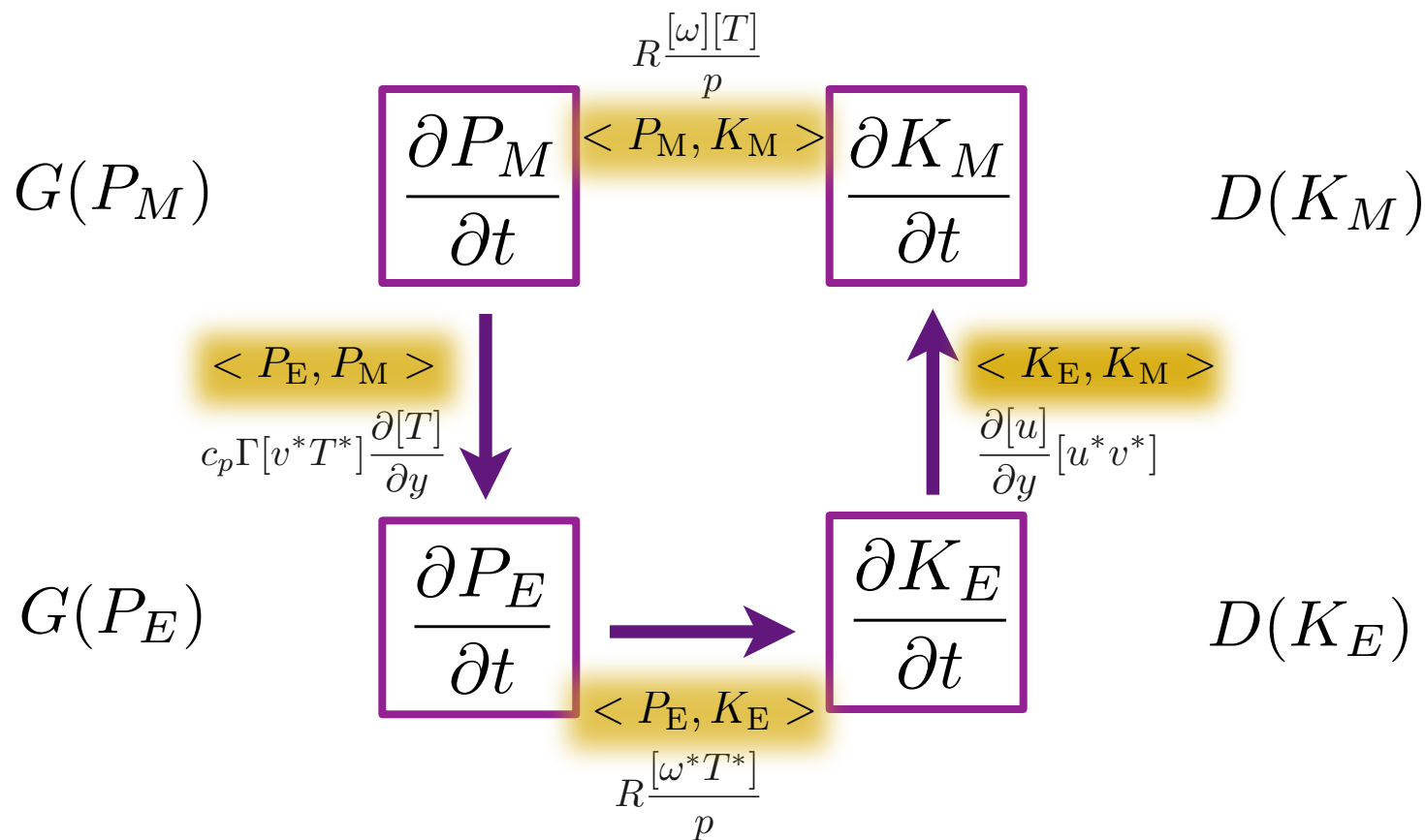


Energy cycles

in the baroclinic eddy-mean flow interactions



- Energy cycles in eddy life cycle: $\boxed{\text{PE}} \rightarrow \boxed{\text{KE}}$



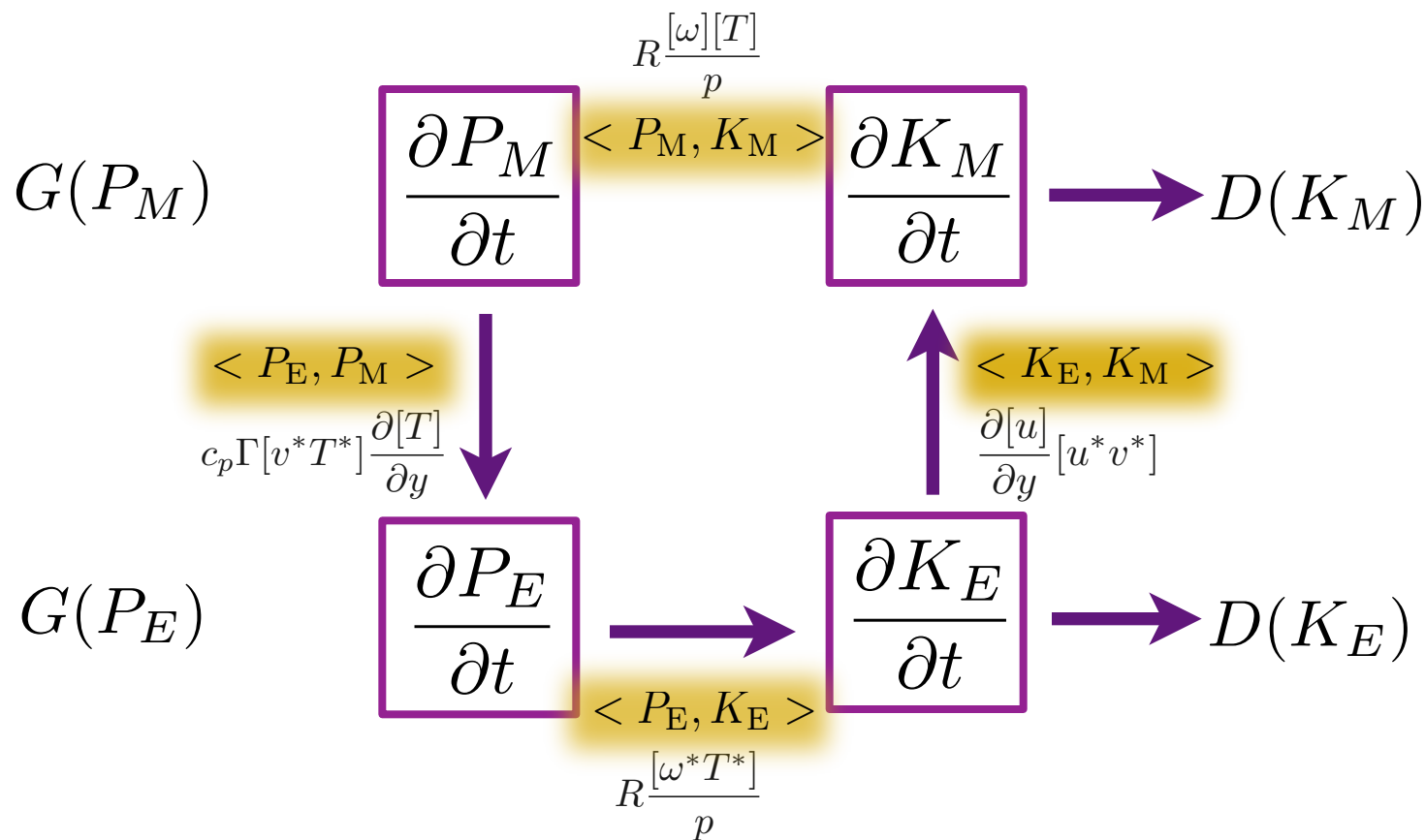


Energy cycles

in the baroclinic eddy-mean flow interactions



- Energy cycles in eddy life cycle: $\boxed{\text{PE}} \rightarrow \boxed{\text{KE}}$





Baroclinic eddy life cycle

- An E-P flux view

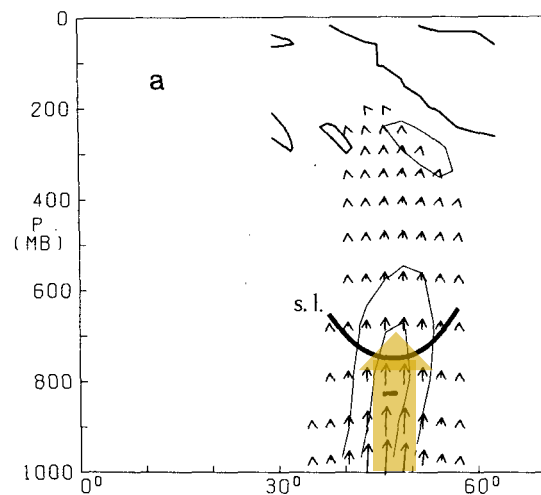


$$\mathcal{F} \equiv -[u^*v^*] \mathbf{j} + f \frac{[v^*\theta^*]}{\partial\theta_s/\partial p} \mathbf{k}$$

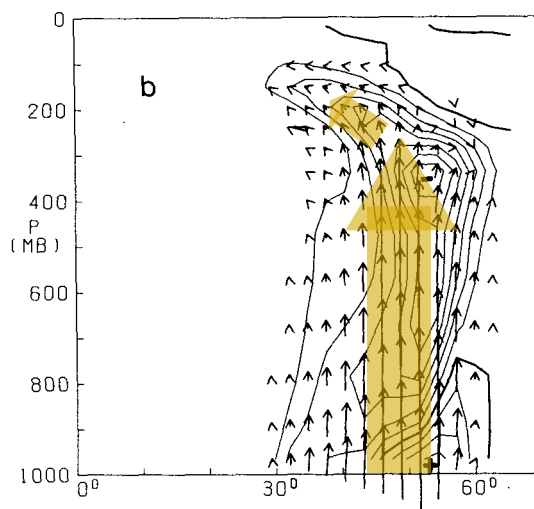
$$\vec{\mathcal{F}} = \mathbf{c}_g \mathcal{A}$$

Eddies: generate at lower level,
propagate **upwards** and **away** from the
eddy source region

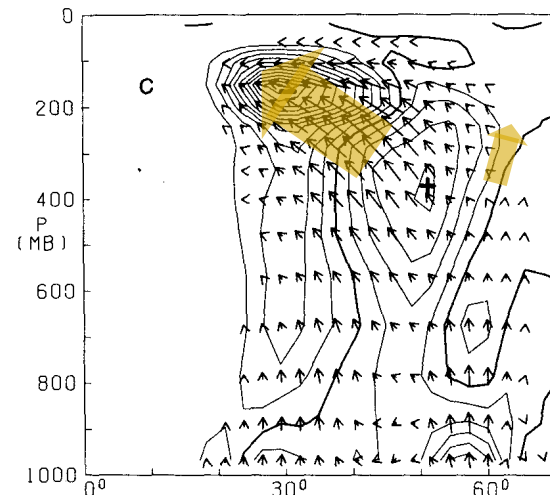
Numerical results from
Simmons and Hoskins,
1978, JAS



