



第五章:

大气环流中的纬向环流系统

5.3 Walker Circulation

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2022. 12. 08



Outline



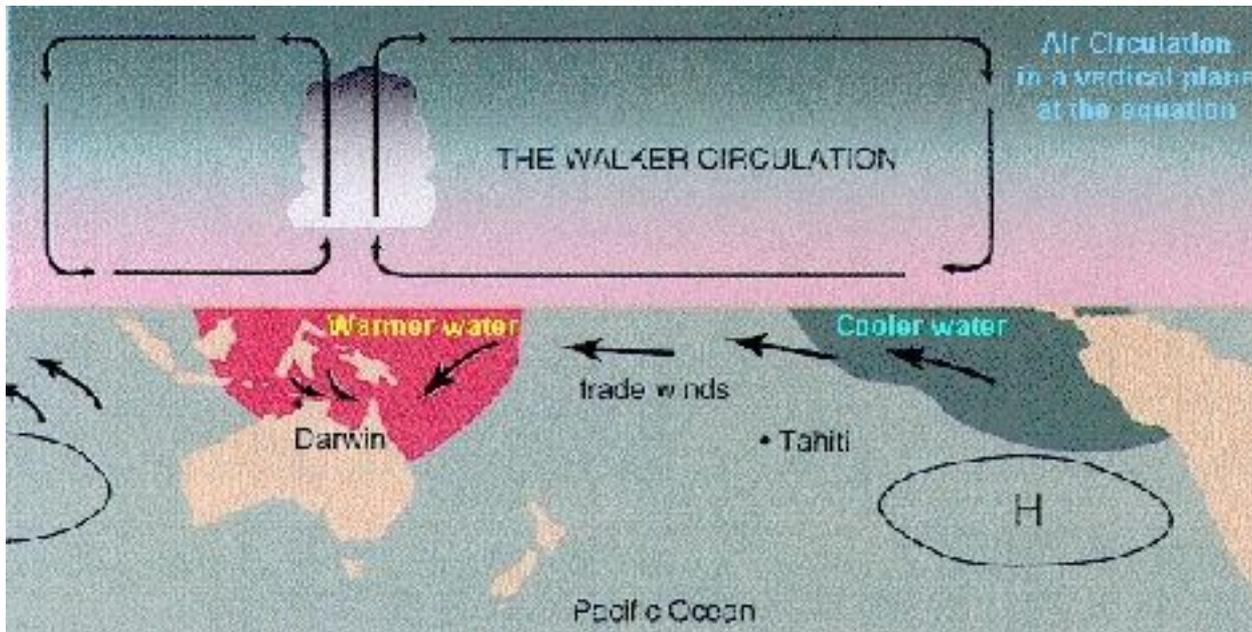
- Introduction
- Features of Walker circulation
- Seasonal variation of walker circulation
- ENSO and Walker circulation (inter-annual variation)
- The two-box model of Walker circulation
- Summary



Introduction



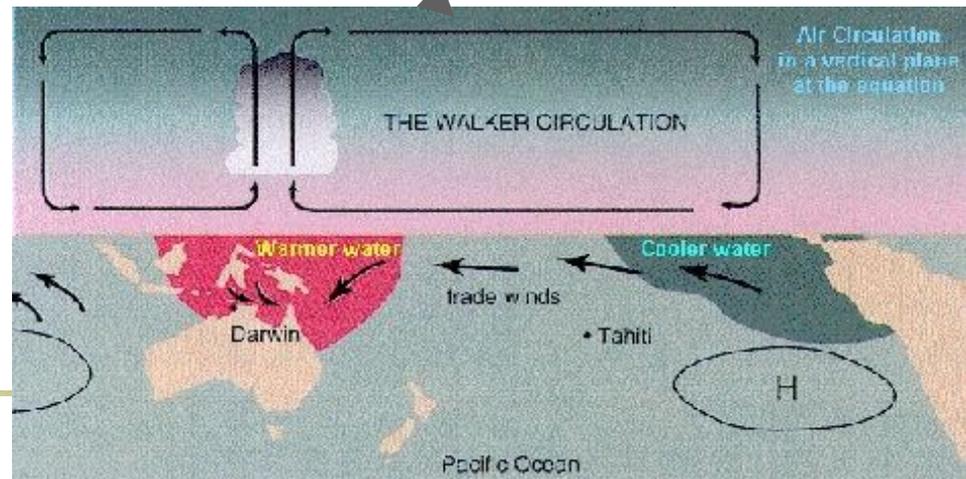
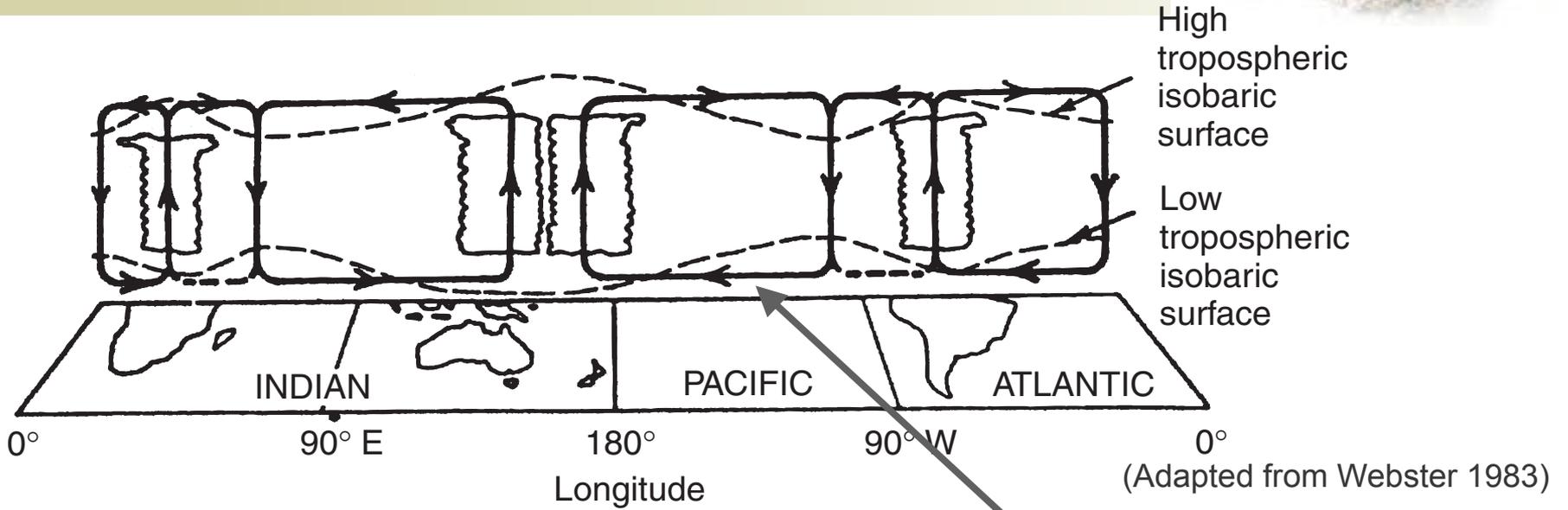
- **Walker Circulation** was first introduced in 1966 by Prof. Jacob Bjerknes, referring to the large scale atmospheric circulation along the **longitude-height plane** over the **equatorial Pacific Ocean**.



Named after Gilbert Walker, a key contributor to Southern Oscillation.