TOTAL E-P FLUX DIVERGENCE
DAY .00



TOTAL E-P FLUX DIVERGENCE
DAY 5.00



TOTAL E-P FLUX DIVERGENCE
DAY 8.00

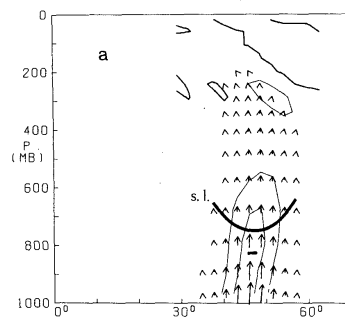
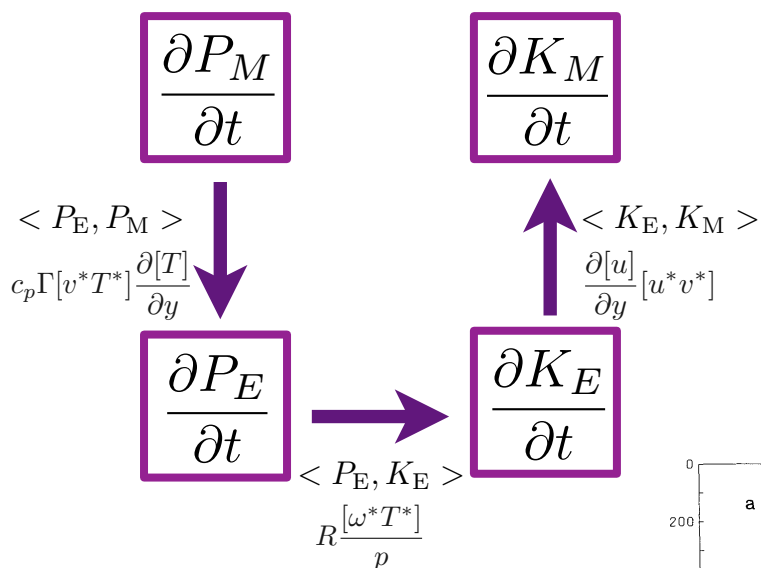


Baroclinic eddies

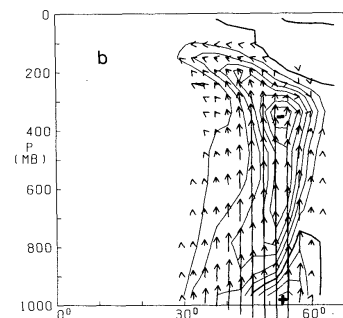
- baroclinic eddy life cycle



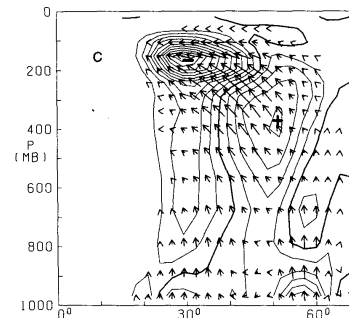
- Westerly jet and energy cycle:



TOTAL E-P FLUX DIVERGENCE
DAY 0.00



TOTAL E-P FLUX DIVERGENCE
DAY 5.00



TOTAL E-P FLUX DIVERGENCE
DAY 8.00

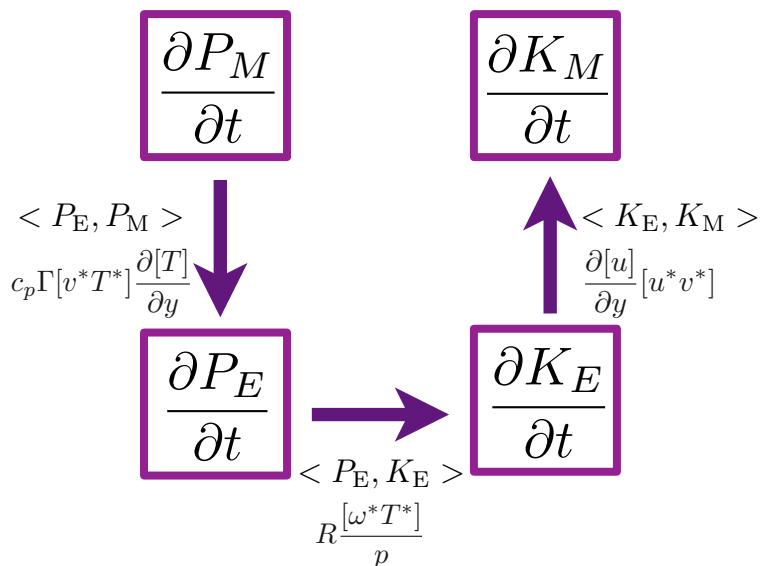


Baroclinic eddies

- baroclinic eddy life cycle



- Westerly jet and energy cycle:



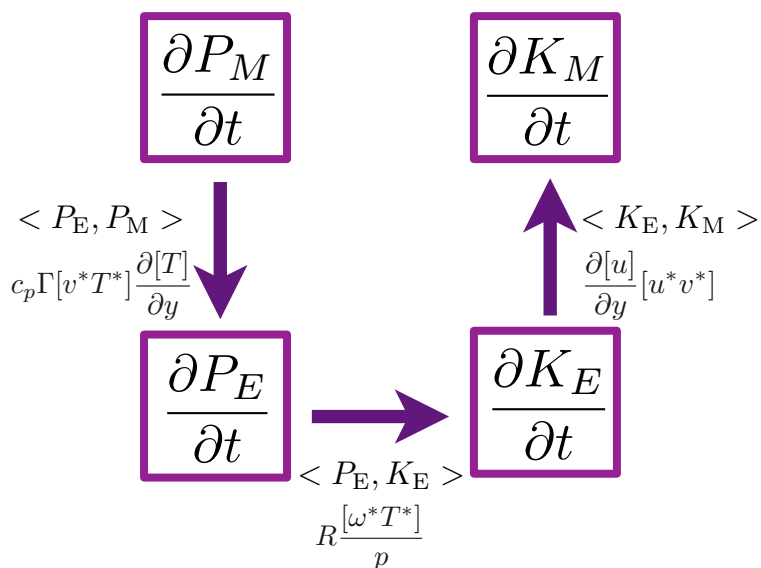


Baroclinic eddies

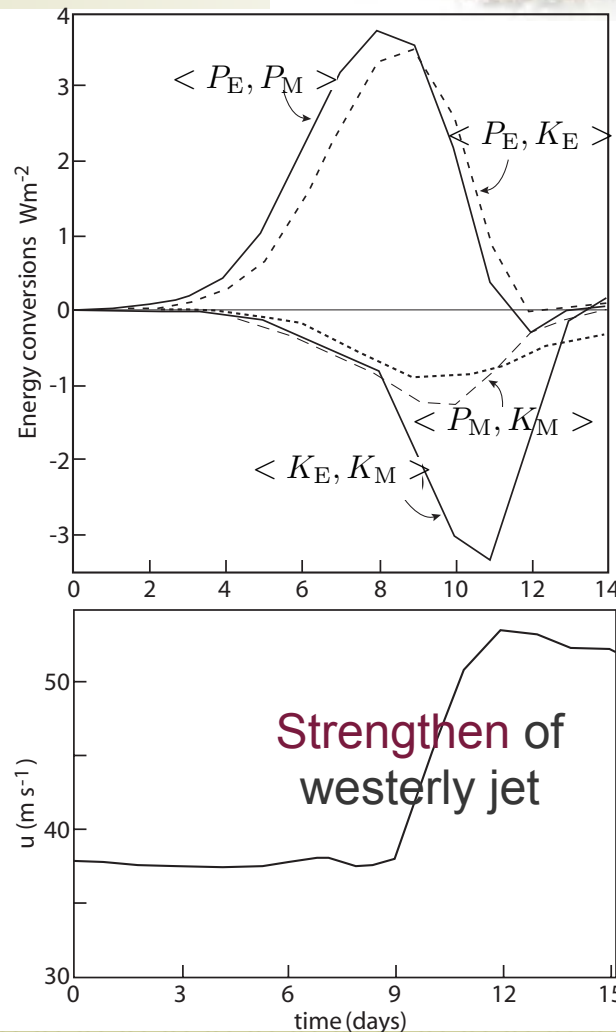
- baroclinic eddy life cycle



- Westerly jet and energy cycle:



Numerical results from
Simmons and Hoskins,
1978, JAS



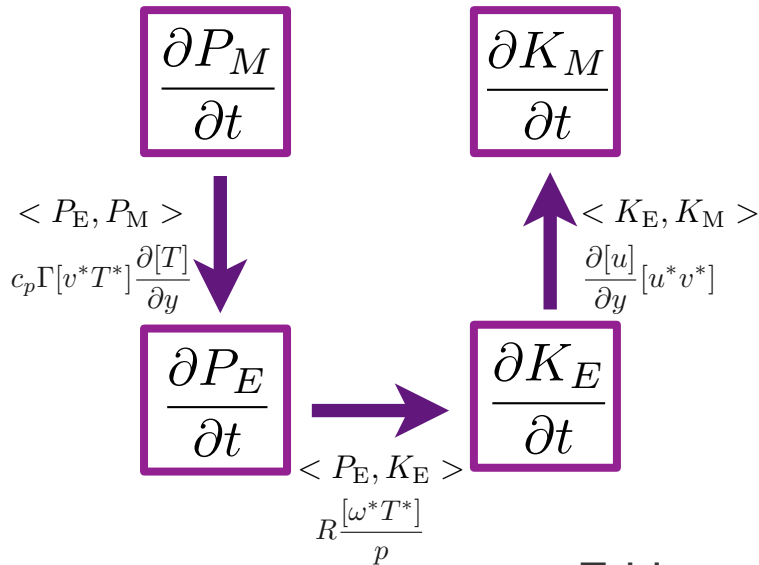


Baroclinic eddies

- baroclinic eddy life cycle

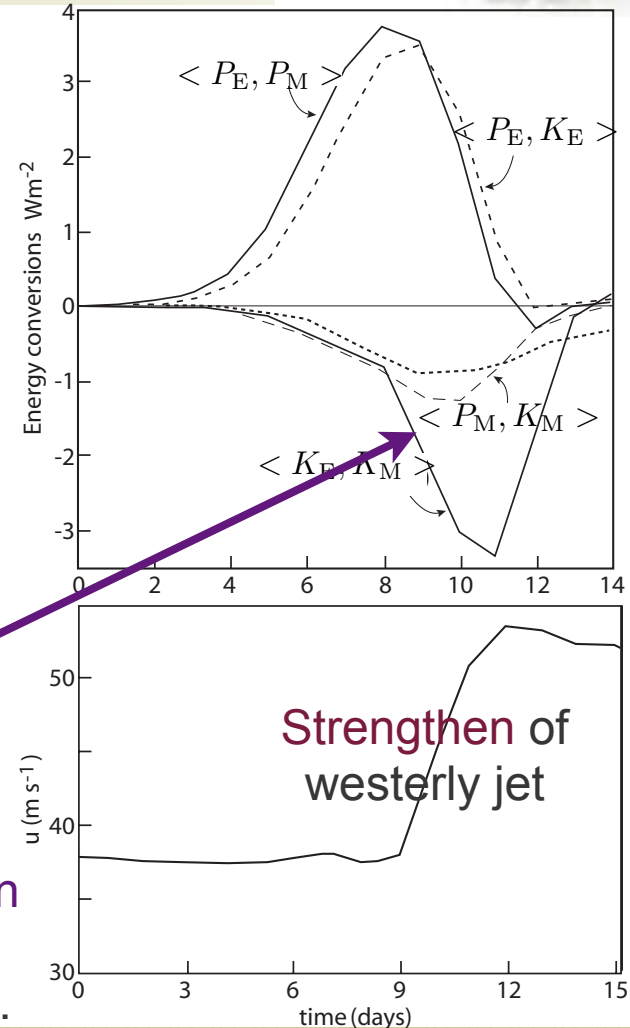


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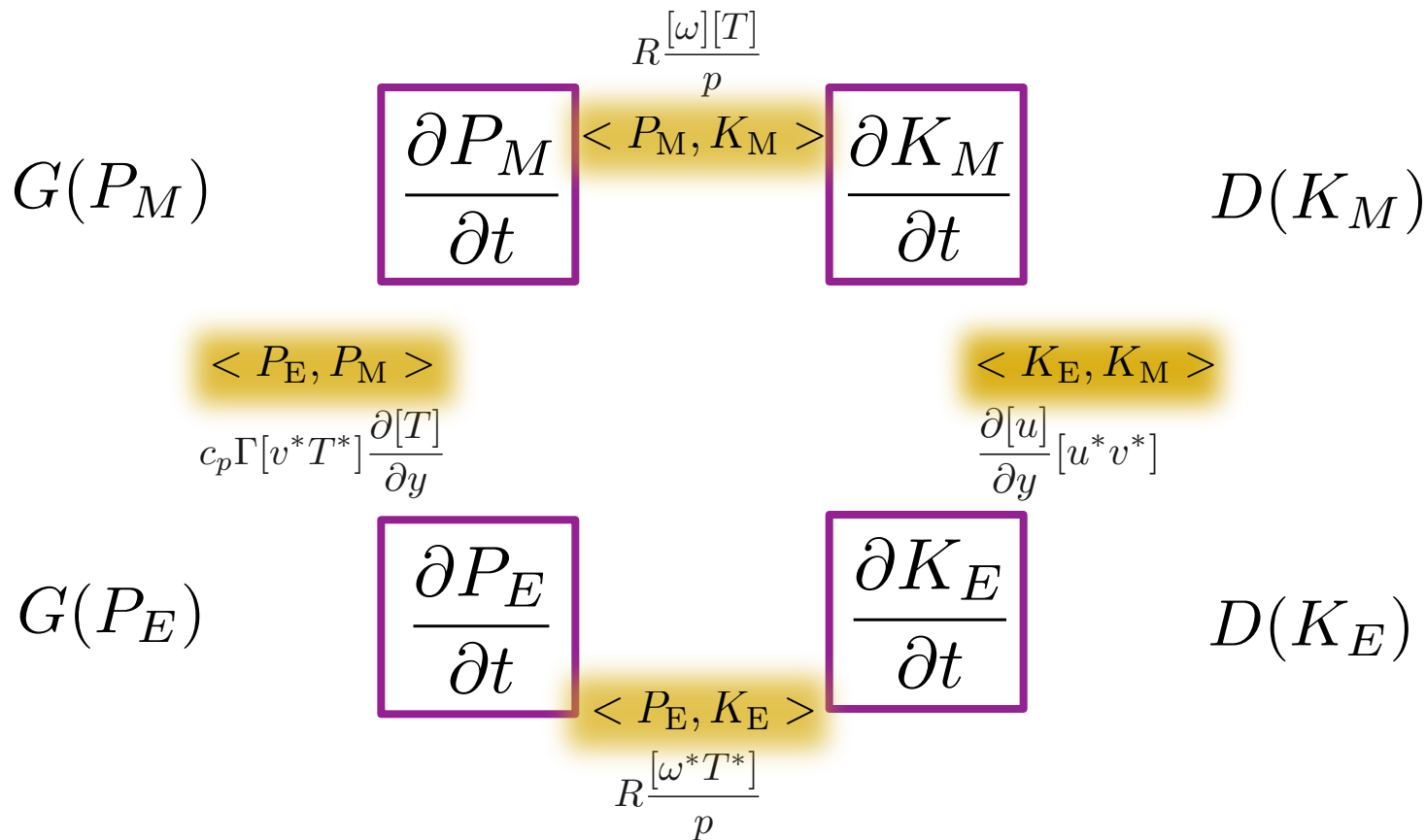
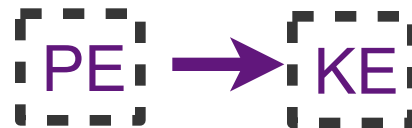
Eddy momentum flux
grows, which extracts
kinetic energy from the
eddies to the zonal mean
flow, then the growth of
the eddy energy ceases.





Energy cycles

in the baroclinic eddy-mean flow interactions



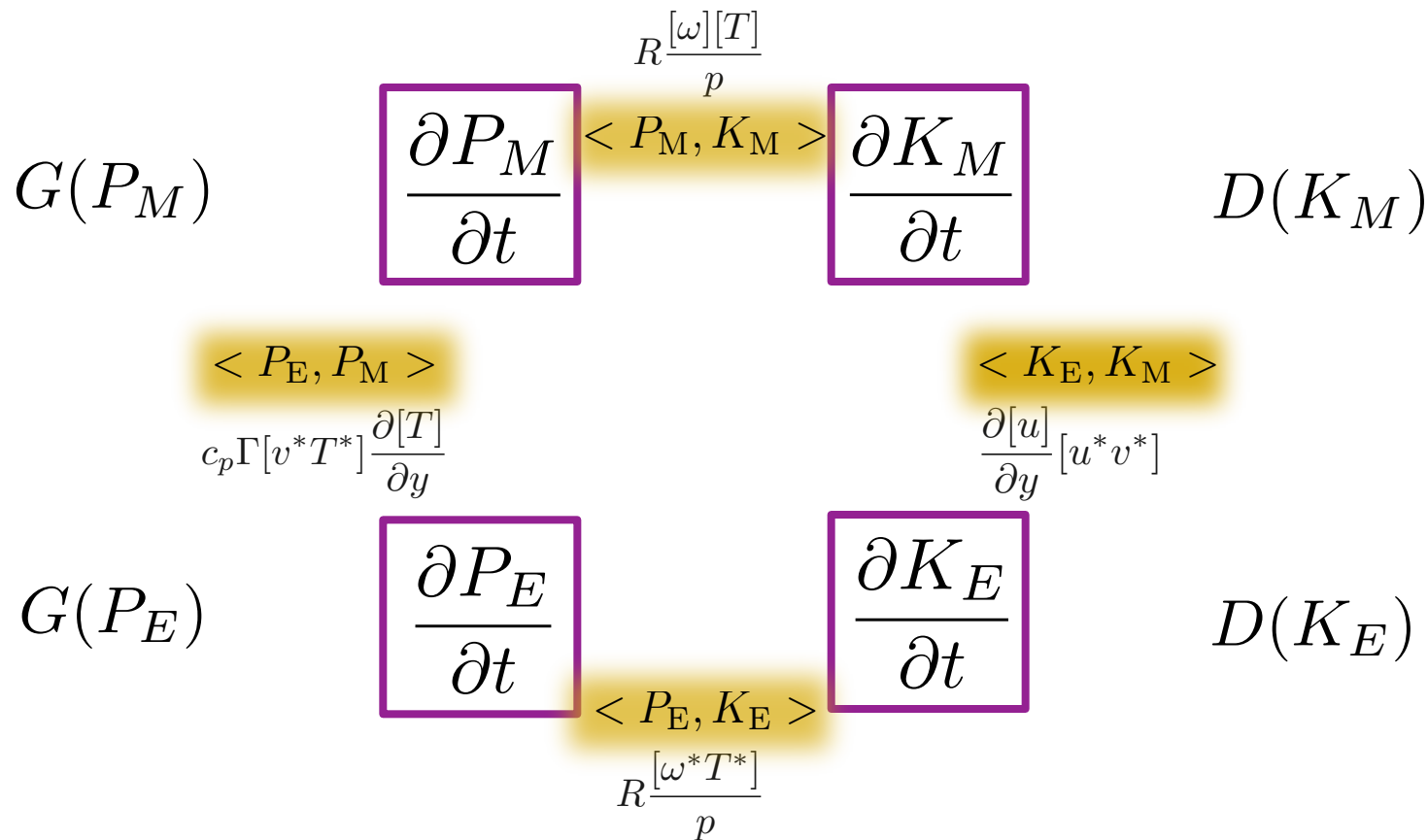
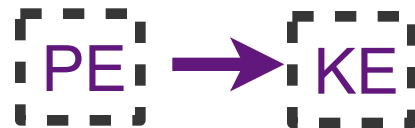


Energy cycles

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- Energy cycles in equilibrium:



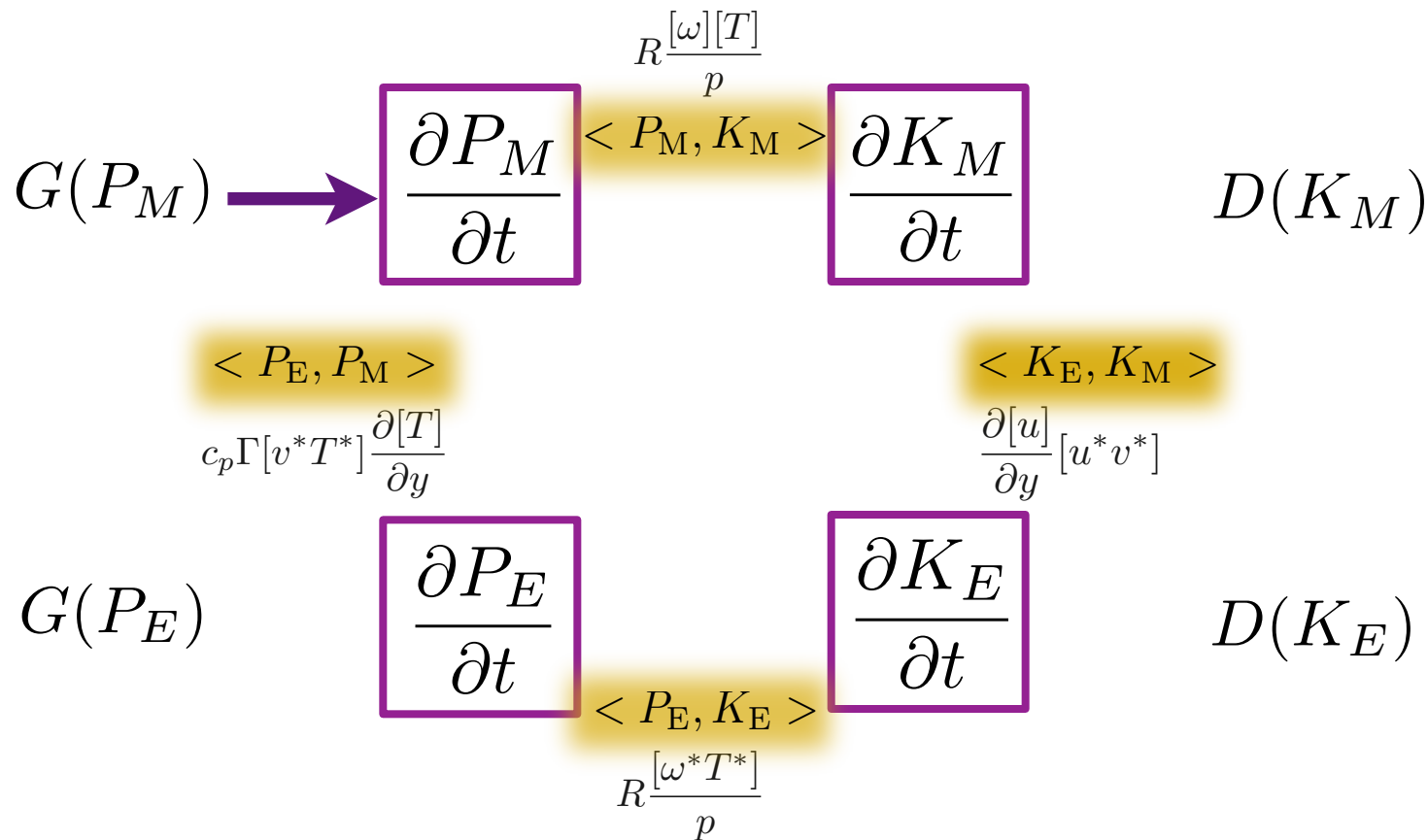
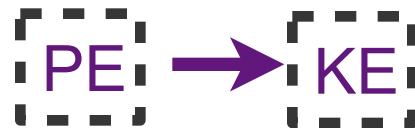