Introduction

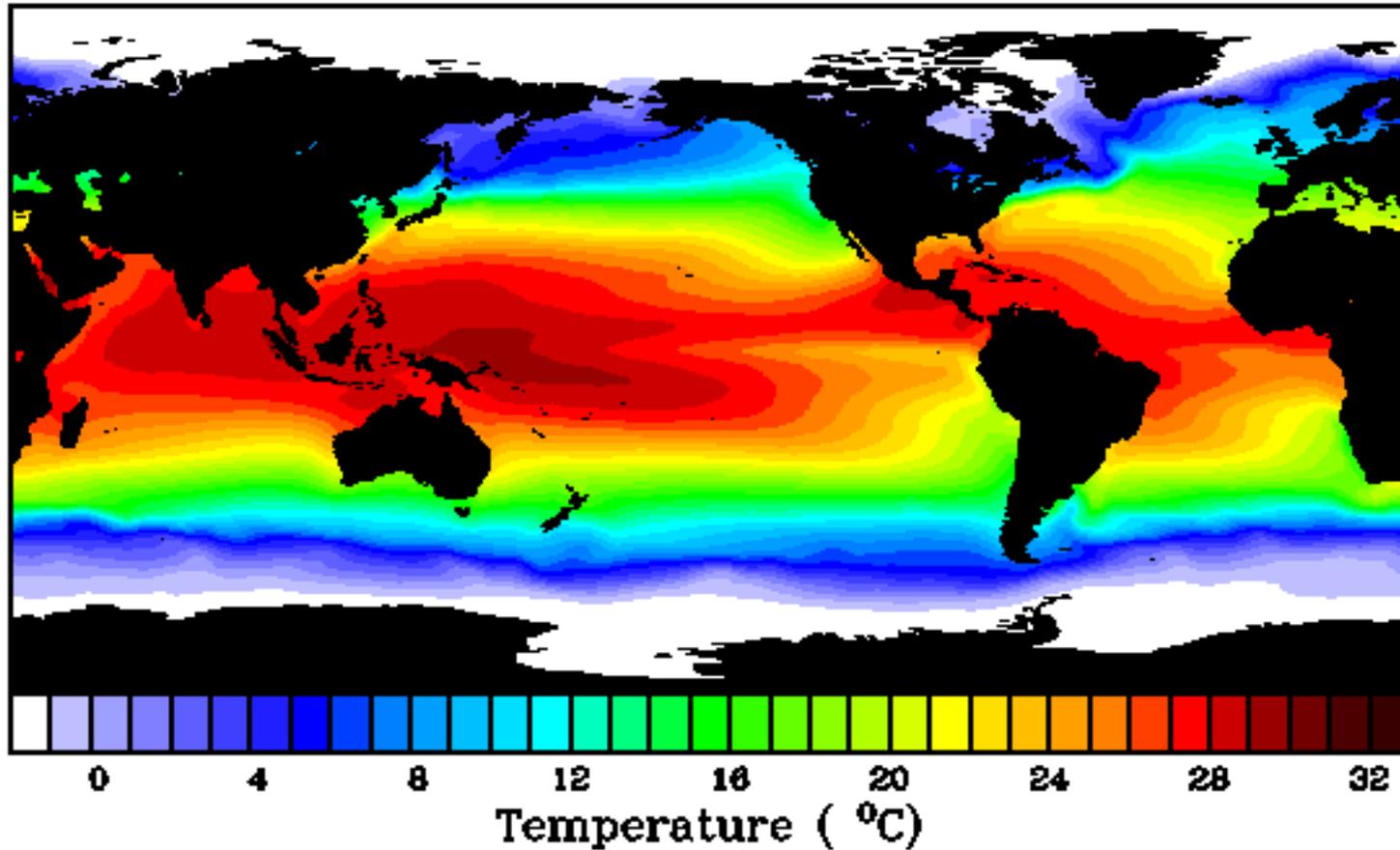




Observed features



ANNUAL MEAN
GLOBAL SEA SURFACE TEMPERATURES

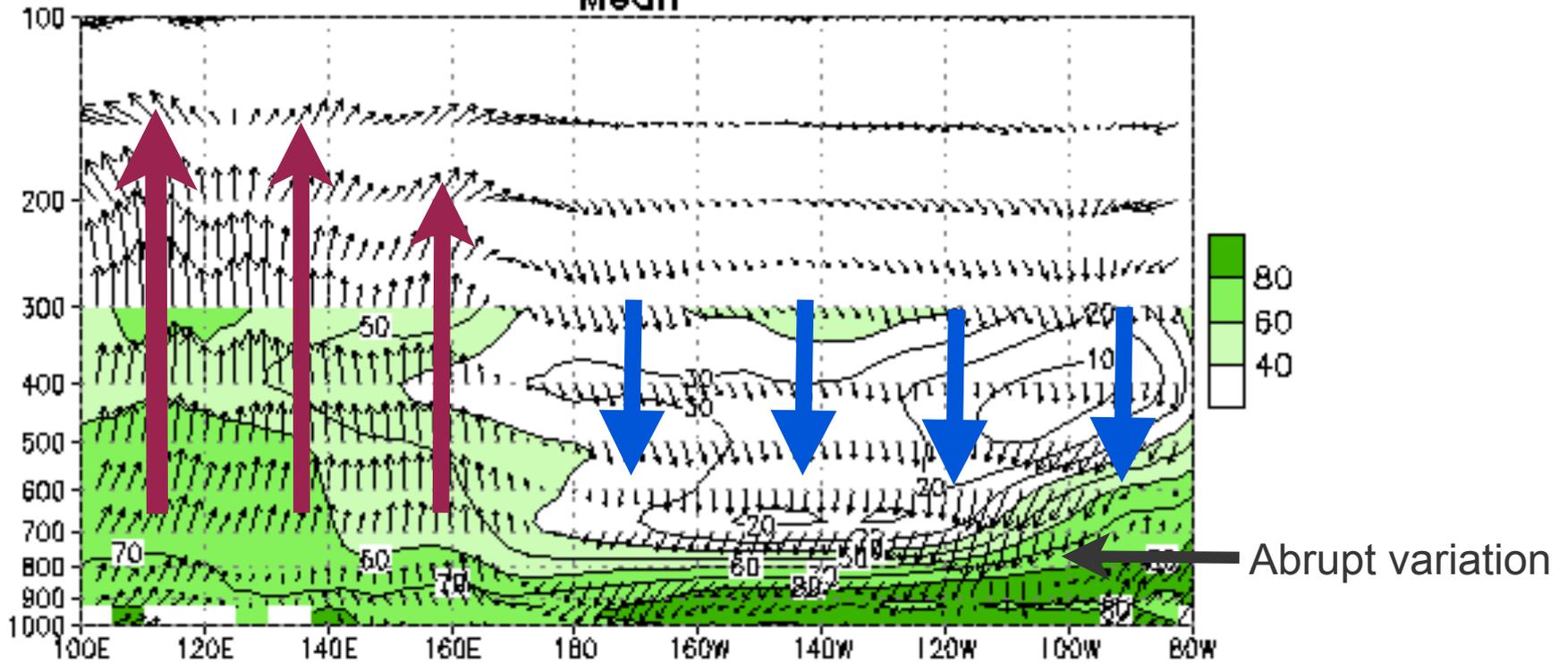




Observed features



RH & Combined uchi and w (5N-5S)
DEC 1999-FEB 2000
Mean



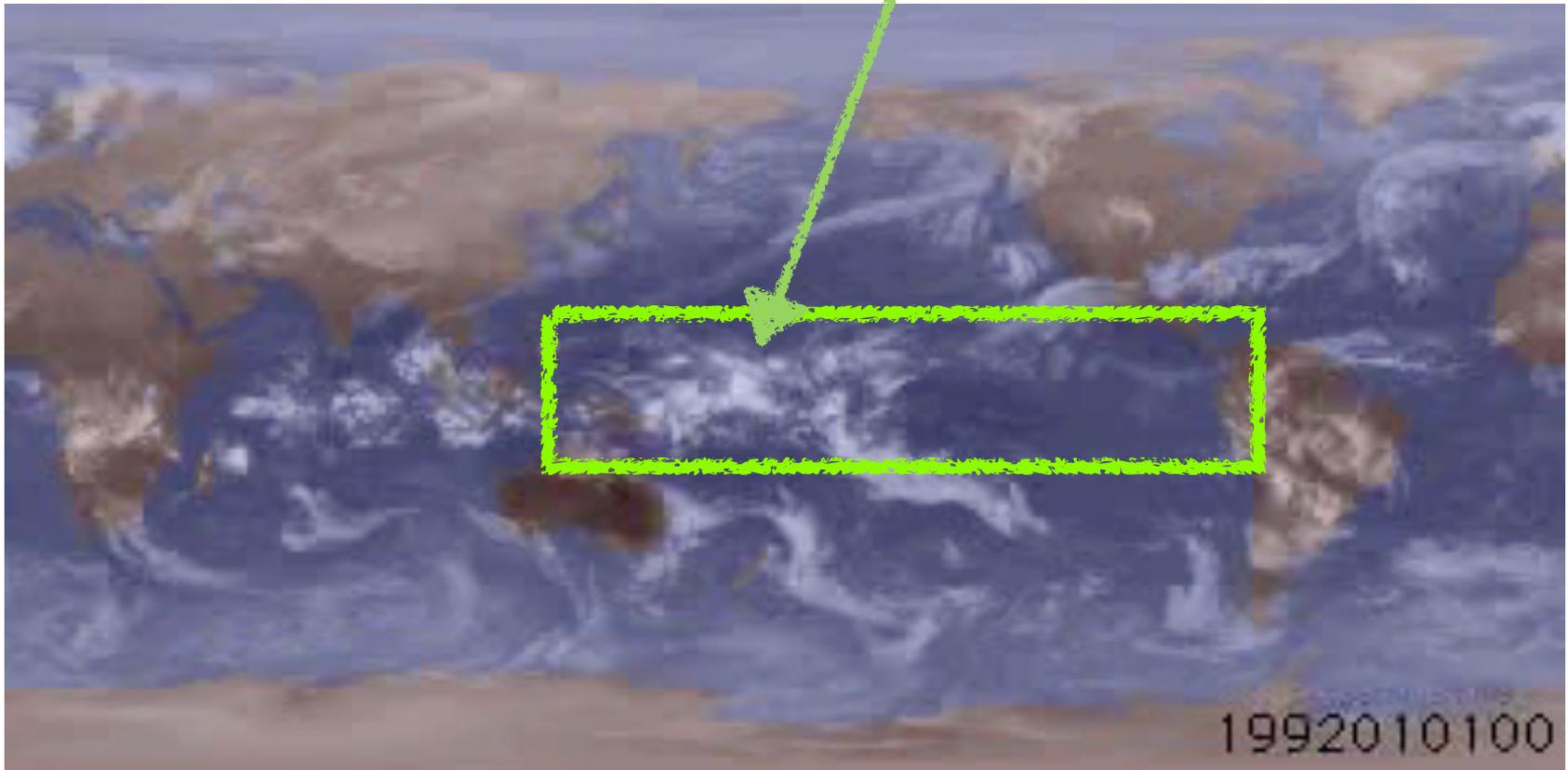
Adapted from Emanuel, 2005



Observed features



great amount of convective cloud in the western pacific

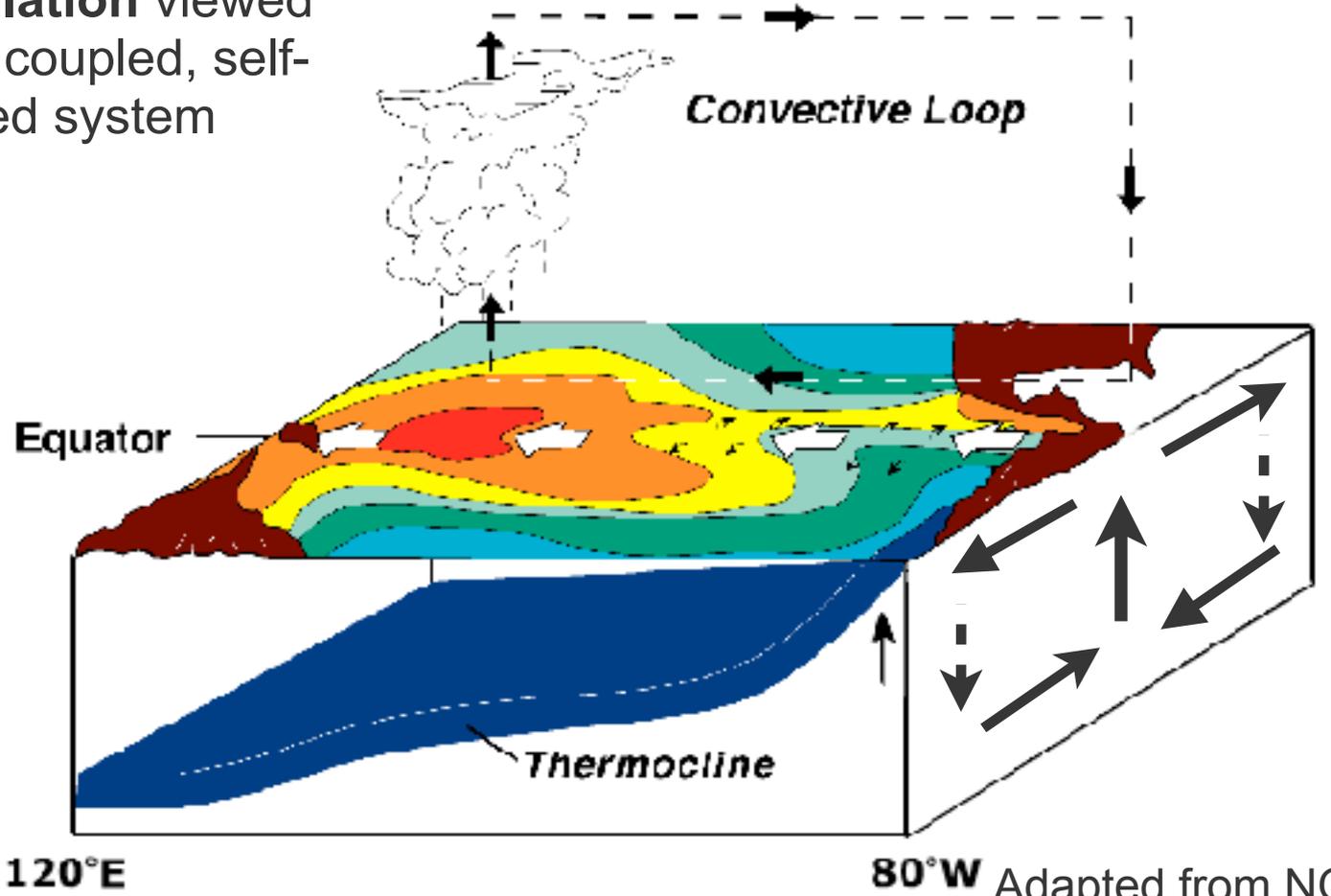




Observed features



Walker Circulation viewed as an air-sea coupled, self-maintained system

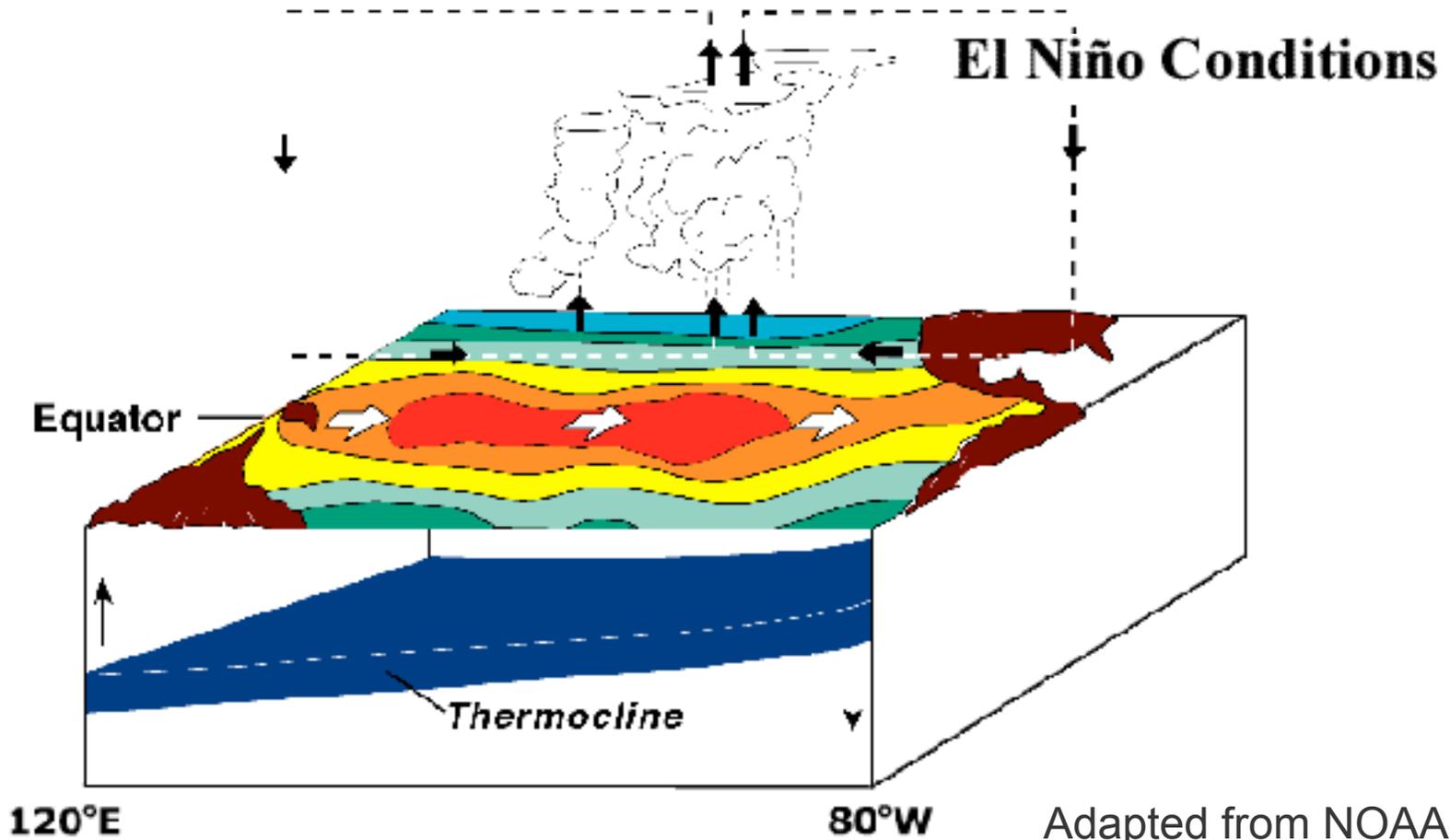




- **El Niño** is defined by prolonged differences in **tropical Pacific Ocean SST** when compared with the average value. The accepted definition is a warming or cooling of at least 0.5 °C averaged over the east-central tropical Pacific Ocean. Typically, this anomaly happens at irregular intervals of 2–7 years and lasts nine months to two years.
- The **Southern Oscillation** is the atmospheric component of **El Niño**. This component is an **oscillation in surface air pressure** between the tropical **eastern and the western Pacific Ocean waters**. The strength of the Southern Oscillation is measured by the **Southern Oscillation Index (SOI)**. The SOI is computed from fluctuations in the surface air pressure difference between Tahiti and Darwin, Australia.