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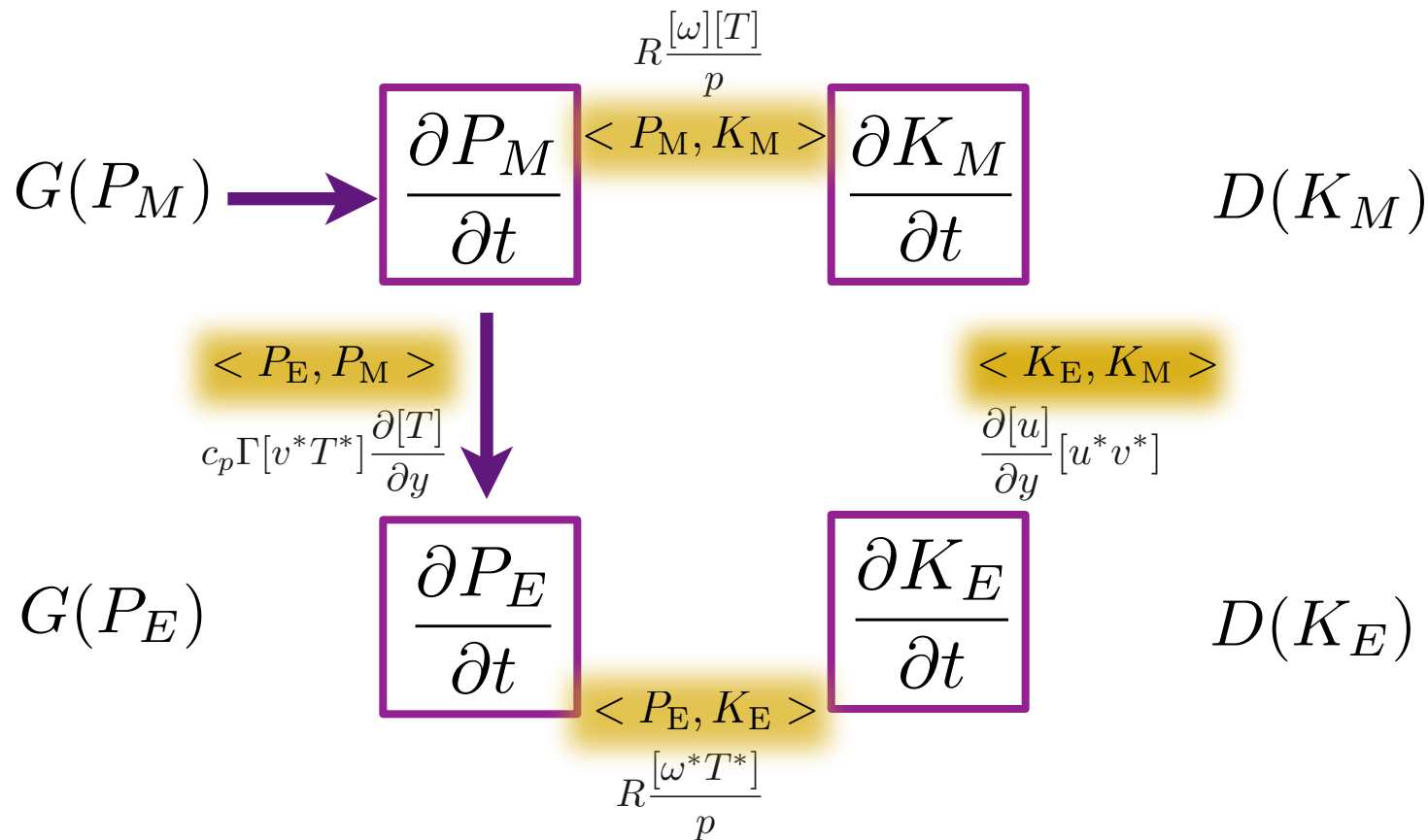
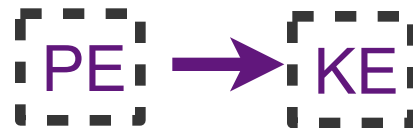


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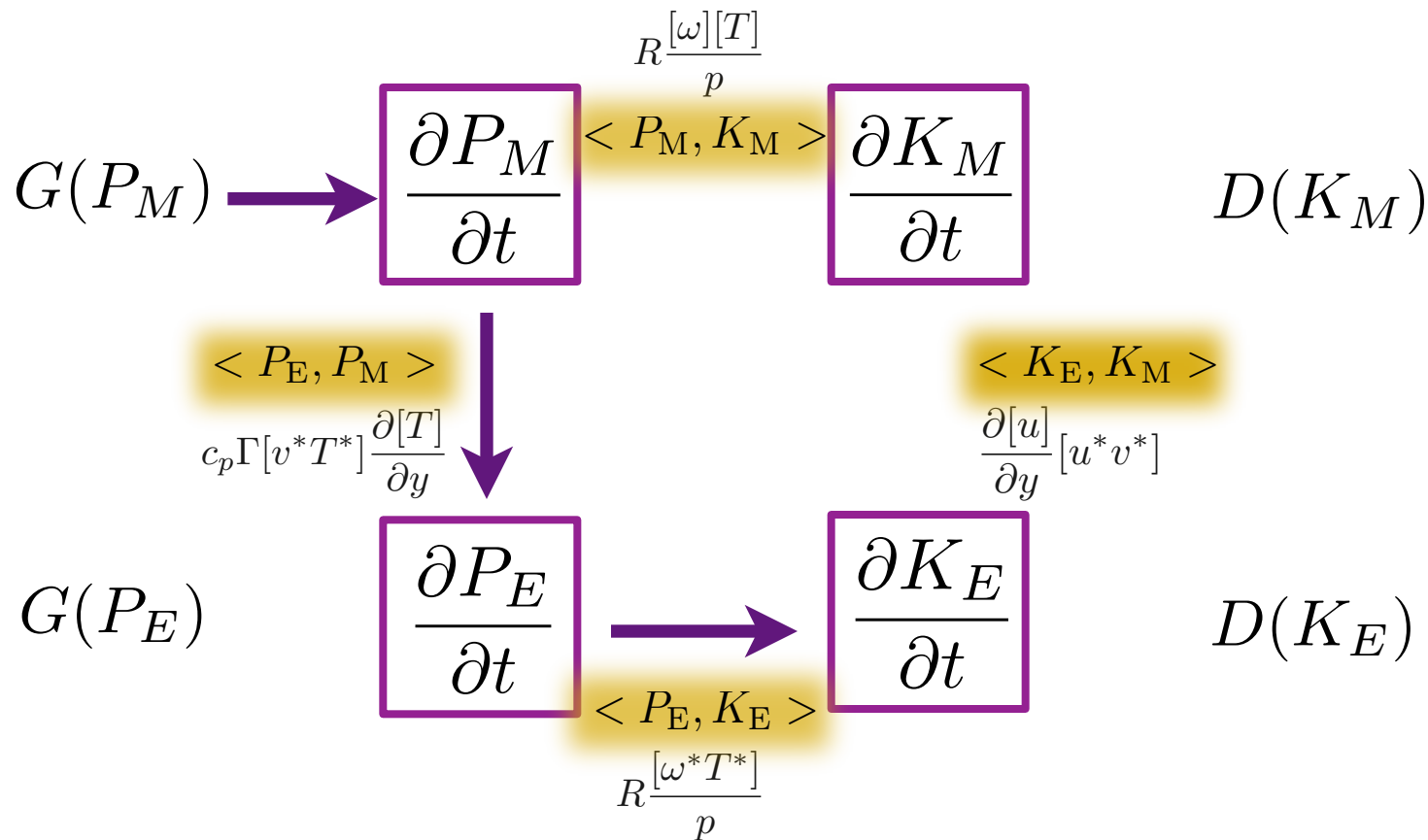
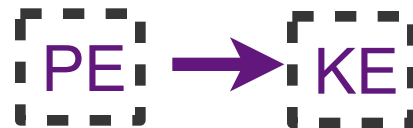


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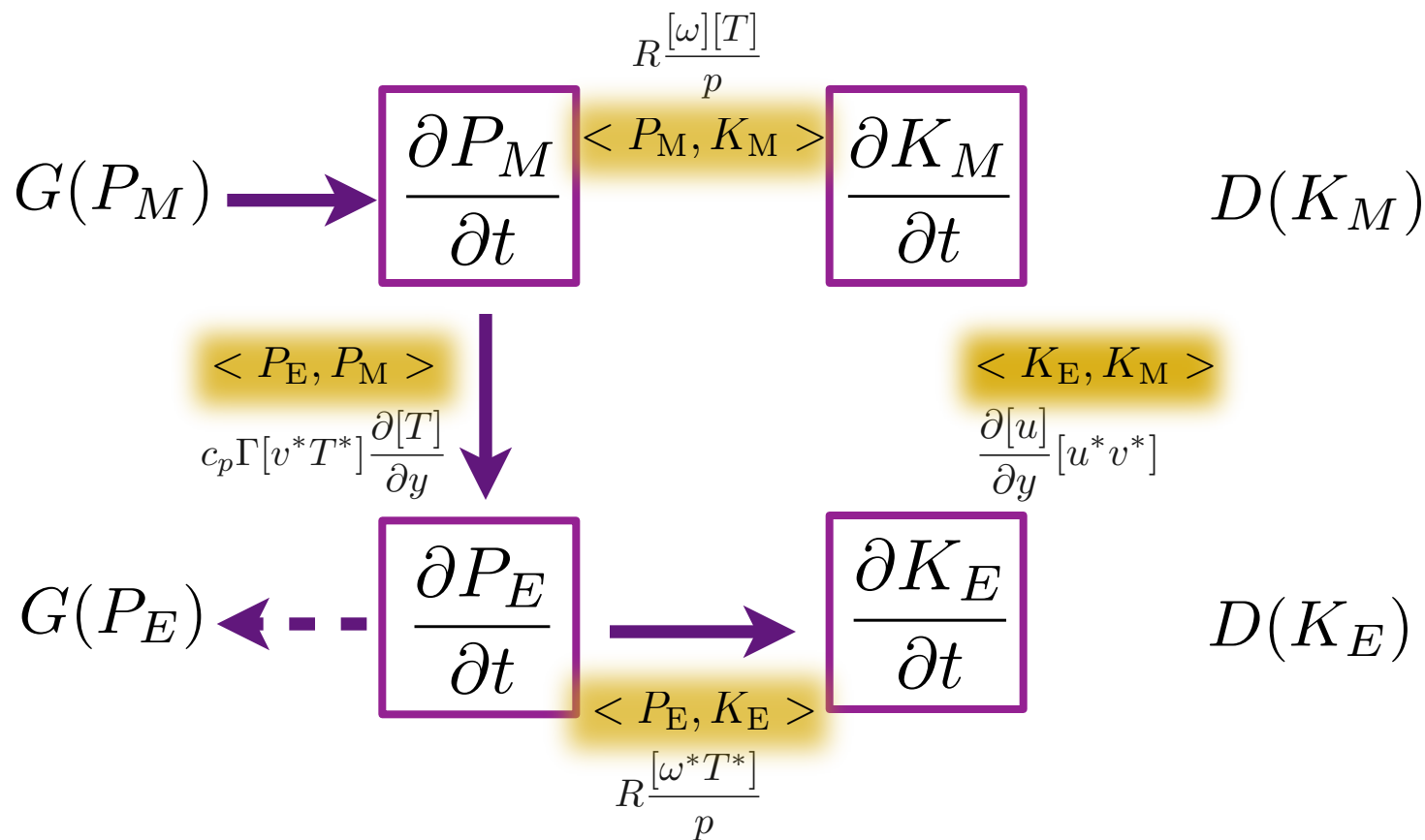
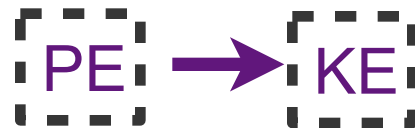