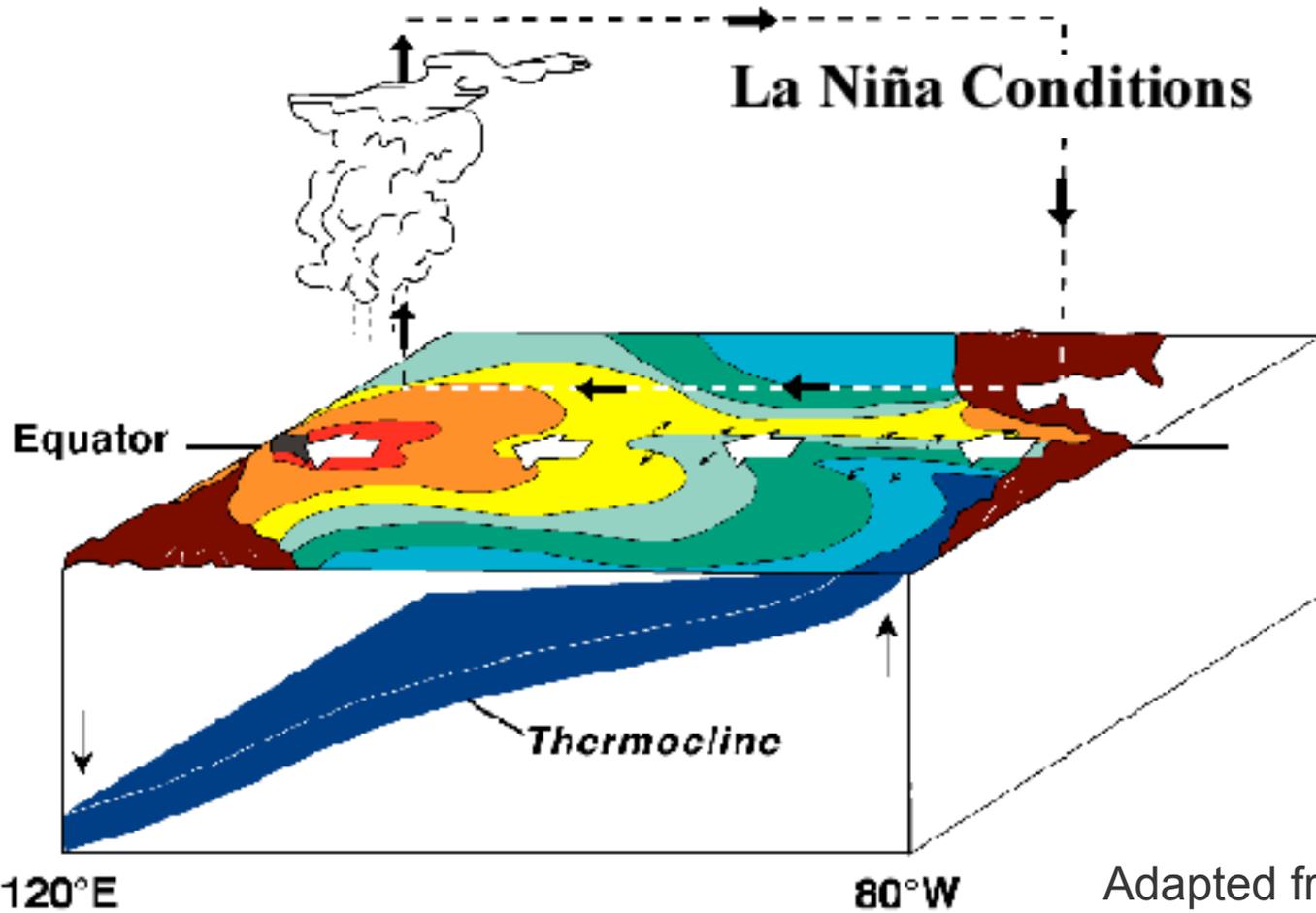


Walker Circulation





Walker Circulation

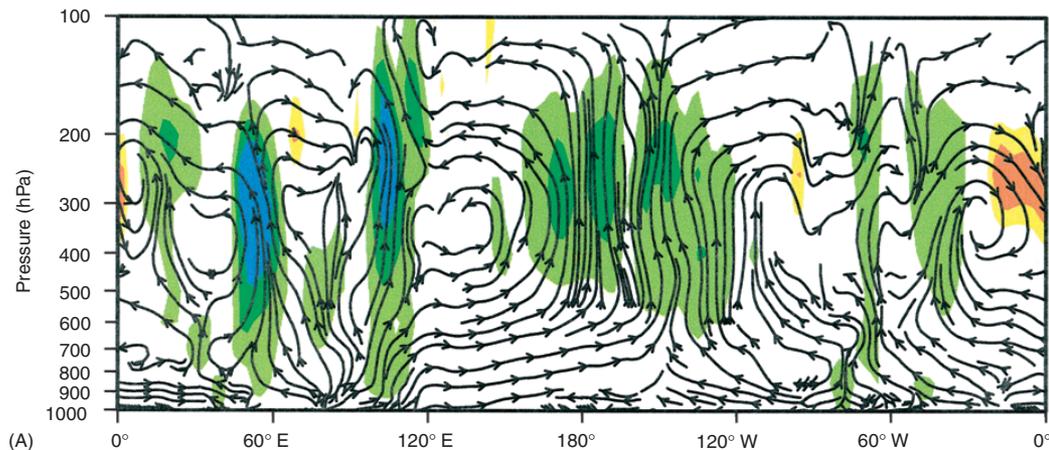


Adapted from NOAA



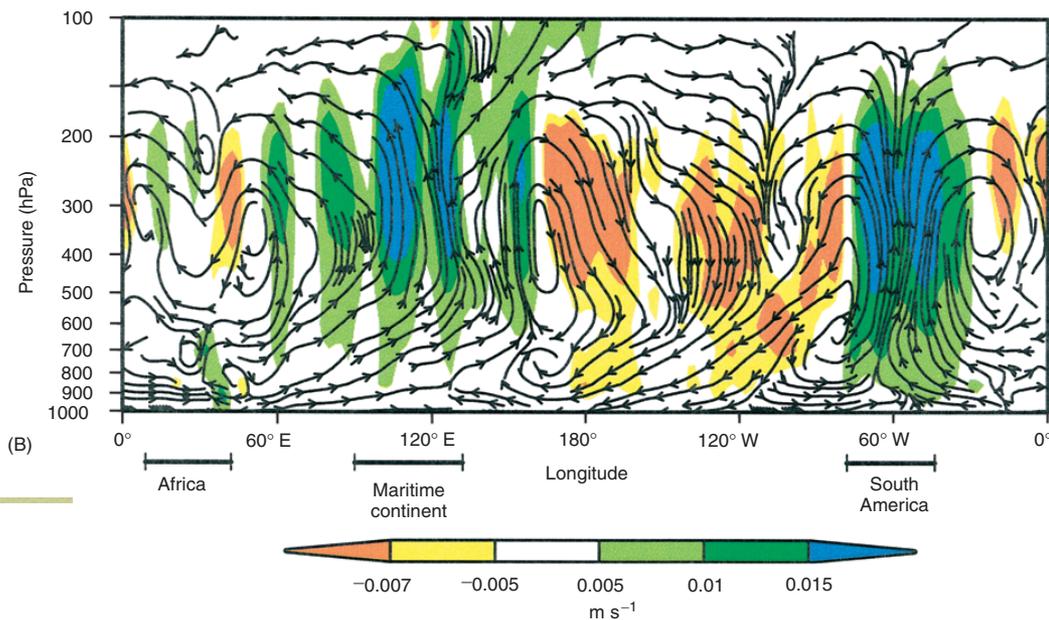
ENSO and

Walker Circulation



El Niño years:

Rising motion prevailed at almost all longitudes with a peak in central Pacific.



La Niña years:

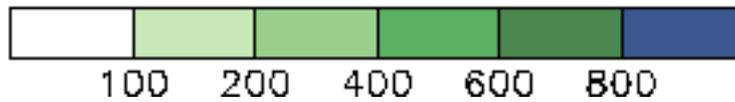
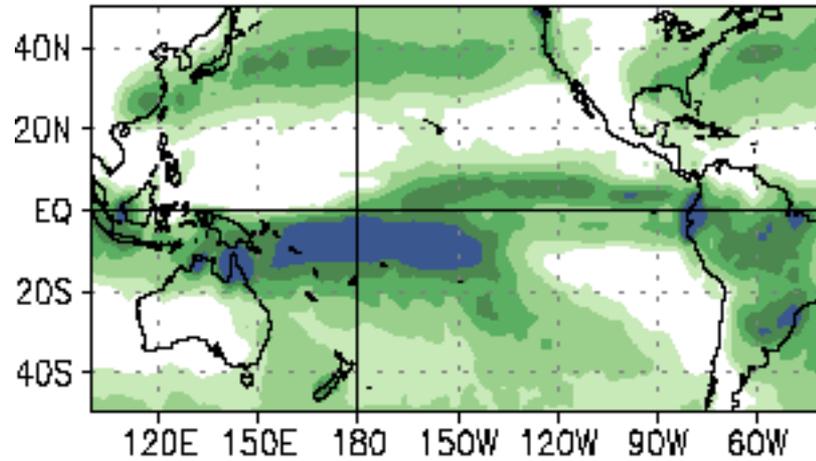
An enhanced Walker Circulation.