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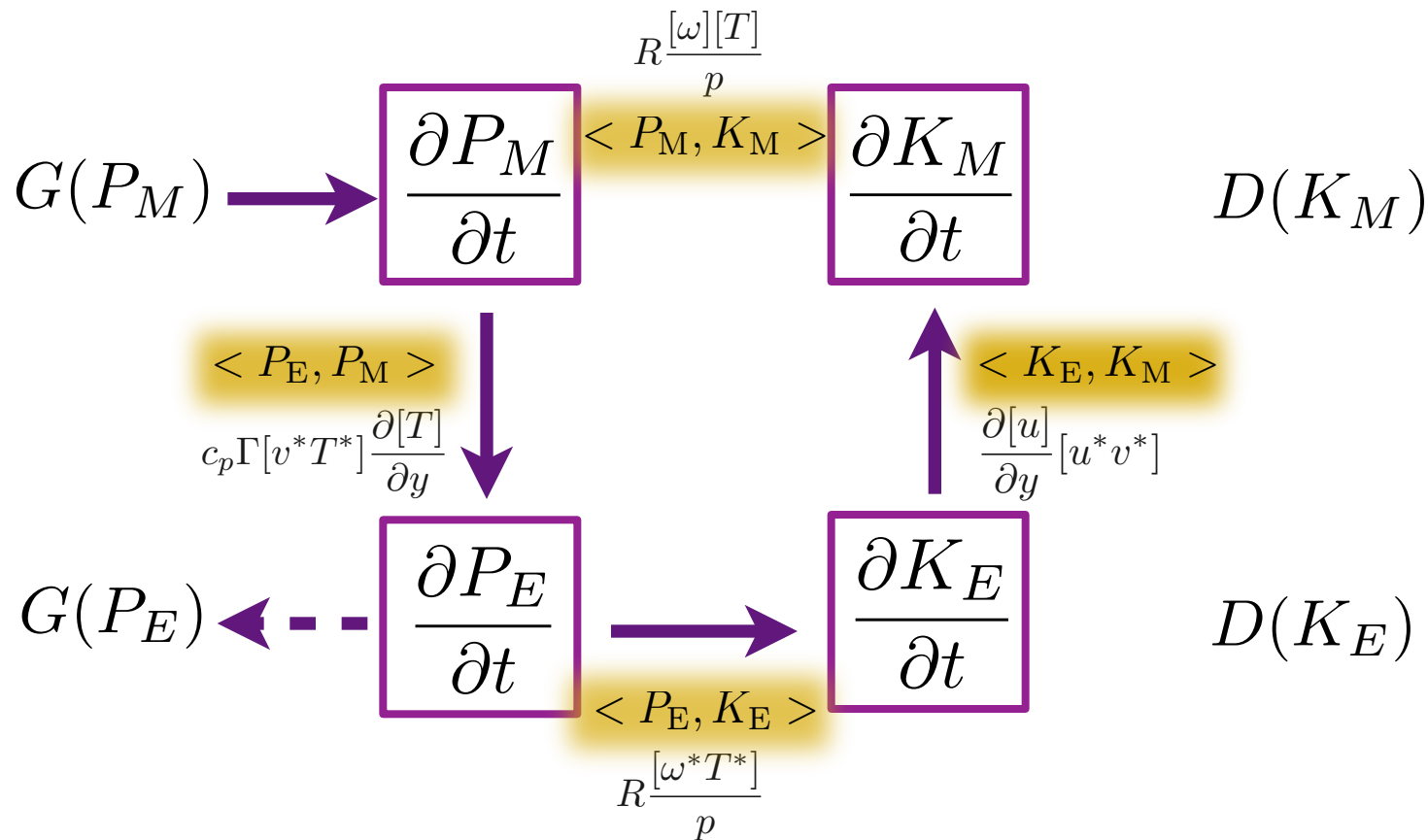
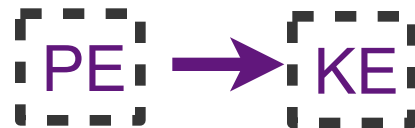


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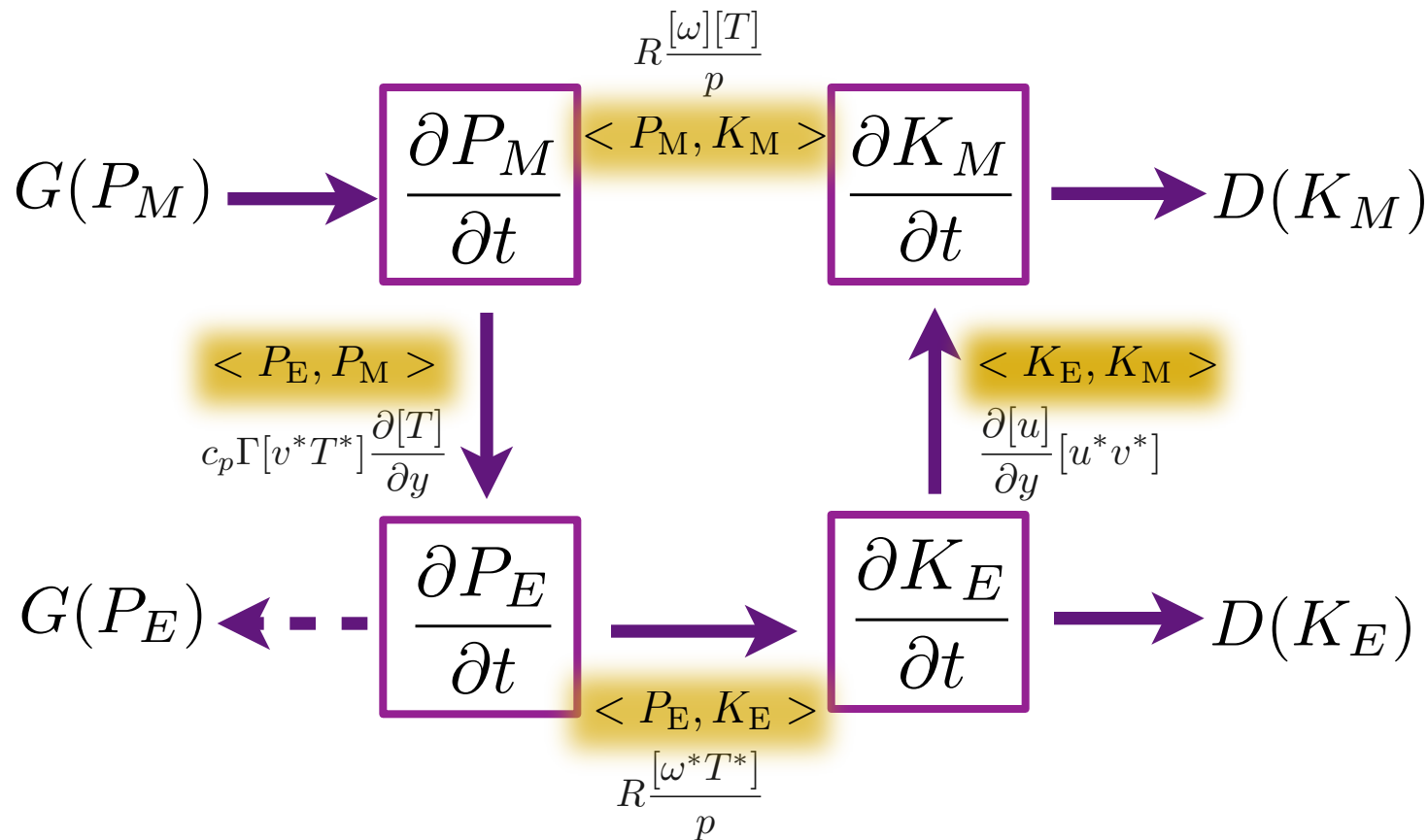
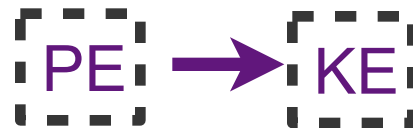


Energy cycles

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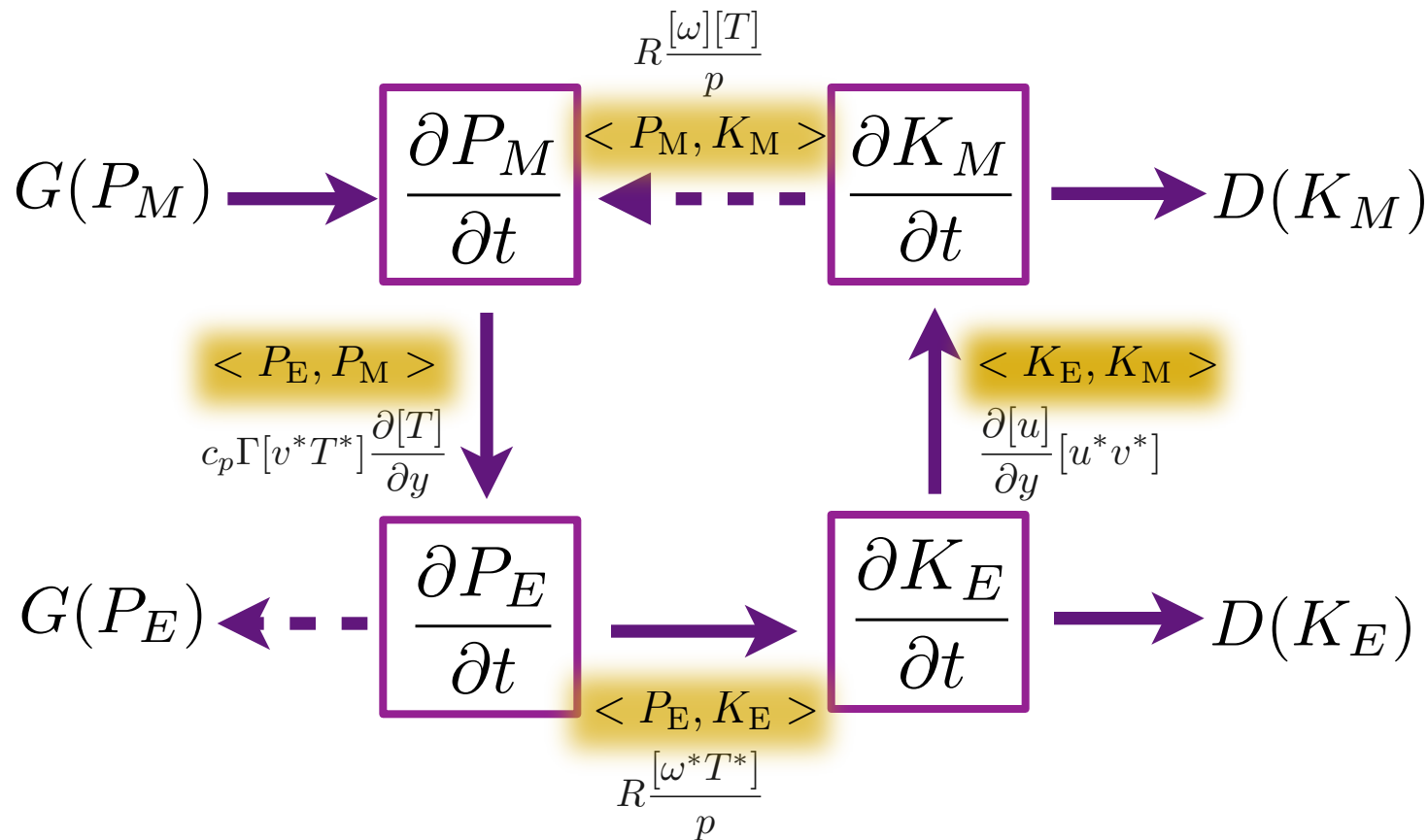
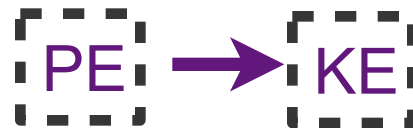