Adapted from Lau et al, 2002

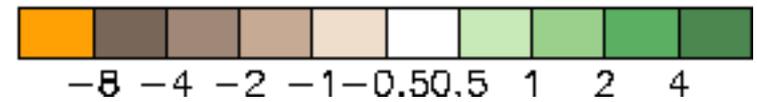
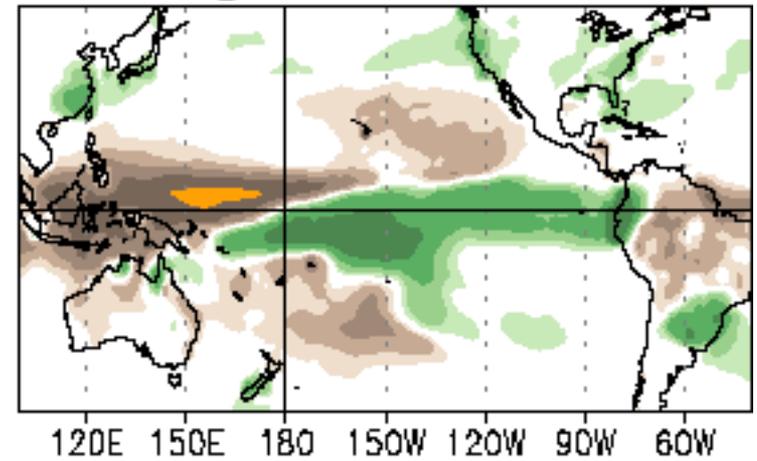


Jan-Mar 1998 Precipitation (mm)

Total



Departures (x100)



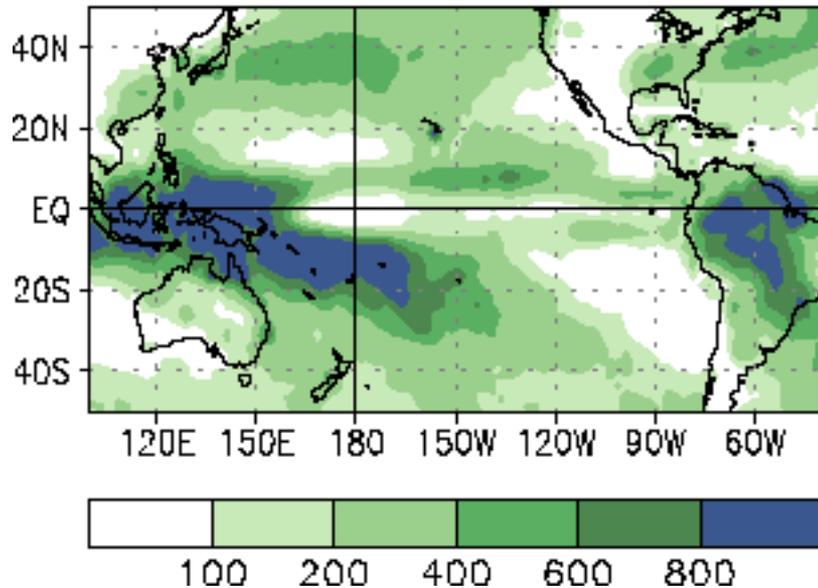
El Nino years

Adapted from NOAA

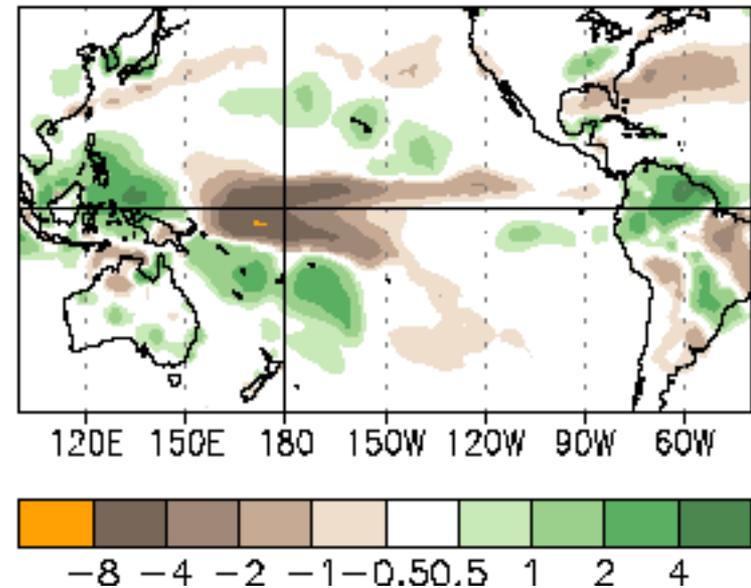


Jan-Mar 1989 Precipitation (mm)

Total



Departures (x100)



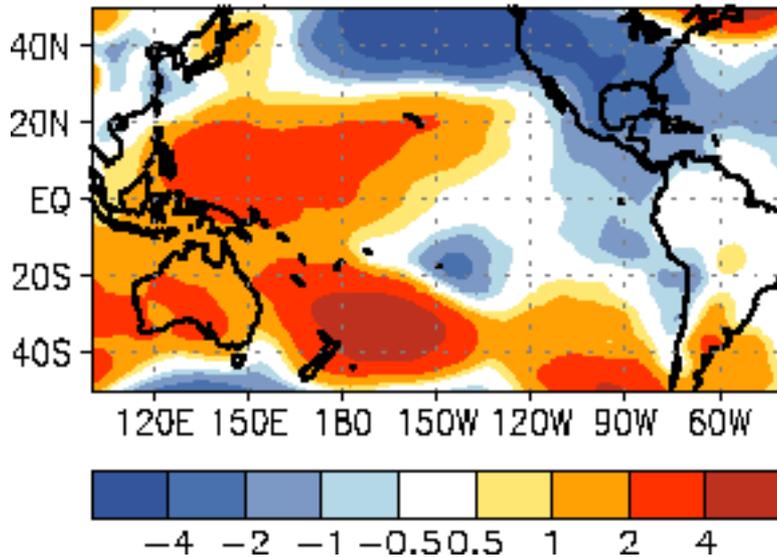
La Nina years

Adapted from NOAA

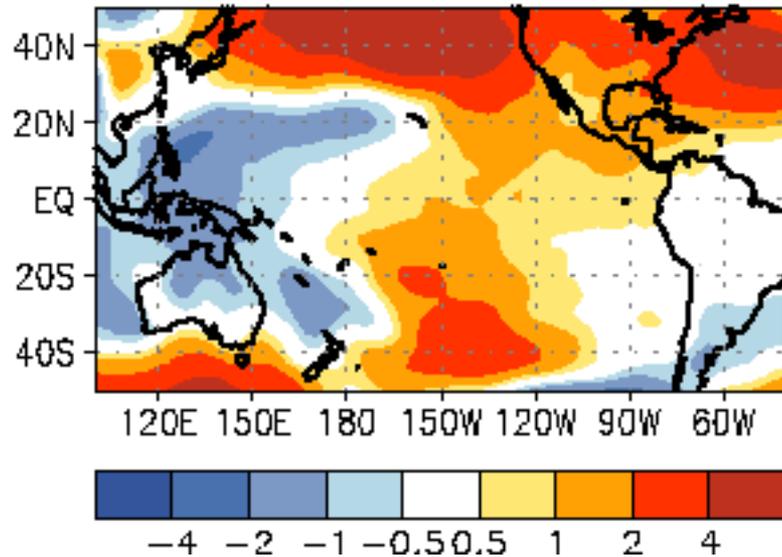


PRESSURE DEPARTURES (mb)