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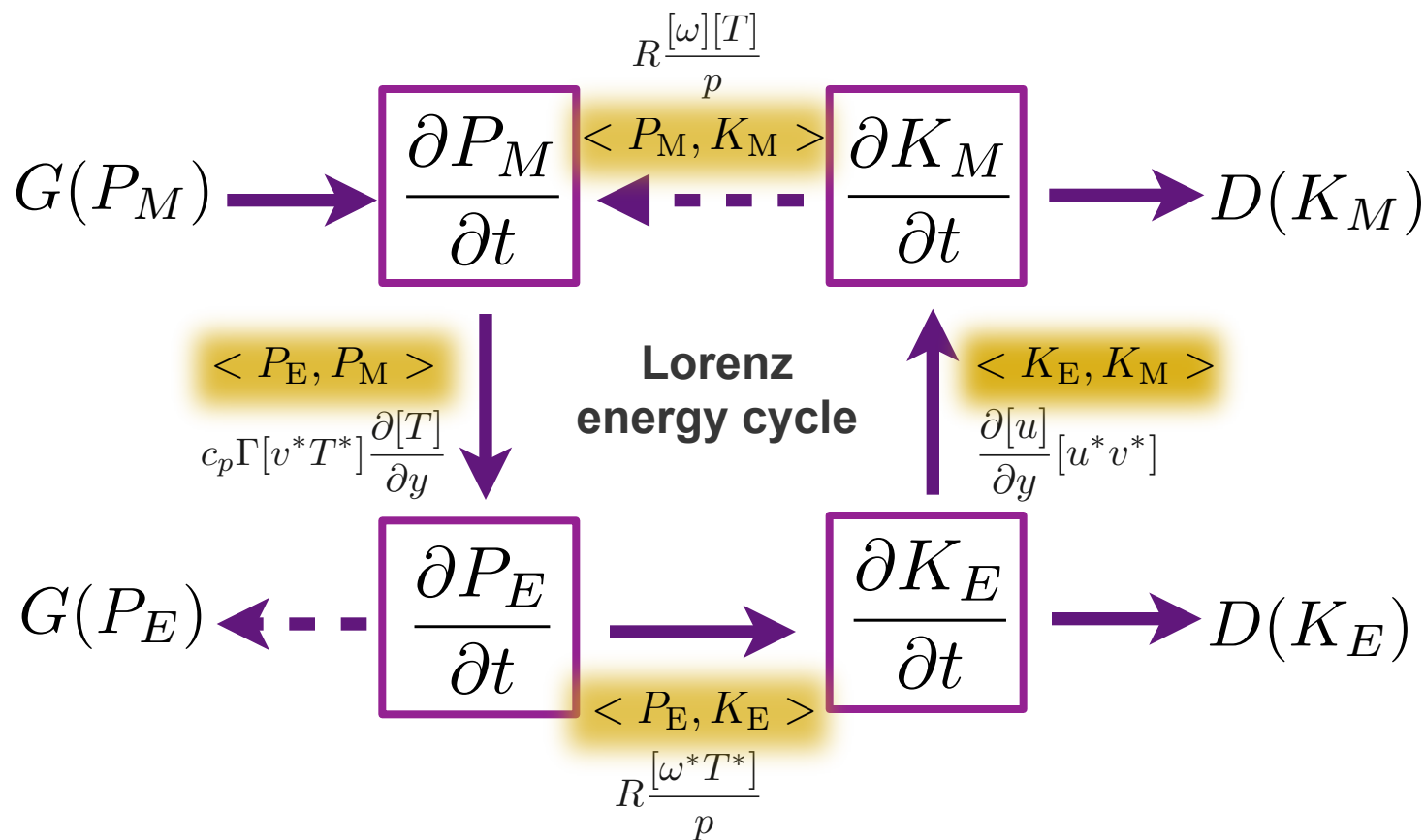
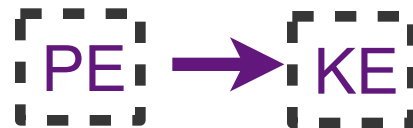


Energy cycles

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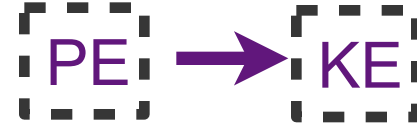


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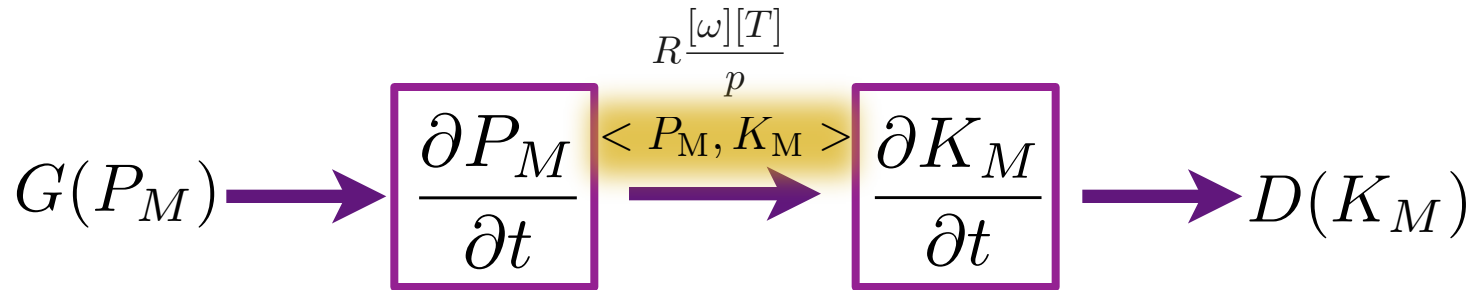
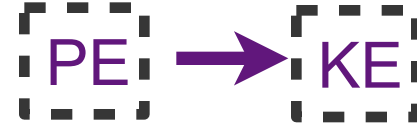


Energy cycles in Hadley Cell





Energy cycles in Hadley Cell

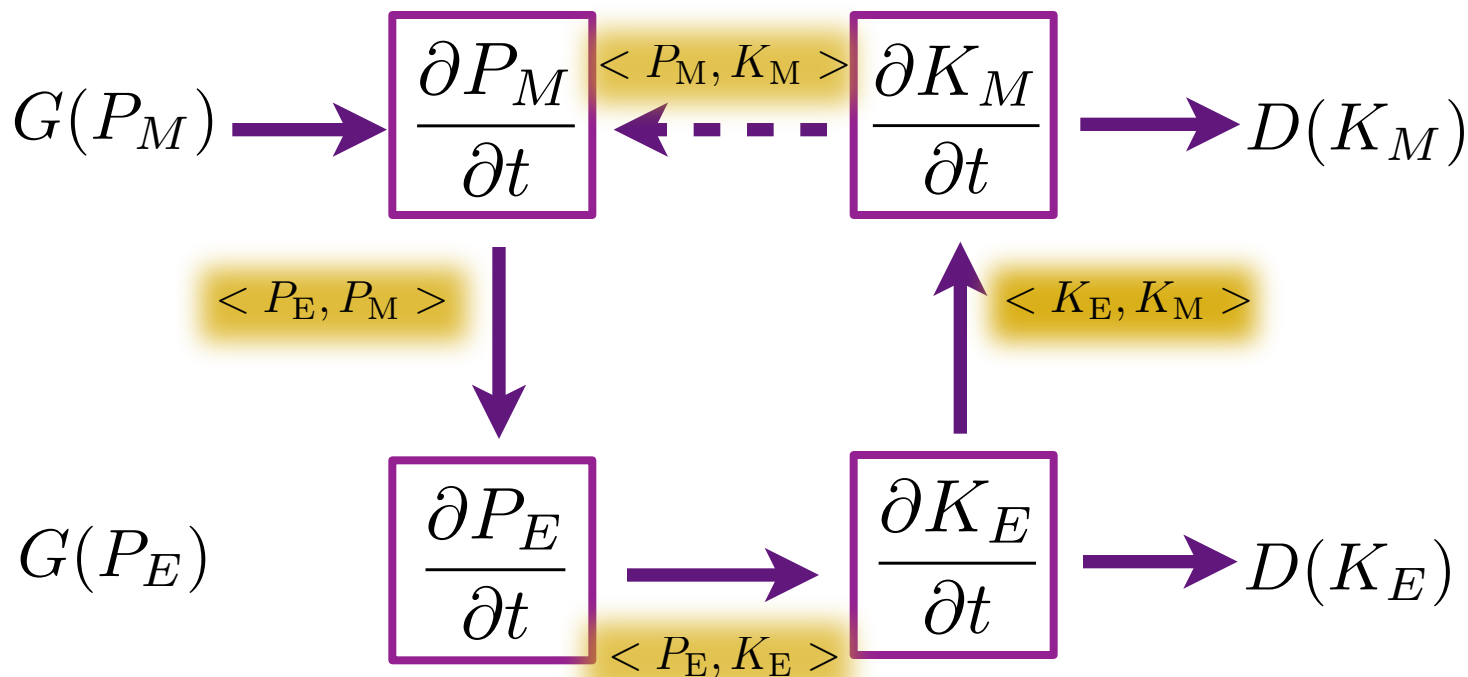


If assume no eddies.