EL NIÑO
Jan-Mar 1998



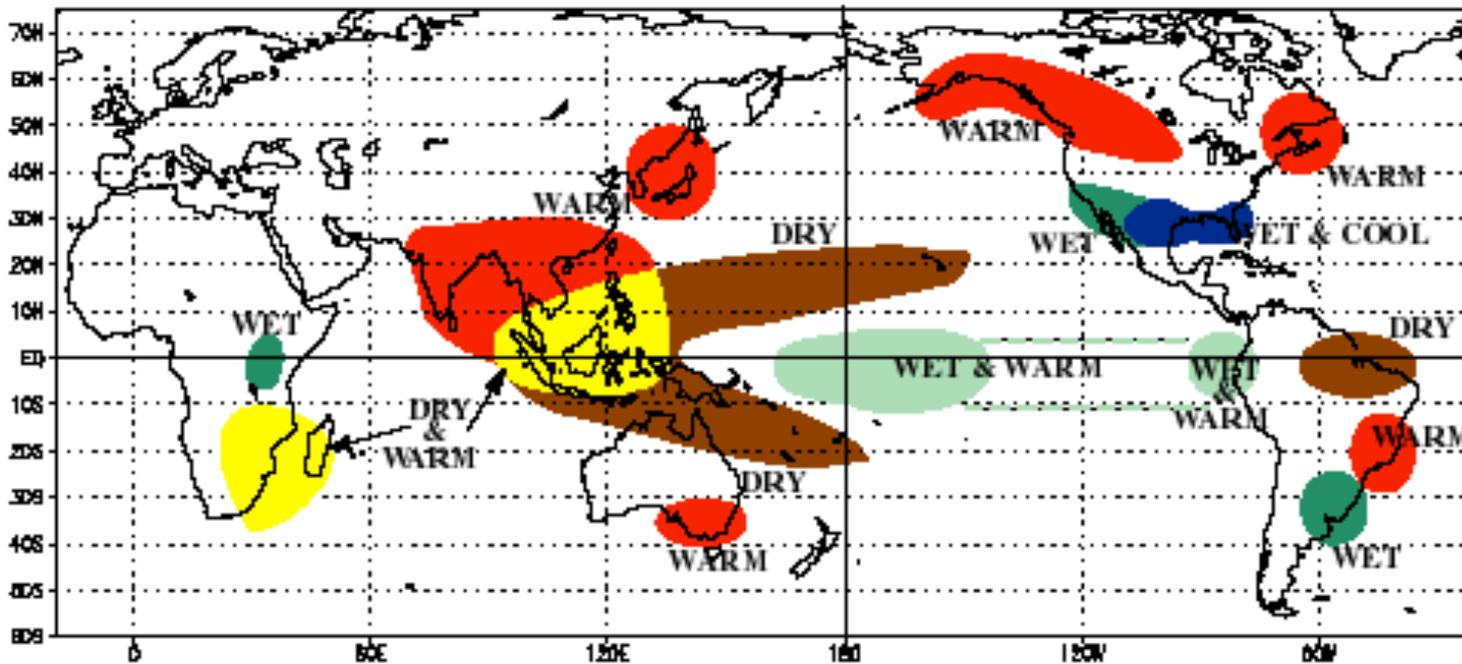
LA NIÑA
Jan-Mar 1989



Adapted from NOAA



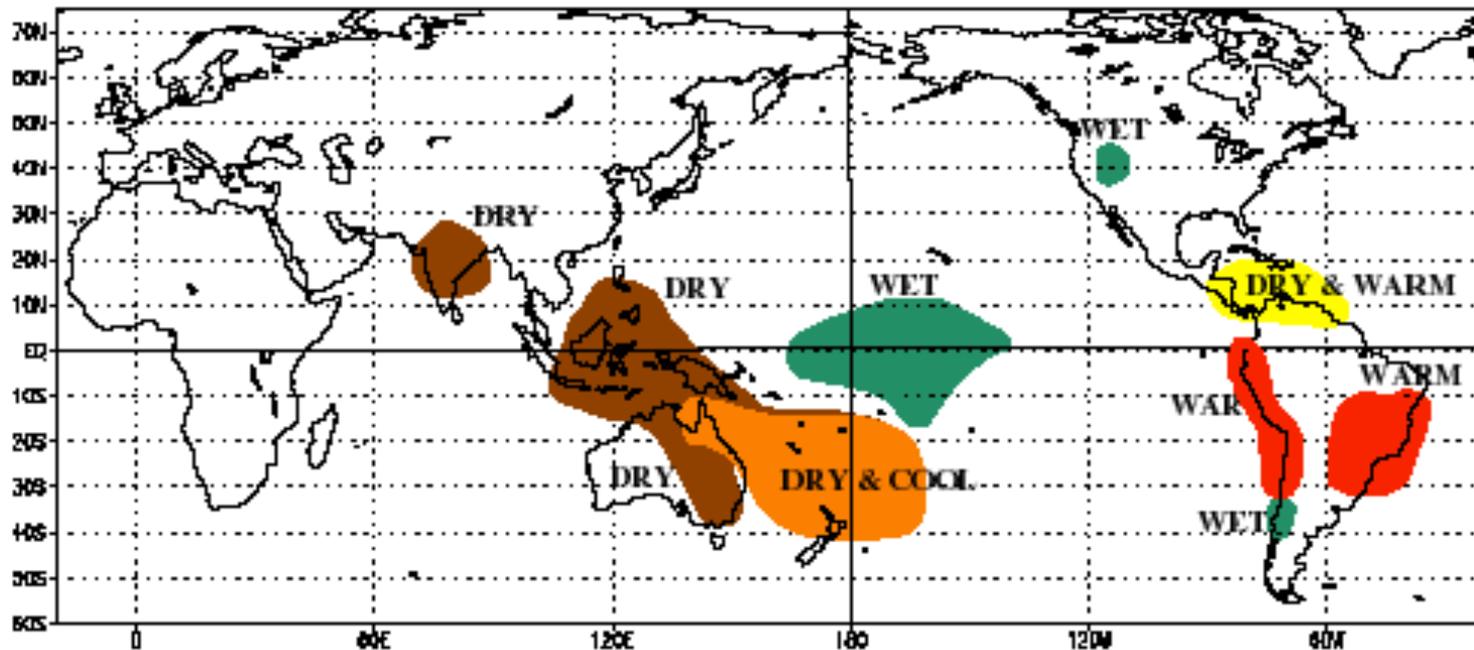
WARM EPISODE RELATIONSHIPS DECEMBER - FEBRUARY



Adapted from NOAA



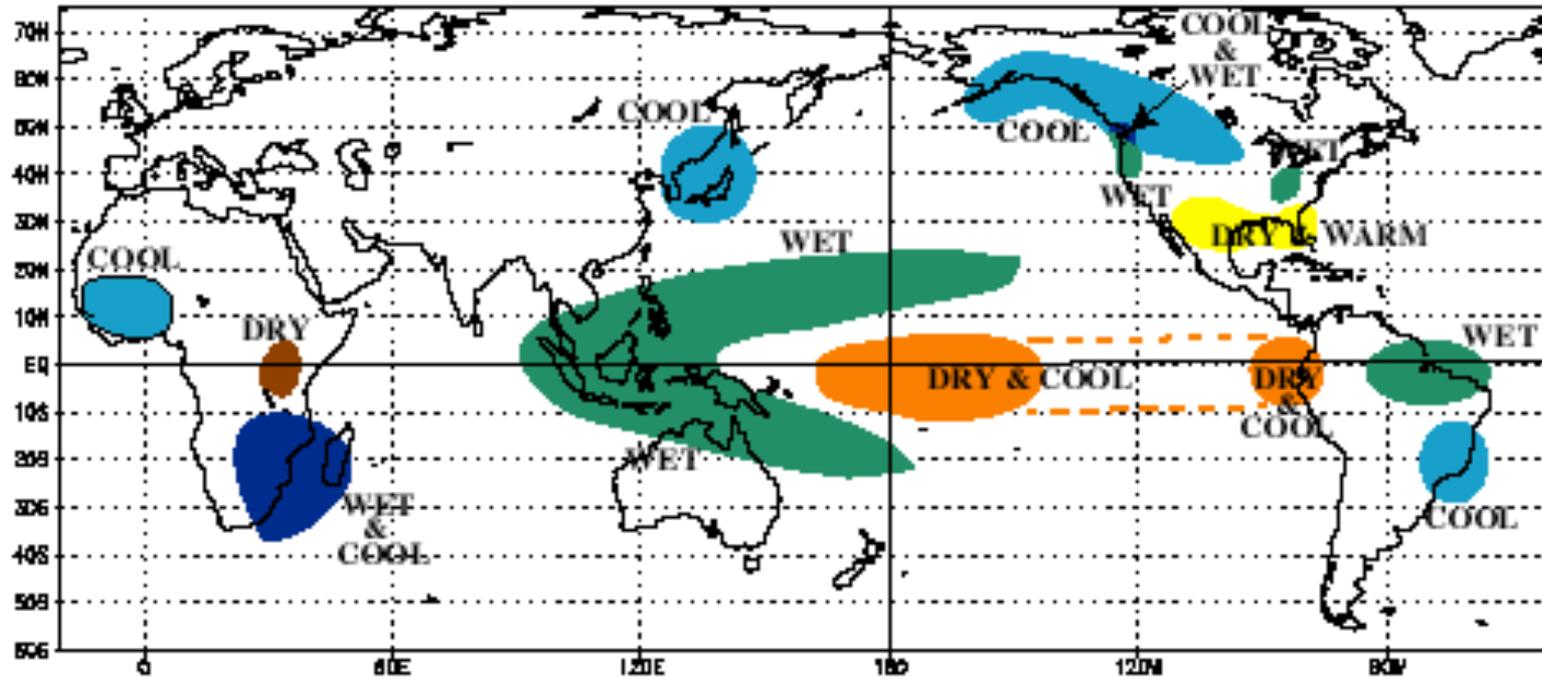
WARM EPISODE RELATIONSHIPS JUNE - AUGUST