Energy cycles

in the baroclinic eddy-mean flow interactions



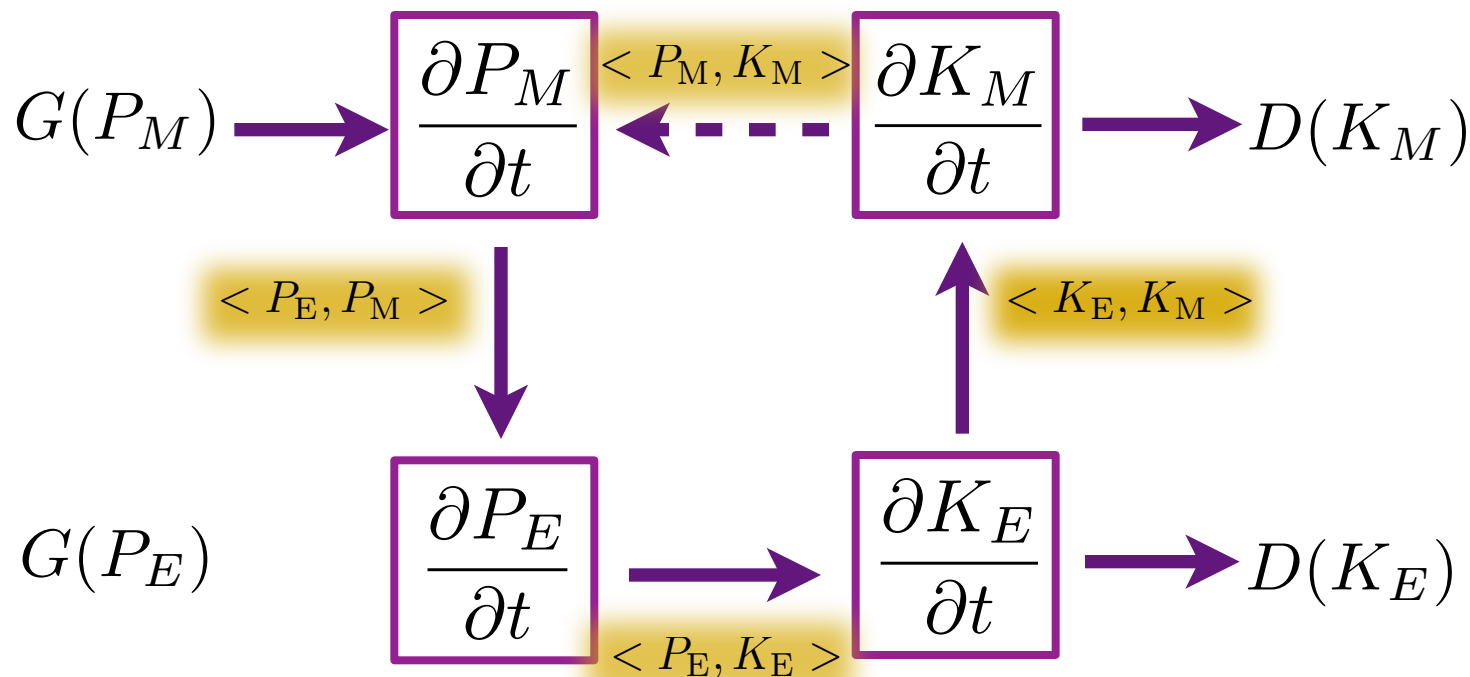


Energy cycles

in the baroclinic eddy-mean flow interactions



- Energy cycles in real atmosphere:



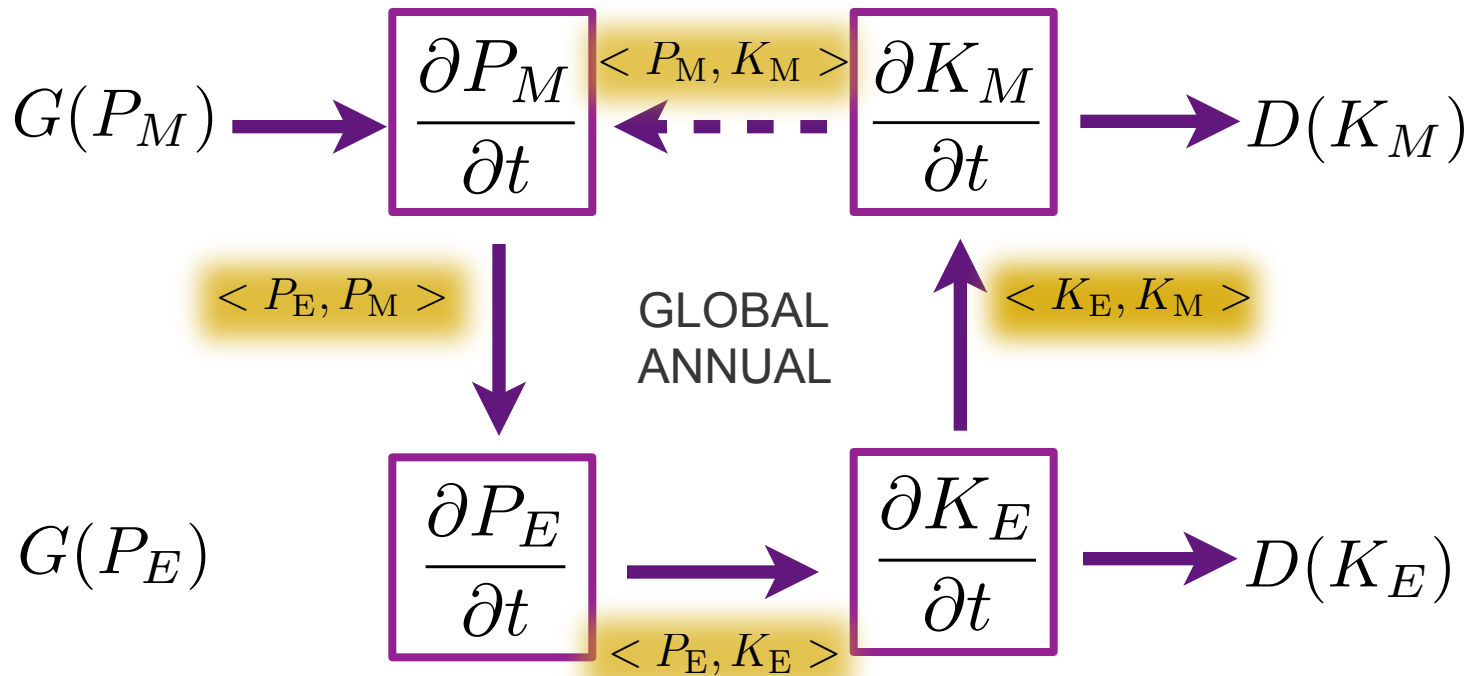


Energy cycles

in the baroclinic eddy-mean flow interactions



- Energy cycles in real atmosphere:



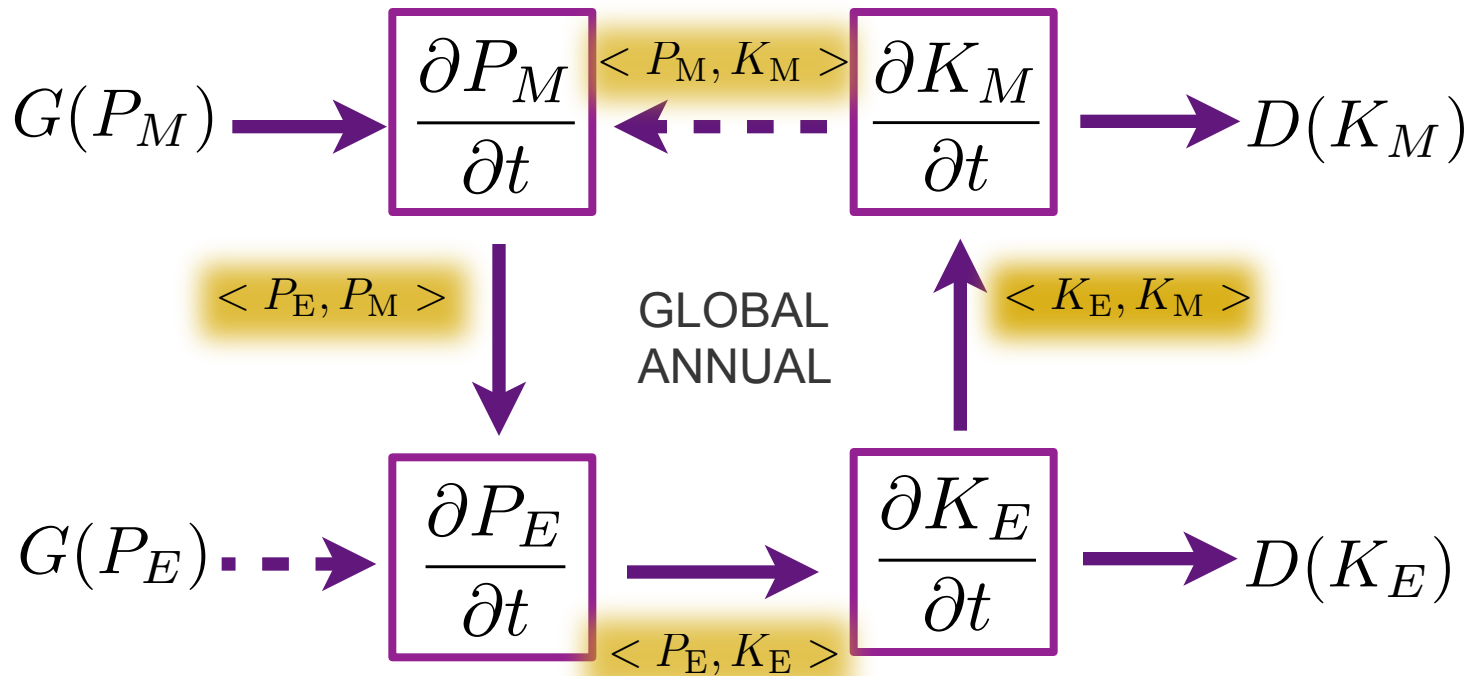


Energy cycles

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- Energy cycles in real atmosphere:





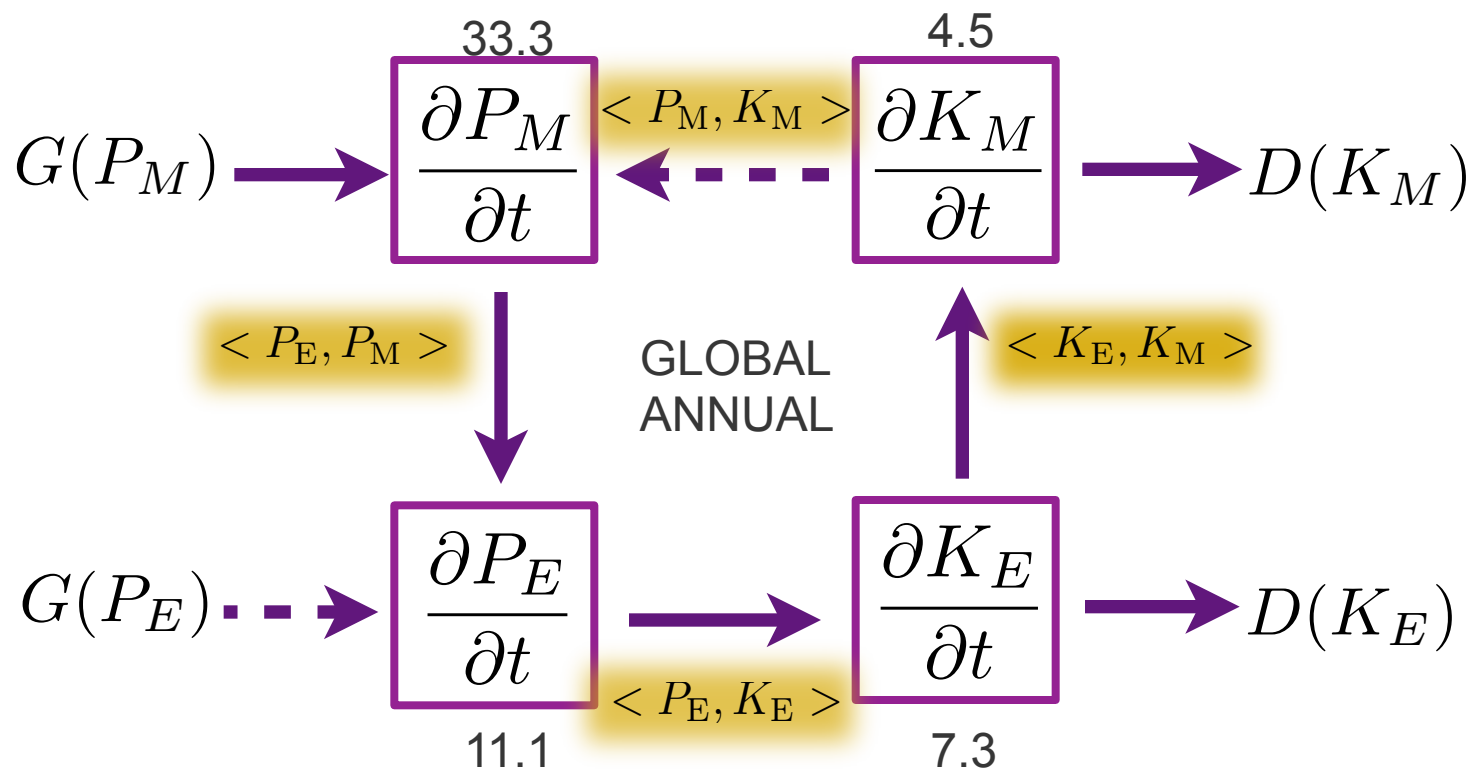
Energy cycles

in the baroclinic eddy-mean flow interactions



- Energy cycles in real atmosphere:

energy: $10^5 Jm^{-2}$





Energy cycles

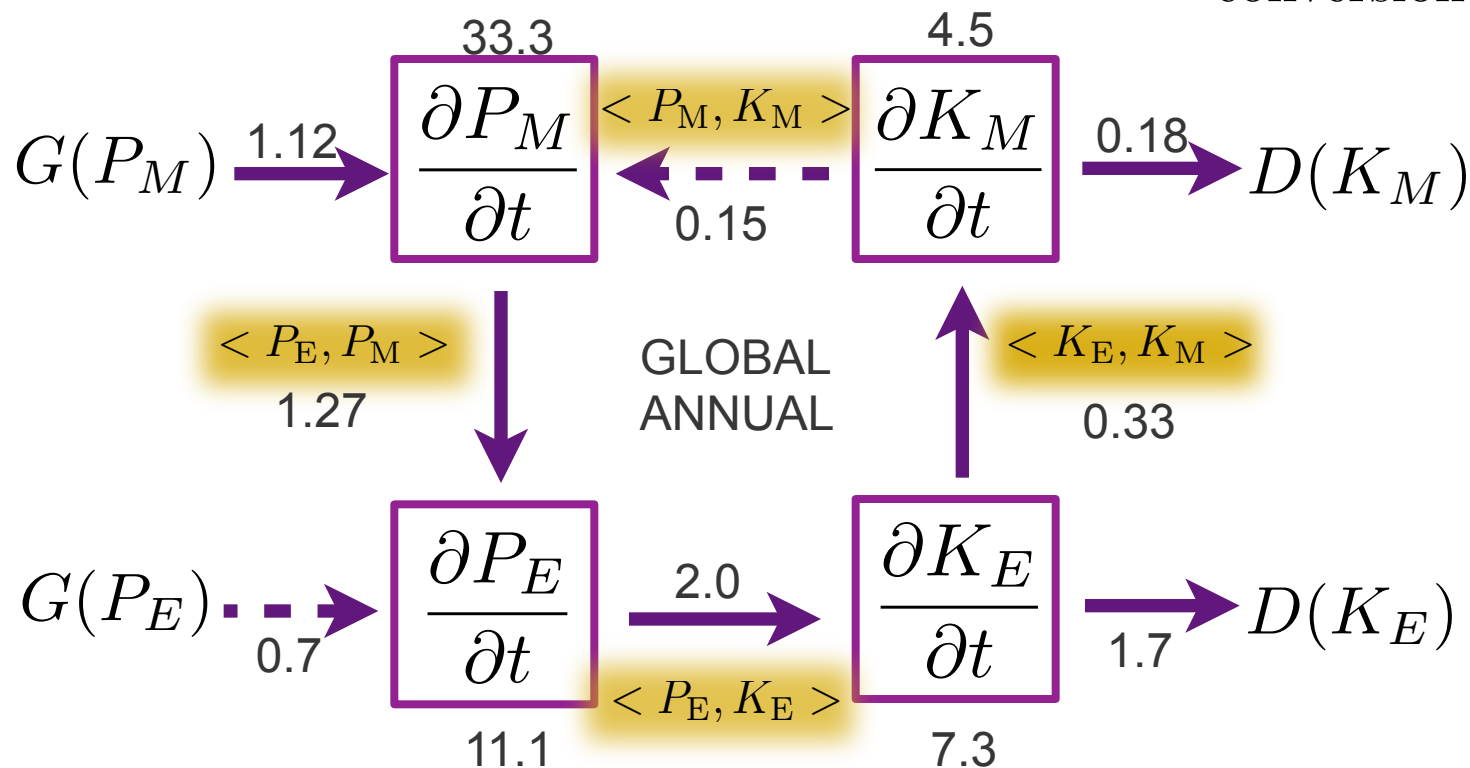
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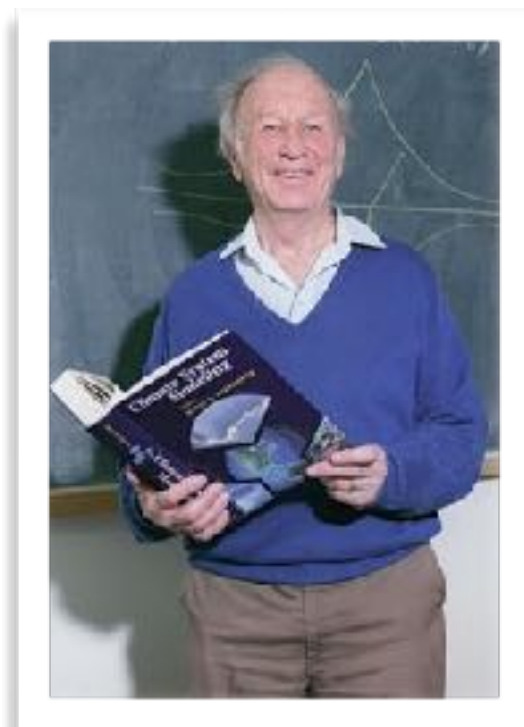


■ Energy cycles in real atmosphere:

energy: $10^5 J m^{-2}$

conversion: $W m^{-2}$





Edward N. Lorenz, a Meteorologist and a Father of Chaos Theory, Dies at 90

By [KENNETH CHANG](#)
Published: April 17, 2008

Edward N. Lorenz, a meteorologist who tried to predict the weather with computers but instead gave rise to the modern field of chaos theory, died Wednesday at his home in Cambridge, Mass. He was 90.



M.I.T. News Office

Edward N. Lorenz

The cause was cancer, said his daughter Cheryl Lorenz.

In discovering “deterministic chaos,” Dr. Lorenz established a principle that “profoundly influenced a wide range of basic sciences and brought about one of the most dramatic changes in mankind’s view of nature since Sir [Isaac Newton](#),” said a committee that awarded him the 1991 Kyoto Prize for basic sciences.

Dr. Lorenz is best known for the notion of the “butterfly effect,” the idea that a small disturbance like the flapping of a butterfly’s wings can induce enormous consequences.

As recounted in the book “Chaos” by James Gleick, Dr. Lorenz’s accidental discovery of chaos came in the winter of 1961. Dr. Lorenz was running simulations of weather using a simple computer model. One day, he wanted to repeat one of the simulations for a longer time, but instead of repeating the whole simulation, he started the second run in the middle, typing in numbers from the first run for the initial conditions.

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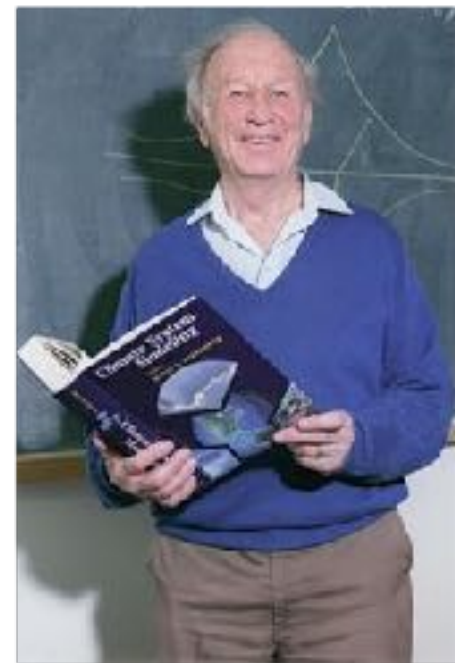
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Summary & Discussion



- Observations
- The Ferrel Cell
- Baroclinic eddies
 - Review: baroclinic instability and baroclinic eddy life cycle
 - Eddy-mean flow interaction, E-P flux
 - Transformed Eulerian Mean equations
- Eddy-driven jet
- The energy cycle



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Summary & Discussion



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1. The role of moisture;

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Summary & Discussion



- Observations
- The Ferrel Cell
- Baroclinic eddies

1. The role of moisture;
2. Quantify (parameterize) the relation between eddies and mean flow;
3. Zonal variations.

- Review: baroclinic instability and baroclinic eddy life cycle
- Eddy-mean flow interaction, E-P flux
- Transformed Eulerian Mean equations
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Due: 2022.12.12 (周一)



Assignment 4, Fall 2022

在第四章中，我们从准地转近似下的纬向平均风场、温度场的趋势方程出发，定义了E-P通量。但是该定义下的E-P通量并没有考虑到大气湿过程的影响。如果从第三章介绍的水汽方程出发，我们可以按照以下步骤定义出一个包含大气大尺度运动中湿过程作用的广义的E-P通量。

1) 在准地转近似下，如果我们按照对热力学方程的简化方法，将比湿 (specific humidity) q ，分解成一个标准比湿 q_s (reference specific humidity) 和变化量 q' ，并且同样假设 $\partial q / \partial p$ 的水平变化很小，请证明在准地转近似下 p 坐标系下的纬向平均比湿 $[q]$ 的变化方程为：

$$\frac{\partial [q']}{\partial t} + \frac{\partial q_s}{\partial p} [\omega] = -[C - S] - \frac{\partial}{\partial y} [v' q'],$$

其中 $C - S$ 为水汽方程在准地转近似下的源汇项，表征由大尺度运动所带来的净凝结率。

2) 如果重新定义一个非绝热加热项 Q_m ，使得 $Q_m = Q - L[C - S](\frac{p}{p_0})^{R/c_p}$ ，请推导出一个关于 $[\theta + \frac{L}{c_p} q']$ 的变化方程。

3) 根据以上推导出的新方程和准地转近似下 $[v]$ 的变化方程，请重新定义一个广义的E-P通量 \mathcal{F}_m ，使得新的E-P通量中包含了eddy对水汽输送的作用；并且证明，在湿绝热 ($Q_m = 0$) 和无摩擦的情况下，平衡状态下的 \mathcal{F}_m 满足 $\nabla \cdot \mathcal{F}_m = 0$ ，并请根据水汽输送的空间分布讨论：在实际大气中，新定义的E-P通量的 $\nabla \cdot \mathcal{F}_m$ 应该有怎样的变化？eddy 对水汽的输送作用将对维持 Ferrel 环流起到怎样的作用？

4) 请根据新定义出的E-P通量，定义出新的剩余环流(residual circulation, $[\bar{v}_m]$, $[\bar{\omega}_m]$)，并讨论此时剩余环流的含义是什么？相对于新的剩余环流，新的TEM方程(Transformed Eulerian Mean Equations)应该是什么？同时，也请写出，如果用剩余环流来表述，(1)问中推导出的水汽方程将如何改写，eddy强迫项应变为什么？

作业用到的eddy对水汽输送的空间分布