Adapted from NOAA



COLD EPISODE RELATIONSHIPS DECEMBER - FEBRUARY



Adapted from NOAA

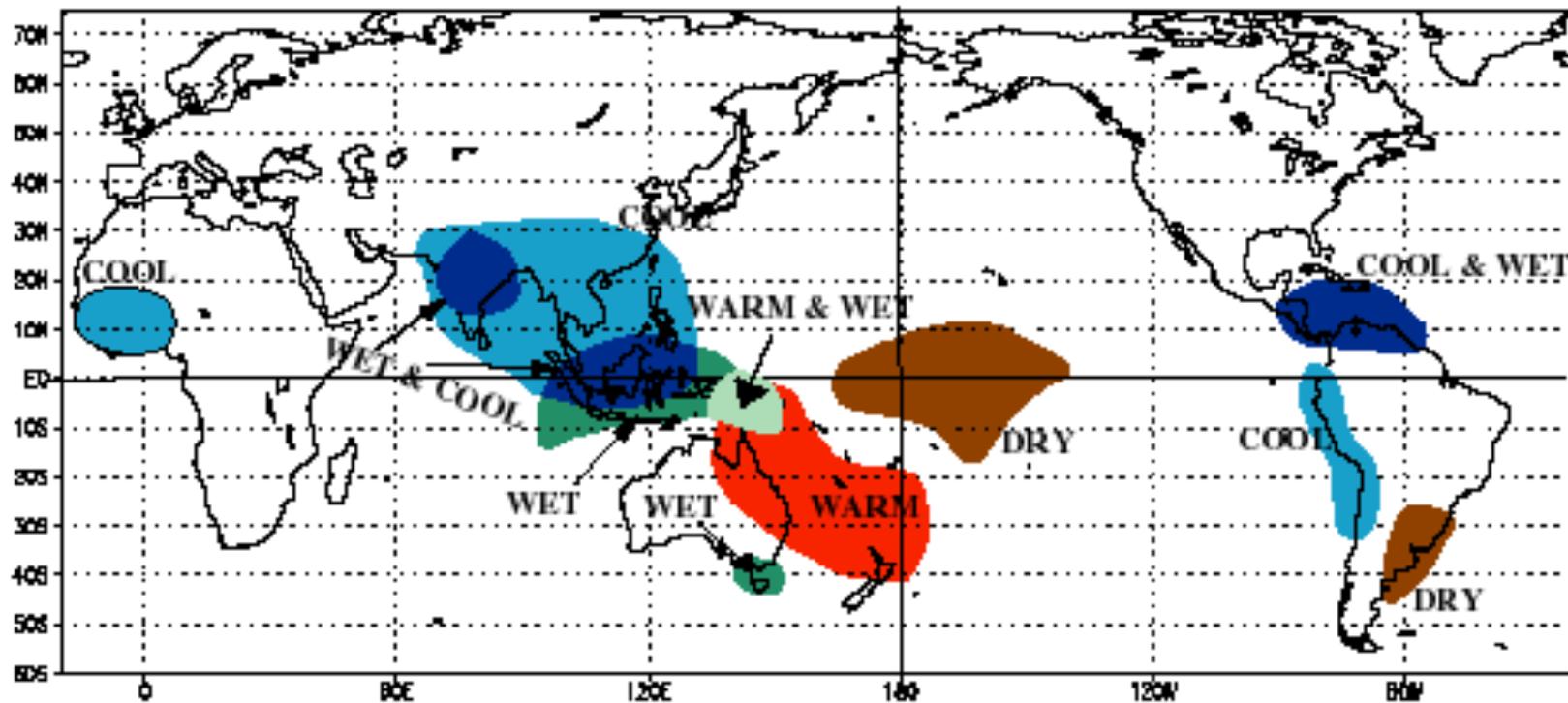


ENSO and

Walker Circulation



COLD EPISODE RELATIONSHIPS JUNE - AUGUST



Adapted from NOAA



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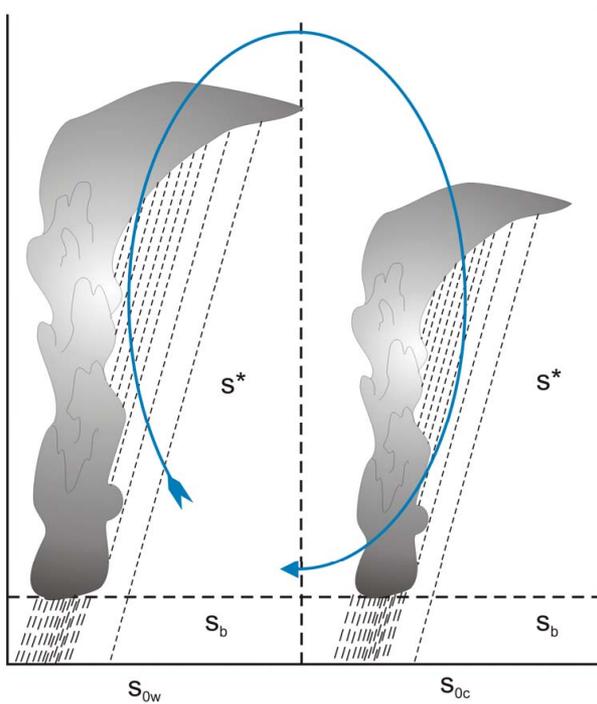
Two-box model

Walker Circulation

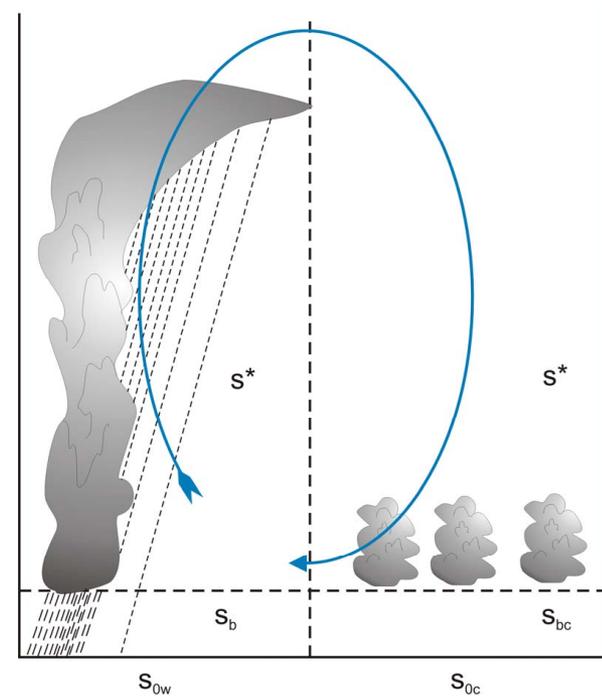


Interaction between convection and large-scale circulation

Two-Box Model:



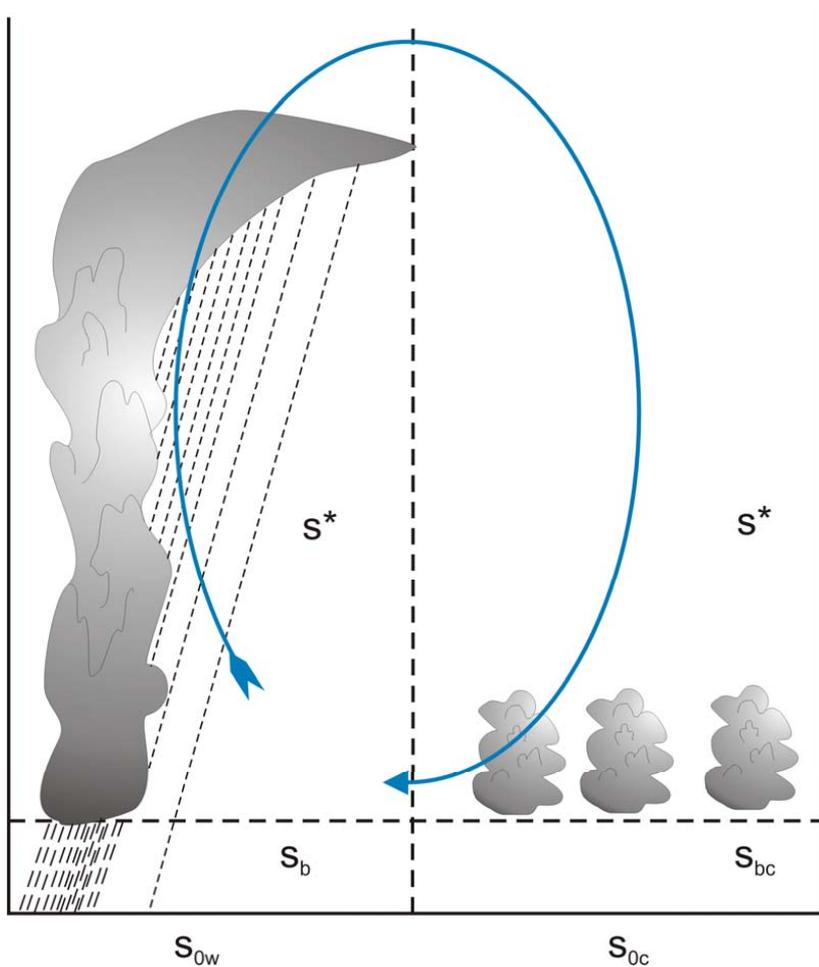
Weak circulation: Deep convection in both boxes



Strong circulation: Deep convection only in warm box



Walker Circulation



$$s \equiv c_p(T_s - \bar{T}) \ln(\theta_e^*)$$

θ_e^* -- saturation equivalent potential temperature

s_{0c} - entropy on the cold ocean surface

s_{bc} - entropy in the boundary layer over the cold ocean surface

s^* - entropy in the free troposphere, recall the weak temperature gradient in the tropics

Assume:

Quasi-equilibrium of the entropy in the boundary layer

Mass continuity

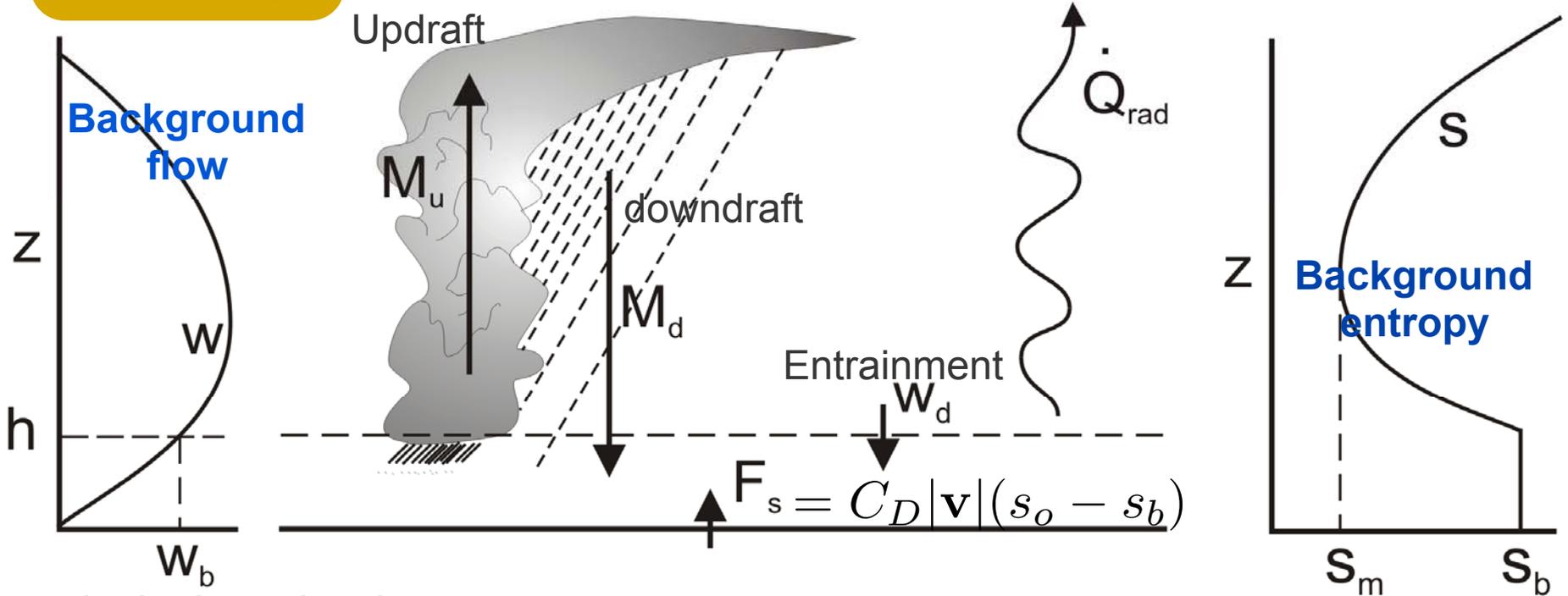


Two-box model

Walker Circulation



IN EACH BOX:



In the boundary layer of the convecting zone, assume equilibrium :

$$h \frac{\partial s_b}{\partial t} \approx 0 = F_s - (M_d + (1 - \sigma) w_d) (s_b - s_m)$$

Mass continuity:

$$M_u - M_d - (1 - \sigma) w_d = w_b = \gamma w$$

area covered by cloud

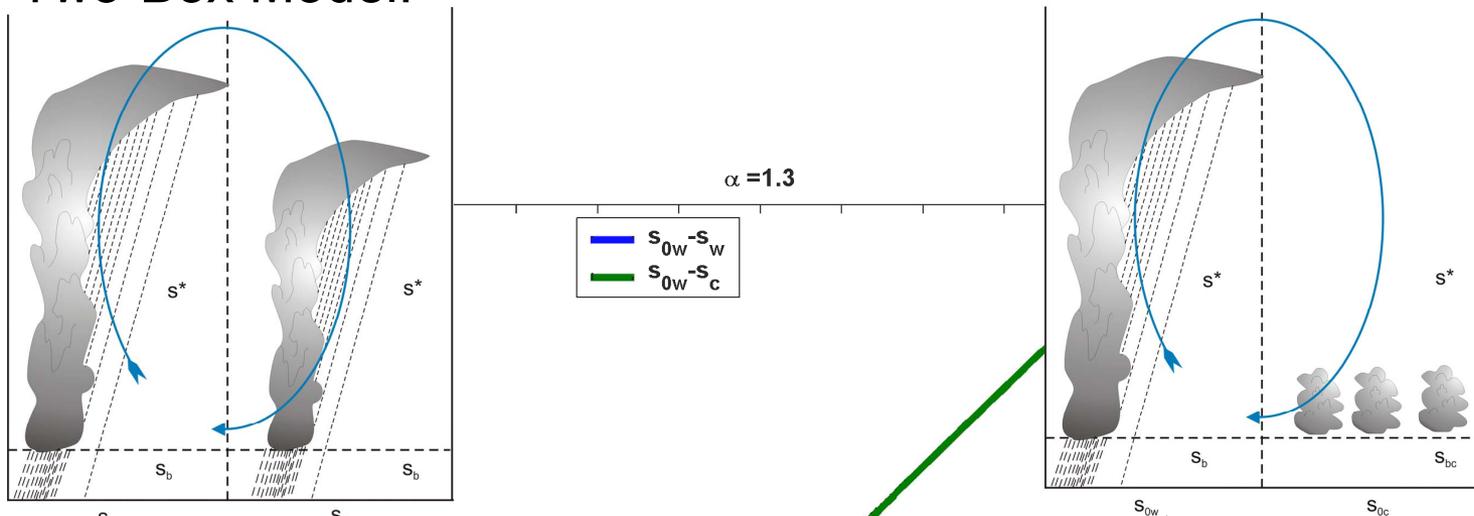


Two-box model

Walker Circulation

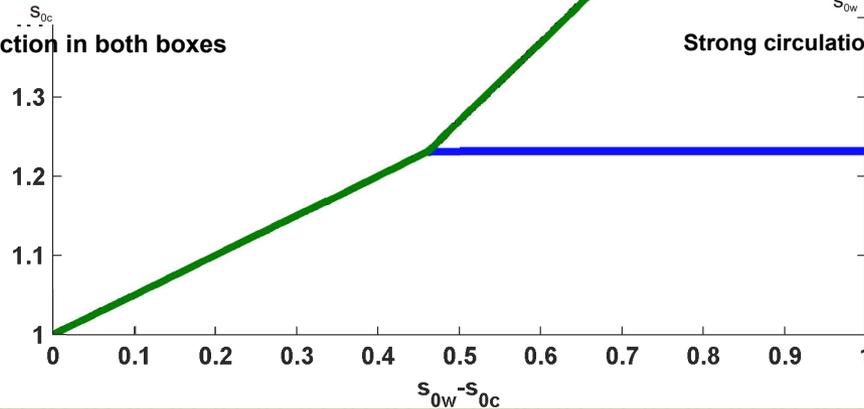


Two-Box Model:



Weak circulation: Deep convection in both boxes

Strong circulation: Deep convection only in warm box





Summary



- Walker circulation is the large-scale atmospheric overturning circulation along the **longitude-height plane** over the **equatorial Pacific Ocean**.
- Walker circulation is **a zonal temperature gradient driven, air-sea coupled phenomena**.
- Walker circulation shows **seasonal migration** in its location.
- Walker circulation is strongly affected by the **ENSO events**, with broad impact on the climate in the tropical and extratropical regions.
- **The two-box model**: the strength of walker circulation is determined by the **entropy difference** between the warm pool and the cold tongue, and the **precipitation efficiency** in the convective zone.